

An approach for modelling volume change of fine grained soil subjected to thermal cycles

Authors: Q. J. Ma*, C. W. W. Ng, D. Mašín and C. Zhou

*Corresponding author

Information of the authors

Corresponding author: Mr Q. J. Ma

Research student, Department of Civil and Environmental Engineering, Hong Kong University
of Science and Technology, Clear Water Bay, Kowloon, Hong Kong.

E-mail: qmaah@connect.ust.hk

Co-author: Dr C. W. W. Ng

Chair Professor, Department of Civil and Environmental Engineering, Hong Kong University of
Science and Technology, Clear Water Bay, Kowloon, Hong Kong.

E-mail: charles.ng@ust.hk

Co-author: Dr D. Mašín

Associate Professor, Faculty of Science, Charles University in Prague, Czech Republic.

E-mail: masin@natur.cuni.cz

17 **Co-author:** Dr C. Zhou

18 Visiting Assistant Professor, Department of Civil and Environmental Engineering, Hong Kong

19 University of Science and Technology, Clear Water Bay, Kowloon, Hong Kong.

20 E-mail: czhou@connect.ust.hk

21 **Abstract:** In consequence of cyclic thermal loads under drained condition, normally
22 consolidated and lightly over-consolidated fine grained soils experience accumulation of
23 irreversible volumetric contraction. Most of existing thermo-mechanical models were
24 developed for one heating-cooling cycle and not suitable for multiple thermal cycles. An
25 approach is proposed to simulate the volume change of fine grained soil induced by thermal
26 cycles. In the proposed approach a thermal stabilization line is introduced to control the
27 stabilized volumetric contraction under thermal cycles. Comparison with experimental results
28 shows that the proposed approach can reproduce well the cumulative feature of volumetric
29 contraction of fine-grained soil subjected to thermal cycles.

30 *Key words:* thermal, cyclic, fine grained soil, constitutive modelling.

31 **Introduction**

32 Thermal effects on soil behaviour have drawn attention from researchers throughout the past
33 decades (Campanella & Mitchell, 1968; Leroueil & Marques, 1996; Gens, 2010). Under drained
34 condition, monotonic temperature increase can induce irreversible volumetric contraction of
35 normally consolidated (NC) and lightly over-consolidated (OC) fine grained soils (e.g., Baldi et
36 al., 1988; Sultan et al., 2002; Cekerevac & Laloui, 2004; Abuel-Naga et al., 2007; Uchaipichat
37 & Khalili, 2009; Ng et al., 2016). When subjected to thermal cycles, they can experience
38 accumulation of irreversible volumetric contraction, which stabilizes after a few thermal cycles
39 (e.g., Campanella & Mitchell, 1968; Vega & McCartney, 2014; Di Donna & Laloui, 2015).

40 To model the thermally induced volume change of fine grained soil, the critical state
41 framework was extended incorporating the shrinkage of yield surface with increasing
42 temperature (e.g., Hueckel & Baldi, 1990; Graham et al., 2001; Laloui & Cekerevac, 2003). In
43 some other models (e.g., Cui et al., 2000; Abuel-Naga et al., 2007), an extra thermal yield
44 surface was introduced to improve the simulation of soil with relatively high over-
45 consolidation ratio (OCR). It should be noted that these models focus on one heating-cooling
46 cycle and cannot give the cumulative trend of irreversible volumetric contraction with thermal
47 cycles. For the thermo-mechanical models based on the concept of hypoplasticity (e.g., Mašín
48 & Khalili, 2012; Zhou & Ng, 2015), although they can capture the cumulative trend, they may

overestimate the accumulated irreversible volumetric contraction. Di Donna & Laloui (2015) furthered the work by Laloui & François (2009) to account for cyclic thermal loading through modifying the rule governing the plasticity mobilization during cooling.

The objective of this study is to propose a new approach for simulating the volume change of fine grained soil subjected to thermal cycles. The sign convention used herein is positive for compression and negative for tension.

Proposed approach

Schematic illustration

As suggested by the experimental results on soil volume change under thermal cycles (e.g., Campanella & Mitchell, 1968; Vega & McCartney, 2014; Di Donna & Laloui, 2015), a thermal stabilization line (TSL) is proposed, which controls the stabilized soil state under cyclic thermal loading. For simplicity, it is assumed to be a straight line in the $\ln v - \ln p'$ space as shown in Fig. 1, where the normal compression line (NCL) and the thermal stabilization line (TSL) correspond to a given temperature higher than the reference temperature. The point O represents the state of an NC soil specimen at the end of first heating, and the point O' represents the stabilized state after several thermal cycles. The distance OO' reflects the accumulated irreversible volumetric contraction of the NC soil specimen after the first heating. If it equals 0, there is no accumulation of irreversible volumetric contraction. It is assumed

that (1) As temperature changes the TSL shifts together with the NCL which is temperature dependent according to the experimental results (e.g., Campanella & Mitchell, 1968); (2) During heating, if the soil state is above the current TSL there is heating induced irreversible volumetric contraction; Otherwise soil response is thermo-elastic; (3) During cooling, soil response is thermo-elastic. Based on the assumptions, it can be deduced that if the distance OO' equals 0 and the slope of the TSL equals that of the NCL, the proposed approach is reduced to that proposed by Hueckel & Baldi (1990) and Laloui & Cekerevac (2003).

An NC soil specimen subjected to thermal cycles at constant effective stress is analysed to illustrate how the proposed approach works. Fig. 2 shows the state evolution of the NC soil specimen during thermal cycles. The open and solid circles represent the initial state and the current state of the soil specimen, respectively. The dashed and solid lines correspond to the initial temperature ($\Delta T = 0^\circ\text{C}$) and the higher temperature ($\Delta T > 0^\circ\text{C}$), respectively. During the first heating, both the NCL and the TSL shift downwards, and the soil state stays on the NCL. Fig. 2(a) shows the soil state after the first heating. During cooling it is assumed that soil response is thermo-elastic, and thus the soil state remains unchanged while the NCL and TSL return back to their initial positions. The corresponding soil volume change during the first heating-cooling cycle is qualitatively represented by the curve in Fig. 2(b), which shows a continuous irreversible volumetric contraction during heating and elastic thermal contraction during cooling.

Upon re-heating, initially the soil state is below the current TSL and the soil response is thermo-elastic according to the assumptions made previously. As the TSL continues to move downwards and crosses the soil state point, it attracts the soil state point to move downwards. Therefore, there occurs more irreversible volumetric contraction. Instead of staying exactly on the TSL, at the end of the second heating the soil state point may stay slightly above the current TSL as shown in Fig. 2(c). This is based on the consideration of the possibly viscous nature of soil behaviour under thermal loading (Leroueil & Marques, 1996). Compared to the state after the first heating (see Fig. 2(a)), the vertical distance $\Delta\epsilon_v$ between the current soil state point and current TSL is reduced. The degree of reduction is influenced by the rate of irreversible volumetric contraction development with respect to the rate of temperature increase. Qualitatively, the corresponding soil volume change during the second thermal cycle is demonstrated by the curve in Fig. 2(d). At the beginning of heating, soil response is thermo-elastic and after temperature reaches some value it turns to be thermo-plastic. It is easy to predict that after several thermal cycles the soil state eventually approaches the TSL and the vertical distance $\Delta\epsilon_v$ decreases to zero as shown in Fig. 2(e). Thus, during subsequent heating there is no irreversible volumetric contraction if the temperature does not exceed the history maximum value, and soil response turns to be thermo-elastic as shown in Fig. 2(f).

Mathematical formulation

104 The mathematical formulation for isotropic stress condition is presented first and then
 105 extended to anisotropic stress condition. According to Mašín & Khalili (2012), the temperature
 106 dependent NCLs can be expressed as

$$\ln (1 + e) = N(T) - \lambda(T) \cdot \ln (p'/p_r) \quad (1)$$

107 where e is the void ratio; p' is the mean effective stress; p_r is the reference pressure (1 kPa);
 108 $\lambda(T)$ and $N(T)$ are the temperature dependent slope and intercept of the NCL, respectively.
 109 They are assumed to follow

$$N(T) = N(T_r) + n_T \cdot \ln (T/T_r) \quad (2)$$

$$\lambda(T) = \lambda(T_r) + l_T \cdot \ln (T/T_r) \quad (3)$$

110 where n_T and l_T are model parameters controlling the shift and slope change of the NCL with
 111 temperature, respectively; T_r is the reference temperature. According to the experimental
 112 results (e.g., Campanella & Mitchell, 1968), λ is almost independent of temperature, and thus
 113 l_T can often be chosen as 0.

114 The newly introduced TSL is determined by its slope k_T and the position of point O'
 115 ($\ln p'_o, \ln (1 + e_{o'})$) as shown in Fig. 1. It is expressed by

$$\ln (1 + e) = \ln (1 + e_{o'}) - k_T \cdot \ln (p'/p'_o) \quad (4)$$

116 where p'_o is the pre-consolidation pressure; $\ln (1 + e_{o'})$ can be calculated from

$$\ln (1 + e_{o'}) = \ln (1 + e_o) + c_T \cdot n_T \cdot \ln (T/T_r) \quad (5)$$

117 where e_o and $e_{o'}$ represent the void ratio corresponding to point O and point O', respectively.

118 The newly introduced coefficient c_T determines the accumulated volumetric contraction from
 119 the second thermal cycle on. If it equals 0, there is no accumulation. The slope parameter k_T
 120 influences the simulated irreversible volumetric contraction of OC soil. As k_T decreases, the
 121 simulated irreversible volumetric contraction of OC soil increases because it is more likely for
 122 the TSL to cross the soil state point during heating.

123 To extend the mathematical formulation from isotropic condition to anisotropic condition, the
 124 expression of NCL for the general anisotropic condition needs to be adopted

$$\ln (1 + e) = N^*(T) - \lambda(T) \cdot \ln (p'/p_r) \quad (6)$$

125 with

$$N^*(T) = N(T) - (\lambda(T) - \kappa) \cdot \ln(p'_o/p^{SBS}) \quad (7)$$

126 where p^{SBS} is the mean effective stress obtained by radial mapping of the stress state onto
 127 the state boundary surface from stress origin corresponding to the current void ratio. For
 128 isotropic condition p'_o is equal to p^{SBS} and equation (6) reduces to equation (1).

129 **Implementation of the proposed approach**

130 **Implementation**

131 The proposed approach is combined with the hypoplastic framework. The basic hypoplastic
 132 model for fine grained soil developed by Mašín (2005) takes a nonlinear relationship between
 133 the Jaumann stress rate tensor $\dot{\boldsymbol{\sigma}}$ and the Euler strain rate tensor $\dot{\boldsymbol{\varepsilon}}$ as

$$\dot{\boldsymbol{\sigma}} = f_s(\mathcal{L} : \dot{\boldsymbol{\varepsilon}} + f_d \mathbf{N} \|\dot{\boldsymbol{\varepsilon}}\|) \quad (8)$$

134 where \mathcal{L} and \mathbf{N} are fourth-order and second-order constitutive tensors, respectively; f_s and
 135 f_d are scalar factors to consider the effects of stress level and void ratio, respectively; $:$ stands
 136 for double contraction and $\|\mathbf{X}\|$ denotes the Euclidean norm of the tensor \mathbf{X} .

137 To model thermally-induced volume change of soil, Mašín & Khalili (2012) introduced a
 138 thermal term $f_u \mathbf{H}_T$ into equation (8) as

$$\dot{\boldsymbol{\sigma}} = f_s[\mathcal{L} : (\dot{\boldsymbol{\varepsilon}} - \dot{\boldsymbol{\varepsilon}}^{\text{TE}}) + f_d \mathbf{N} \|\dot{\boldsymbol{\varepsilon}} - \dot{\boldsymbol{\varepsilon}}^{\text{TE}}\|] + f_u \mathbf{H}_T \quad (9)$$

139 where $\dot{\boldsymbol{\varepsilon}}^{\text{TE}}$ is the isotropic thermal elastic strain rate tensor. Mathematical expression for \mathbf{H}_T
 140 was derived by considering that when subjected to heating at constant effective stress, the
 141 NC soil stays on the current NCL as it moves with temperature change. Detailed mathematical
 142 expressions and discussions of the terms involved can be found in Mašín (2005) and Mašín &
 143 Khalili (2012).

144 The collapse potential factor $0 \leq f_u \leq 1$ controls the heating induced irreversible volumetric
 145 contraction. The larger the collapse potential factor f_u the more heating induced irreversible
 146 volumetric contraction is. For an NC soil specimen during the first heating f_u equals 1. When

147 f_u equals 0 it implies soil response is thermo-elastic.

148 To implement the proposed approach within the hypoplastic framework, the collapse
149 potential factor f_u is modified as

$$f_u = \left\langle \frac{e - e_T^*}{e_T - e_T^*} \right\rangle^\gamma \quad (10)$$

150 where $\langle x \rangle$ is an operator obtaining the positive part of the scalar variable x , $\langle x \rangle =$
151 $(x + |x|)/2$; e is the void ratio; e_T and e_T^* are the void ratios on the current NCL and TSL
152 corresponding to current mean effective stress, respectively (see point C in Fig. 1). They can
153 be calculated from equations (6) and (4), respectively. γ is a new parameter controlling the
154 rate of irreversible volumetric contraction development with respect to the heating rate. As a
155 consequence, it controls the number of thermal cycles required to get the soil volumetric
156 contraction stabilized.

157 **Calibration of model parameters**

158 In this section, calibration of the three newly introduced parameters c_T , k_T and γ is discussed.
159 For the calibration of other relevant model parameters, please refer to Mašín (2005) and
160 Mašín & Khalili (2012). To calibrate the parameter c_T , volume change test of an NC soil
161 specimen subjected to thermal cycles until stabilization is required. Regarding the slope
162 parameter k_T , test results of soil specimens with different OCRs are necessary. It can be

determined based on the threshold value of OCR, corresponding to which there is no heating induced irreversible volumetric contraction for a given temperature increase. The crossing point of the thermal stabilization line for a given temperature and the unloading line corresponds to the threshold OCR. Alternatively, it can be determined by trial and error to achieve an overall reasonable simulation of OC soil.

A sensitivity analysis was conducted to study the effect of the parameter γ on the accumulation of irreversible volumetric contraction of an NC soil specimen with thermal cycles. Model parameters used are these for the soil tested by Uchaipichat & Khalili (2009) (see Table 1). Obtained results are shown in Fig. 3, which indicates that as γ decreases the soil volumetric contraction stabilizes within less number of thermal cycles. Specifically, for $\gamma = 0.1$ it stabilizes around the fifth thermal cycle, which is consistent with the experimental results (e.g., Vega & McCartney, 2014). Based on the sensitivity analysis, a default value of 0.1 is suggested for γ , which will be further validated against experimental results later. It should be noted that the parameter γ does not affect the irreversible volumetric contraction corresponding to the first thermal cycle. For an NC soil specimen during the first heating, the soil state remains on the current NCL, and thus its volumetric contraction is completely determined by the parameter n_T controlling the shift of NCL with temperature increase (see equation (2)).

Typical results

Fig. 4 presents the computed volume change of soil specimens with different OCRs (1, 1.3, 2 and 4) subjected to 15 thermal cycles (25 – 60 °C) using model parameters for the soil tested by Uchaipichat & Khalili (2009) (see Table 1). It can be seen that as the OCR increases, both the irreversible volumetric contraction after the first thermal cycle and that corresponding to the stabilized state decrease. The irreversible volumetric contraction stabilizes within roughly five thermal cycles for all the soil specimens except that with OCR of 4. For the soil specimen with OCR of 4, there is no irreversible volumetric contraction, which indicates a thermo-elastic soil response. It can be predicted that for soil specimen with even higher OCRs, the simulated soil response is also going to be thermo-elastic. These trends are in good agreement with experimental results of saturated silt with different OCRs from Vega & McCartney (2014).

Evaluation of the proposed approach

Isotropic condition

By using the proposed approach, experimental test of remoulded illite from Campanella & Mitchell (1968) was simulated. The soil specimen was consolidated under isotropic condition to 200 kPa. Then three heating-cooling cycles (from about 60 °C to 5 °C) were applied under drained condition with constant mean effective stress. The initial temperature of the soil specimen was around 20 °C. Adopted model parameters are summarized in Table 1.

Comparison of the experimental results and the computed results is shown in Fig. 5. Overall,

it shows a reasonably good correlation between them. The irreversible volumetric contraction accumulates at a decreasing rate. The temperature at which irreversible volumetric contraction occurs increases cycle after cycle. However, for the cooling and re-heating phase there are some discrepancies. One possible reason is the temperature dependence of the thermal expansion coefficient α_s of soil skeleton (Laloui & Cekerevac, 2003), which is assumed to be a constant in this study for simplicity. In addition, the thermo-elastic assumption corresponding to the cooling and initial re-heating process can also result in some inaccuracies. This assumption is adopted also for the convenience of modelling, and can be dropped in future if necessary.

1D condition

One oedometer test from Di Donna & Laloui (2015) on natural silty-clay subjected to thermal cycles was also simulated. The in-situ condition was considered to be normally consolidated with the vertical effective stress around 149 kPa. Test procedure was that the soil specimen was loaded under 1D condition to a vertical effective stress of 125 kPa. Then three thermal cycles (from 60 °C to 5 °C) were applied under drained condition with the initial temperature around 20 °C. The vertical effective stress remained constant during the thermal loading. Model parameters used are summarized in Table 1.

Fig. 6 compares the experimental results to the computed results. It can be seen that the

proposed approach can simulate the main features of volume change behaviour under thermal cycles reasonably well. It should be noted that at the beginning of the first heating phase the experimental results show a continuous volumetric contraction compared to the computed results, which show slight volumetric expansion followed by large volumetric contraction. This computed initial volumetric expansion is due to the adopted assumption of elastic response when the soil state is below the TSL. According to some other literatures (e.g., Towhata et al., 1993; Sultan et al., 2002; Abuel-Naga et al., 2007), lightly over-consolidated soils experience initial volumetric expansion and then volumetric contraction after temperature exceeds some threshold value, which is consistent with the computed results. If the continuous volumetric contraction for lightly over-consolidated soil is intrinsic, which may depend on the soil type, the proposed approach can result in some discrepancies for the initial phase of first heating.

Summary

Based on the experimental results, a thermal stabilization line is introduced in the $\ln v - \ln p'$ space, which controls the stabilized soil state under cyclic thermal loading. Two parameters are needed to characterize the thermal stabilization line. One determines the accumulated irreversible volumetric contraction for an NC soil specimen, and the other controls the simulated results of OC soil. By taking use of the introduced thermal stabilization line, a

method is proposed to model the volume change of fine grained soil subjected to thermal cycles. The proposed method is realized within the hypoplastic framework and tested against experimental results. The comparison shows that measured and computed results are fairly consistent. The proposed method is able to simulate the overall trend of accumulation and stabilization of irreversible volumetric contraction with thermal cycles.

Acknowledgements

The first author greatly appreciates the HKPFS scholarship offered by the Research Grants Council (RGC) of the HKSAR. The financial support provided by the RGC of the HKSAR (grant no. GRF 617213 and 16209415), the HKUST (grant no. FP204) and the National Science Foundation of China (grant no. 51378178) are also gratefully acknowledged.

References

Abuel-Naga, H. M., Bergado, D. T., Bouazza, A. & Ramana, G. V. 2007. Volume change behaviour of saturated clays under drained heating conditions: experimental results and constitutive modelling. *Canadian Geotechnical Journal*, **44**(8): 942–956. doi: 10.1139/t07-031.

Baldi, G., Hueckel, T. & Pellegrini, R. 1988. Thermal volume changes of the mineral-water system in low-porosity clay soils. *Canadian Geotechnical Journal*, **25**(4): 807–825. doi: 10.1139/t88-089.

253 Campanella, R. G. & Mitchell, J. K. 1968. Influence of temperature variations on soil behavior.
 254 Journal of the Soil Mechanics and Foundations Division, *ASCE* **94**(3): 709–734.

255 Cekerevac, C. & Laloui, L. 2004. Experimental study of thermal effects on the mechanical
 256 behaviour of a clay. International journal for numerical and analytical methods in
 257 geomechanics, **28**(3): 209-228. doi: 10.1002/nag.332.

258 Cui, Y. J., Sultan, N. & Delage, P. 2000. A thermomechanical model for saturated clays.
 259 Canadian Geotechnical Journal, **37**(3): 607–620. doi: 10.1139/t99-111.

260 Di Donna, A. & Laloui, L. 2015. Response of soil subjected to thermal cyclic loading:
 261 experimental and constitutive study. Engineering Geology, **190**: 65–76. doi:
 262 10.1016/j.enggeo.2015.03.003.

263 Gens, A. 2010. Soil–environment interactions in geotechnical engineering. Géotechnique,
 264 **60**(1): 3–74. doi: 10.1680/geot.9.P.109.

265 Graham, J., Tanaka, N., Crilly, T. & Alfaro, M. 2001. Modified Cam-Clay modelling of
 266 temperature effects in clays. Canadian Geotechnical Journal, **38**(3): 608–621. doi:
 267 10.1139/t00-125.

268 Hueckel, T. & Baldi, G. 1990. Thermoplasticity of Saturated Clays: Experimental Constitutive
 269 Study. Journal of Geotechnical Engineering, **116**(12): 1778–1796. doi:

270 10.1061/(ASCE)0733-9410(1990)116:12(1778).

271 Towhata, I., Kuntiwattanakul, P., Seko, I. and Ohishi, K. 1993. Volume change of clays induced
 272 by heating as observed in consolidation tests. *Soils and Foundations*, **33**(4): 170–183. doi:
 273 10.3208/sandf1972.33.4_170.

274 Laloui, L. & Cekerevac, C. 2003. Thermo-plasticity of clays: an isotropic yield mechanism.
 275 *Computers and Geotechnics*, **30**(8): 649–660. doi: 10.1016/j.compgeo.2003.09.001.

276 Laloui, L. & François, B. 2009. ACMEG-T: soil thermoplasticity model. *Journal of engineering*
 277 *mechanics*, **135**(9): 932-944. doi: 10.1061/(ASCE)EM.1943-7889.0000011.

278 Leroueil, S. & Marques, M.E.S. 1996. Importance of strain rate and temperature effects in
 279 geotechnical engineering. In *Measuring and modeling time-dependent soil behavior*.
 280 Edited by T.C. Sheahan and V.N. Kaliakin. Geotechnical Special Publication 61, ASCE, New
 281 York. pp. 1–60.

282 Mašín, D. 2005. A hypoplastic constitutive model for clays. *International journal for numerical*
 283 *and analytical methods in geomechanics*, **29**(4): 311–336. doi: 10.1002/nag.416.

284 Mašín, D. & Khalili, N. 2012. A thermo-mechanical model for variably saturated soils based on
 285 hypoplasticity. *International journal for numerical and analytical methods in*
 286 *geomechanics*, **36**(12): 1461–1485. doi: 10.1002/nag.1058.

287 Ng, C. W. W., Cheng, Q., Zhou, C. & Alonso, E. E. 2016. Volume changes of an unsaturated clay
 288 during heating and cooling. *Géotechnique Letters*, **6**(3): 1-7. doi: 10.1680/jgele.16.00059.

289 Sultan, N., Delage, P. & Cui, Y. J. 2002. Temperature effects on the volume change behaviour
 290 of Boom clay. *Engineering Geology*, **64**(2): 135–145. doi: 10.1016/S0013-7952(01)00143-
 291 0.

292 Uchaipichat, A. & Khalili, N. 2009. Experimental investigation of thermo-hydro-mechanical
 293 behaviour of an unsaturated silt. *Géotechnique*, **59**(4): 339-353. doi:
 294 10.1680/geot.2009.59.4.339.

295 Vega, A. & McCartney, J. S. 2014. Cyclic heating effects on thermal volume change of silt.
 296 *Environmental Geotechnics*, **2**(5): 257–268. doi: 10.1680/envgeo.13.00022.

297 Zhou, C. & Ng, C. W. W. 2015. A thermomechanical model for saturated soil at small and large
 298 strains. *Canadian Geotechnical Journal*, **52**(8): 1101–1110. doi: 10.1139/cgj-2014-0229.

Tables and Figures

List of tables

Table 1. A summary of model parameters

List of figures

Fig. 1. Concept of the newly introduced thermal stabilization line (TSL)

Fig. 2. Schematic illustration of the proposed approach for simulating volume change of an NC soil specimen subjected to thermal cycles

Fig. 3. Effect of the parameter γ on simulated volume change of an NC soil specimen subjected to thermal cycles

Fig. 4. Typical results of volume change of soil specimens with different OCRs subjected to thermal cycles from the proposed approach

Fig. 5. Comparison of measured and computed results: Normally consolidated remolded illite under isotropic stress condition

Fig. 6. Comparison of measured and computed results: Lightly over-consolidated natural silty-clay under one-dimensional stress condition

Table 1. A summary of model parameters

Soil tested by	Soil type	Mechanical part					Thermal part						
		φ_c (°)	λ	κ	N	r	h_T	n_T	α_s (°C ⁻¹)	T_r (°C)	k_T	c_T	γ
Uchaipichat & Khalili (2009)	Silt	29.5	0.06	0.002	0.772	0.2	0	-0.01	3.5×10^{-5}	25	0.01	0.5	0.1
Campanella & Mitchell (1968)	Clay	22	0.092	0.027	1.178	Nil*	0	-0.009	3.5×10^{-5}	20	Nil*	0.4	0.1
Di Donna & Laloui (2015)	Silty-clay	24	0.023	0.01	0.676	0.6	0	-0.0055	1.8×10^{-5}	20	0.015	0.4	0.1

* This parameter is irrelevant to the simulations conducted in this study.

FIGURE 1 Concept of the newly introduced thermal stabilization line (TSL)

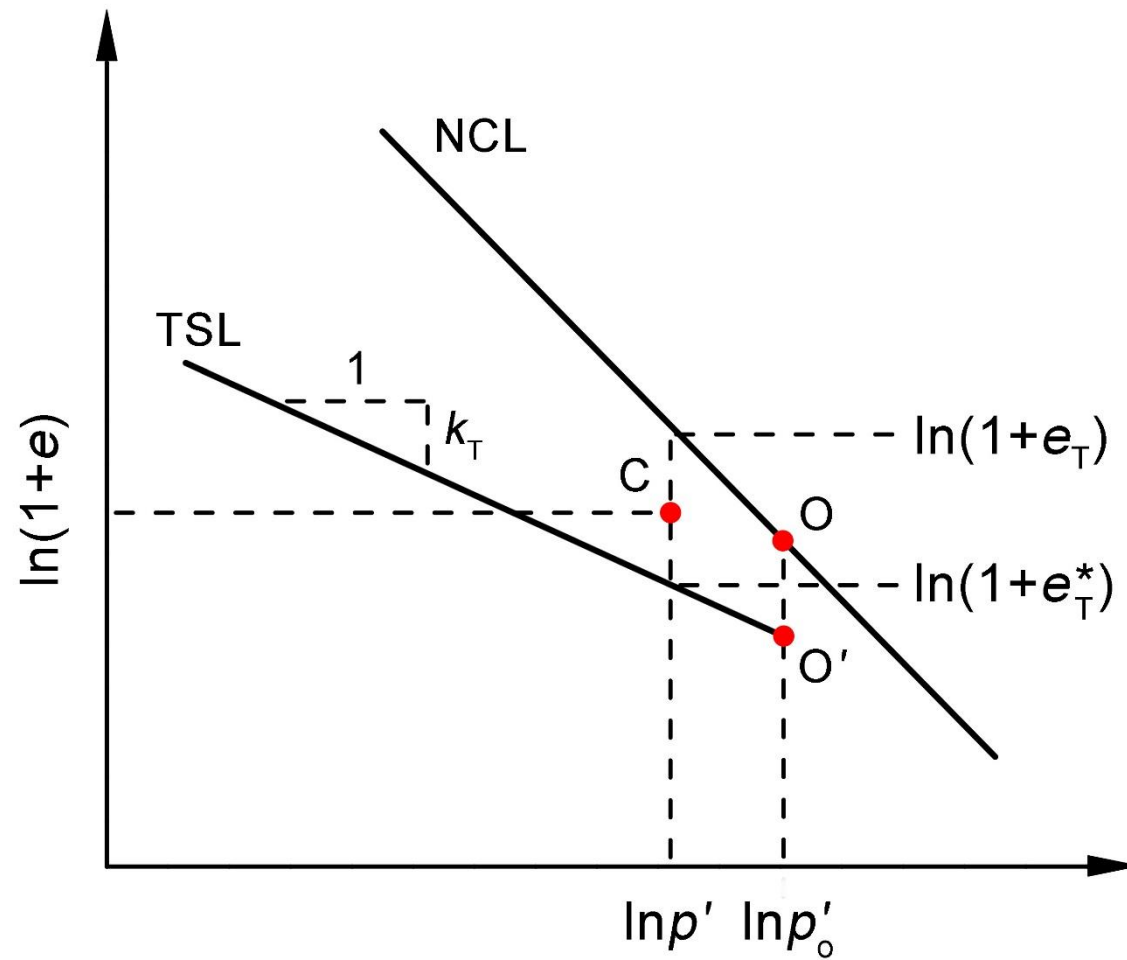


FIGURE 2 Schematic illustration of the proposed approach for simulating volume change of an NC soil specimen subjected to thermal cycles

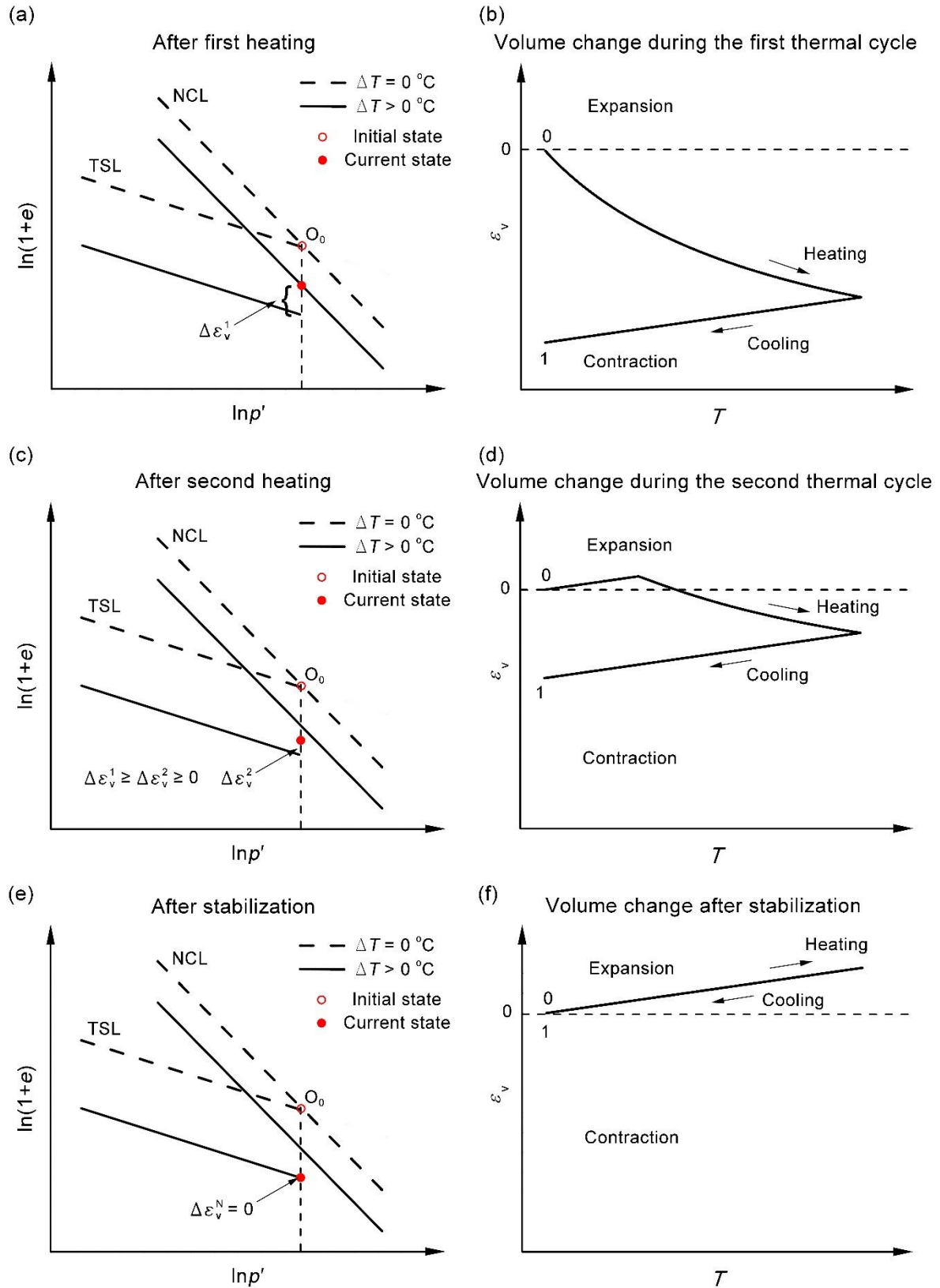


FIGURE 3 Effect of the parameter γ on simulated volume change of an NC soil specimen subjected to thermal cycles

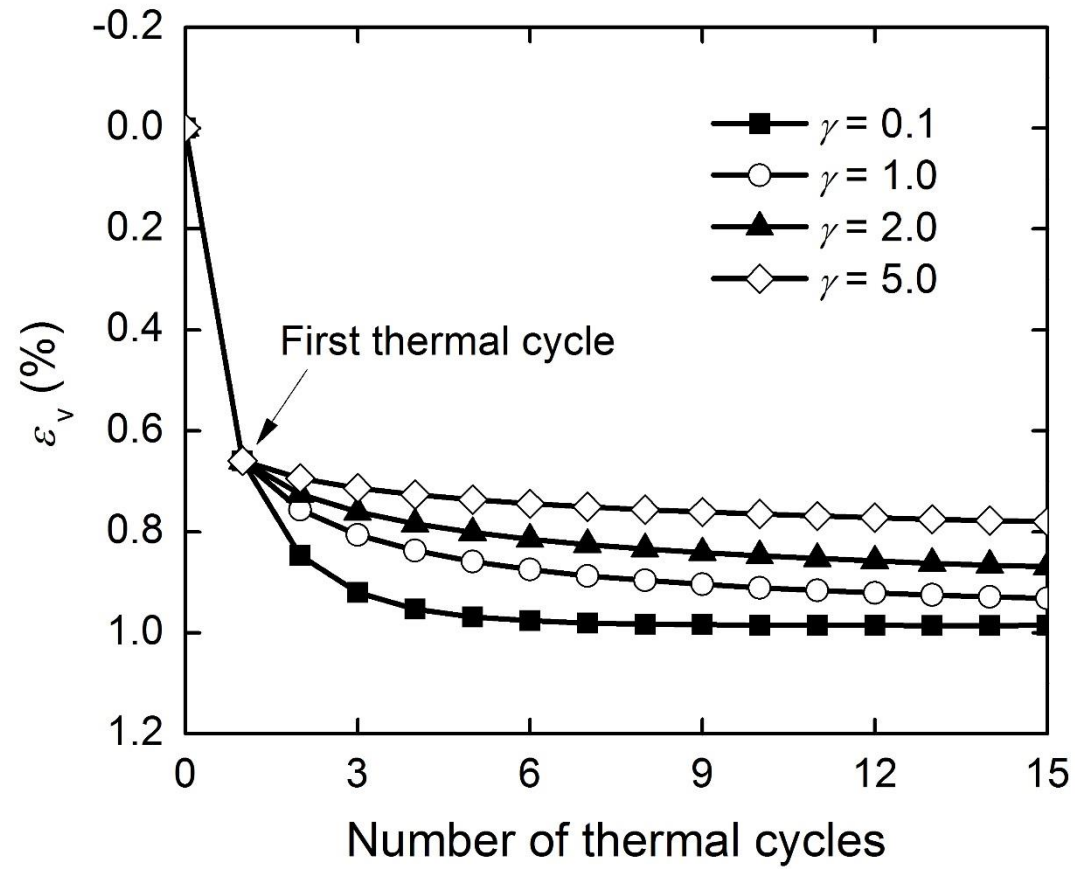


FIGURE 4 Typical results of volume change of soil specimens with different OCRs subjected to thermal cycles from the proposed approach

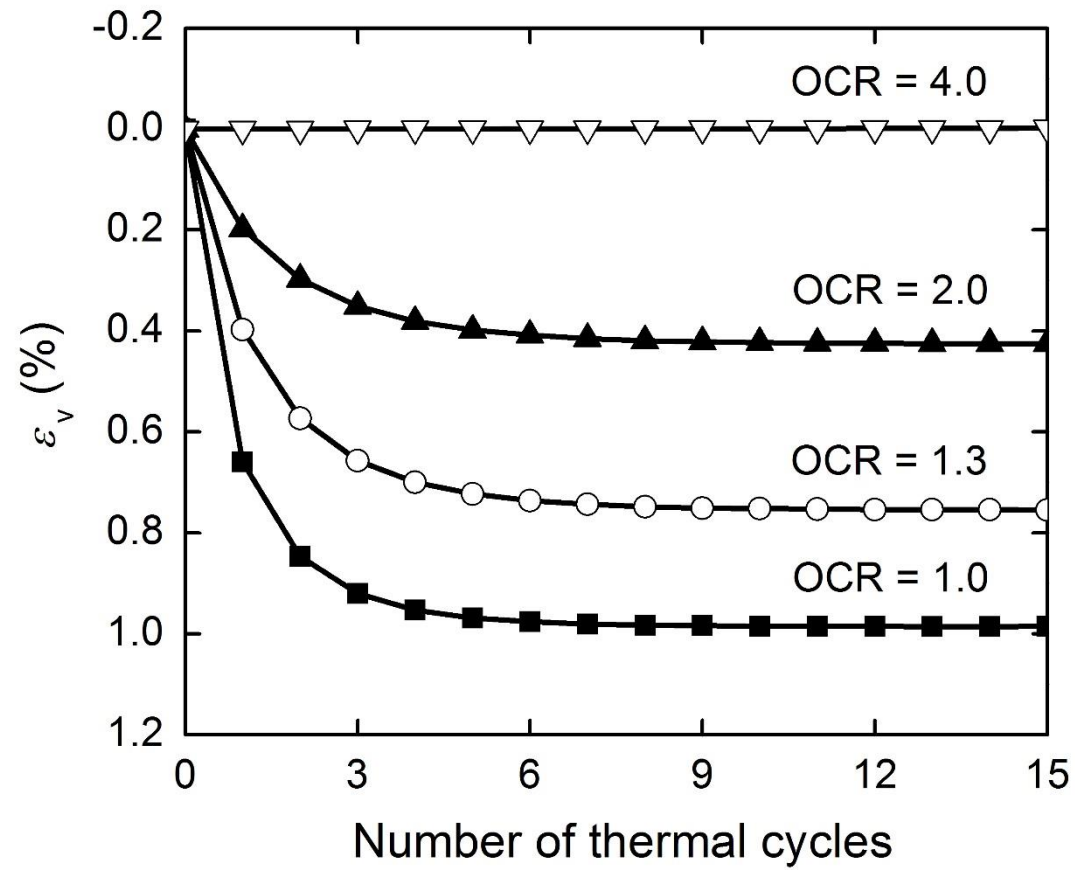


FIGURE 5 Comparison of measured and computed results: Normally consolidated remolded illite under isotropic stress condition

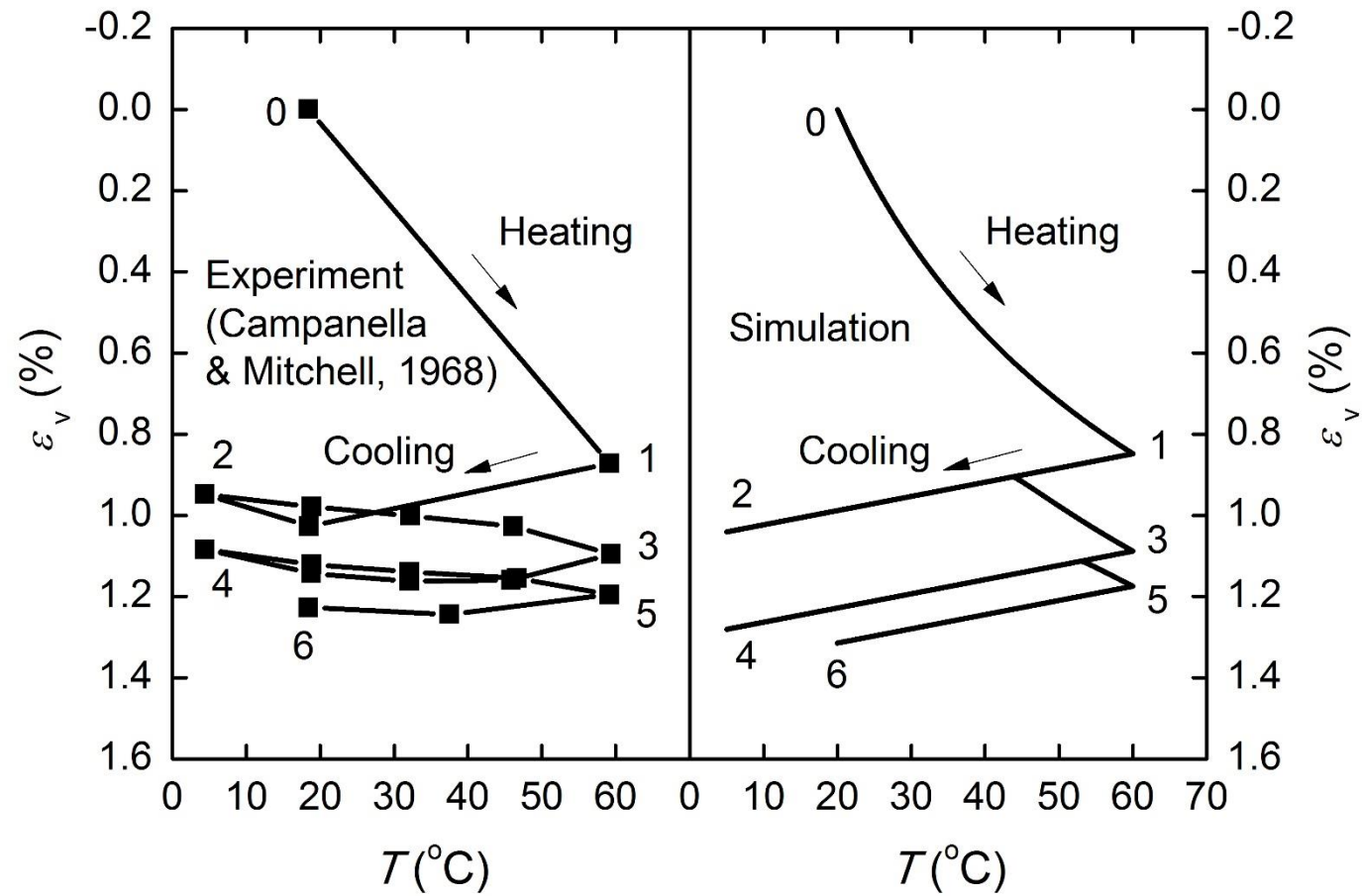


FIGURE 6 Comparison of measured and computed results: Lightly over-consolidated natural silty-clay under one-dimensional stress condition

