

1 Submitted to *Canadian Geotechnical Journal*

2

3 **Incorporating region-variability of model bias into liquefaction-**  
4 **triggering procedures for sandy and gravelly soils through BUS-**  
5 **powered hierarchical Bayesian updating**

6

**Mao-Xin Wang<sup>a</sup>, Y.F. Leung<sup>a</sup>, M.K. Lo<sup>a</sup>**

7

<sup>a</sup> Department of Civil and Environmental Engineering, The Hong Kong Polytechnic

8

University, Hung Hom, Hong Kong.

9

Correspondence to: Y.F. Leung (E-mail: [andy.yf.leung@polyu.edu.hk](mailto:andy.yf.leung@polyu.edu.hk))

10

11

## Abstract

The accuracy of cyclic stress-based liquefaction-triggering assessment procedures can vary systematically from region to region, but it is challenging to regionalize models due to the lack of region-specific data. This paper presents a hierarchical Bayesian modeling (HBM)-based framework for incorporation of inter-region and intra-region variabilities of the bias factor in liquefaction-triggering procedures. A key feature is that the BUS approach (Bayesian Updating with Structural reliability methods) is combined with subset simulation to efficiently update high-dimensional statistics of bias factors. Another feature is a new four-hyperparameter HBM considering both region-specific means and variances of bias factors. This framework is utilized to develop three sets of region-specific liquefaction probability models for practical applications, covering liquefaction-susceptible sandy and gravelly soils. The results show that the four-hyperparameter HBM generally matches better with liquefaction observations and produces larger total variance, compared to the lumped-region modeling and the HBM with only region-specific means. Meanwhile, the population-level distribution and the weighting factor of liquefaction/non-liquefaction occurrence can considerably affect model performance. Furthermore, a discrete integration-based probabilistic method is suggested for liquefaction-triggering hazard assessment. Illustrative examples indicate that different HBM configurations can yield notably different liquefaction hazard results, while neglecting the region-variability tends to be unconservative.

**Keywords:** Earthquake-induced soil liquefaction; Liquefaction-triggering procedure; Inter-region variability; Hierarchical Bayesian modeling; Bayesian evidence; Region-specific liquefaction probability

## 34 **1. Introduction**

35 Earthquake-induced soil liquefaction often causes significant structural damage and disruption  
36 to infrastructure systems. The liquefaction potential of soils is commonly assessed by the semi-  
37 empirical liquefaction-triggering procedures pioneered by Seed and Idriss (1971). Depending  
38 on the types of in-situ tests used to measure liquefaction resistance, these procedures can  
39 generally be classified into standard penetration test (SPT)-, cone penetration test (CPT)-, and  
40 shear wave velocity ( $V_s$ )-based categories (e.g., Robertson and Wride 1998; Youd et al. 2001;  
41 Moss et al. 2006; Boulanger and Idriss 2014; Robertson 2022). While the existing procedures  
42 have been proven to be reliable, many of them follow a deterministic manner by estimating the  
43 nominal factor of safety ( $F_{S0}$ ). Yet, soil liquefaction may still occur if  $F_{S0} > 1$  due to significant  
44 uncertainty involved in the liquefaction triggering analysis. The assessment of liquefaction  
45 potential through probabilistic methods has attracted increasing attention (e.g., Juang et al.  
46 2010; Boulanger and Idriss 2012; Green et al. 2020; Guan and Wang 2022).

47 To enable probabilistic applications of the existing liquefaction-triggering procedures,  
48 extensive efforts have been made to quantify the model uncertainty. For instance, Juang et al.  
49 (2002) developed a generic framework to calibrate the uncertainty of liquefaction-triggering  
50 model using the first-order reliability method. This framework has been subsequently used for  
51 other types of models (e.g., Juang et al. 2012; Duan et al. 2022). In addition, Ku et al. (2012)  
52 evaluated the uncertainty associated with the Robertson and Wride (1998) procedure using the  
53 maximum likelihood method. The Bayesian network has also been introduced to describe the  
54 uncertainty in liquefaction potential assessment (e.g., Cetin et al. 2002; Hu and Liu 2019). The  
55 previous studies generally lumped liquefaction case histories around the world together in

56 analyses, such that the model error (or bias) was implicitly assumed to be consistent across  
57 different regions. Yet, the accuracy of liquefaction-triggering assessment can vary  
58 systematically from region to region (Green 2022), due to the inter-region differences of  
59 ground-motion and geologic/geotechnical characteristics (e.g., Liu and Leung 2018; Baker et  
60 al. 2021; Xiao et al. 2021).

61 An ideal way to improve the model accuracy is to calibrate the model bias using data only  
62 from the region of interest (e.g., Maurer et al. 2019; Green et al. 2020). In most cases, however,  
63 the scarce liquefaction data in a region is statistically not adequate for the calibration. To  
64 address this issue, Zhang et al. (2016) developed a two-hyperparameter hierarchical Bayesian  
65 modeling (HBM) to describe the inter-region variability of the bias factor for a CPT-based  
66 procedure. Subsequently, Zhang et al. (2020) calibrated the inter-region variability for a SPT-  
67 based procedure. These works focused on the sandy soil liquefaction. Yet, documented field  
68 observations have also indicated the consequence of gravel liquefaction (e.g., Brandenberg et  
69 al. 2020; Rollins et al. 2022). Therefore, it is of practical importance to calibrate the region-  
70 specific model bias in liquefaction potential assessments for both sandy and gravelly soils.

71 The HBM regards the group (or region)-specific statistics as random variables that follow  
72 the same population-level distribution parameterized by the hyperparameters. The previously  
73 used two-hyperparameter HBM (termed as HBM-2Hyp hereafter) assumes that only the means  
74 of the bias factor are region-specific following a lognormal distribution (Zhang et al. 2016,  
75 2020). Considering the growing number of documented liquefaction data (Brandenberg et al.  
76 2020; Geyin et al. 2021), it becomes more reasonable to also characterize the inter-region  
77 variability of the standard deviation of the bias factor. The inter-group (e.g., inter-region or

78 inter-site) variability of both the means and the variances of group-level variables has been  
79 incorporated in some previous studies (e.g., Herschtal et al. 2016; Tabarroki et al. 2024). For  
80 instance, Zhang et al. (2014) proposed a four-hyperparameter HBM to describe the inter-site  
81 variability of the mean and the coefficient of variation of the model bias factor for pile  
82 foundation design. Considering the inter-site variability of the mean and covariance, Ching et  
83 al. (2021b) and Wu et al. (2022) developed a computationally efficient HBM framework to  
84 construct quasi-site-specific probability distributions for soil/rock properties. On the other hand,  
85 different types of the population-level distribution such as normal and inverse gamma  
86 distributions have been adopted in HBM applications (e.g., Ching et al. 2021a; Xiao et al. 2021;  
87 Tao et al. 2023). It is thus of interest to examine the effects of the region-specific bias factor  
88 variance and the population-level distribution type on liquefaction potential assessment. A  
89 potential challenge with the incorporation of more HBM hyperparameters and/or target regions  
90 is the high-dimensional Bayesian updating problem, which is difficult to be tackled by classical  
91 Markov Chain Monte Carlo simulation (MCMCS) methods. Meanwhile, a quantitative index  
92 is essential for effectively comparing the performances of different HBM configurations.

93 This study aims (1) to present a Bayesian updating and model selection framework for  
94 calibration of the inter-region and intra-region variabilities of the liquefaction-triggering model  
95 bias; and (2) to develop region-specific liquefaction probability models for both sandy and  
96 gravelly soils. A key feature of this framework is the BUS approach (Bayesian Updating with  
97 Structural reliability methods) (Straub and Papaioannou 2015) with subset simulation (SuS)  
98 (Au and Beck 2001) used for posterior sampling and the Bayesian evidence utilized for model  
99 selection. Besides, a new four-hyperparameter HBM (termed as HBM-4Hyp hereafter) is

100 established to describe the bias factor, where three candidate distribution types for the  
 101 population-level probabilistic modeling are examined. The Bayesian framework is  
 102 subsequently utilized to incorporate the region-variability of model bias into the SPT-based  
 103 Boulanger and Idriss (2012) procedure (SPT-BI12), the CPT-based Robertson and Wride (1998)  
 104 procedure (CPT-RW98), and the Vs-based Rollins et al. (2022) procedure (Vs-R22). Finally, a  
 105 probabilistic liquefaction-triggering hazard analysis method is suggested, and the implication  
 106 of region-feature modeling for liquefaction-triggering assessment is investigated.

## 107 **2. General liquefaction potential assessment**

108 In cyclic stress-based liquefaction-triggering procedures, the cyclic stress ratio (CSR) and  
 109 cyclic resistance ratio (CRR) represent the seismic loading within a soil profile and the  
 110 liquefaction resistance of soils, respectively. The relationship between the actual factor of  
 111 safety ( $F_s$ ) and the nominal  $F_{s0}$  can be expressed as:

$$112 \quad F_s = cF_{s0} = c \frac{\text{CRR}}{\text{CSR}} \quad (1)$$

113 where  $c$  denotes the model bias factor due to imperfect prediction and the model error is  
 114 neglected when using  $c = 1$ ; CSR is usually estimated as (e.g., Youd et al. 2001):

$$115 \quad \text{CSR} = 0.65 \frac{\sigma_v}{\sigma'_v} \frac{a_{\max}}{g} r_d \frac{1}{MSF} \frac{1}{K_\sigma} \quad (2)$$

116 where  $\sigma_v$  and  $\sigma'_v$  are the vertical total and effective stresses at depth  $z$ , respectively;  $a_{\max}$  is  
 117 peak ground acceleration;  $g$  is gravitational acceleration;  $r_d$  is the shear stress reduction factor  
 118 that accounts for the deformability of soil profile;  $MSF$  is the earthquake magnitude scaling  
 119 factor that accounts for the ground motion duration;  $K_\sigma$  is the overburden correction factor.

120 Detailed mathematical expressions of CRR,  $r_d$ ,  $MSF$ , and  $K_\sigma$  for the adopted SPT-BI12,

121 CPT-RW98, and Vs-R22 procedures are presented in the Supplementary material. Although  
 122 developed using post-earthquake case histories across different regions, the existing procedures  
 123 would be directly applied to specific sites, as demonstrated in Fig. 1.

124 The factor  $c$  in Eq. (1) is usually assumed to follow a lognormal distribution with the mean  
 125 of  $\mu_c$  and standard deviation of  $\sigma_c$  (e.g., Juang et al. 2012; Zhang et al. 2016), resulting in the  
 126 normally distributed  $\ln F_S$  as follows:

$$127 \quad \ln F_S \sim \text{Normal} \left\{ \ln \mu_c - 0.5 \ln \left( 1 + \frac{\sigma_c^2}{\mu_c^2} \right) + \ln F_{S0}, \ln \left( 1 + \frac{\sigma_c^2}{\mu_c^2} \right) \right\} \quad (3)$$

128 Therefore, the probability of liquefaction ( $P_{L0}$ ) can be calculated as:

$$129 \quad P_{L0} = P(F_S < 1) = P(\ln F_S < 0) = \Phi \left[ \frac{-\ln \mu_c + 0.5 \ln \left( 1 + \frac{\sigma_c^2}{\mu_c^2} \right) - \ln F_{S0}}{\sqrt{\ln \left( 1 + \frac{\sigma_c^2}{\mu_c^2} \right)}} \right] \quad (4)$$

130 where  $\Phi[\cdot]$  is the cumulative distribution function (CDF) of the standard normal variable.

131 Based on the liquefaction case histories,  $\mu_c$  and  $\sigma_c$  can be estimated from Bayesian inference  
 132 by lumping data across different regions, as illustrated in Fig. 2a.

### 133 **3. Bayesian framework for development of region-specific liquefaction-triggering models**

#### 134 **3.1. Inter- and intra-region variabilities of model bias**

135 The model bias may be similar for different cases in a region (e.g., low intra-region variability)  
 136 while different regions may result in systematically different biases (e.g., high inter-region  
 137 variability). Thus, the total variability of the bias factor can be separated into the inter- and  
 138 intra-region components by defining the region-specific bias factor  $c_i$  ( $i = 1, 2, \dots, r$ ), where  $r$   
 139 denotes the number of regions considered. Similar to Eq. (3),  $c_i$  is assumed to follow a  
 140 lognormal distribution with the mean of  $\mu_{c_i}$  and standard deviation of  $\sigma_{c_i}$ . These statistics can  
 141 be estimated from liquefaction data through two Bayesian modeling ideas, which are shown in

142 Fig. 2b and c. The HBM-2Hyp uses region-specific  $\mu_{ci}$  and region-invariant  $\sigma_c = \sigma_{c1} = \sigma_{c2} = \dots$   
 143  $= \sigma_{cr}$ , while the HBM-4Hyp considers both  $\mu_{ci}$  and  $\sigma_{ci}$  as region-specific variables.

144 The region-specific means or standard deviations are considered to be similar by assuming  
 145 a population-level distribution of  $\mu_{ci}$  (or  $\sigma_{ci}$ ) parametrized by two hyperparameters, which are  
 146 the mean  $\mu_\mu$  (or  $\mu_\sigma$ ) and standard deviation  $\sigma_\mu$  (or  $\sigma_\sigma$ ) of  $\mu_{ci}$  (or  $\sigma_{ci}$ ). As a result, there are two  
 147 hyperparameters ( $\mu_\mu$  and  $\sigma_\mu$ ) and four hyperparameters ( $\mu_\mu, \sigma_\mu, \mu_\sigma,$  and  $\sigma_\sigma$ ) in HBM-2Hyp and  
 148 HBM-4Hyp, respectively (see Fig. 2). The hyperparameter  $\mu_\mu$  represents the ‘‘average’’ model  
 149 bias across different regions, whereas  $\sigma_\mu$  measures the inter-region variability of the average  
 150 bias. Also,  $\mu_\sigma$  represents the ‘‘average’’ intra-region variability and  $\sigma_\sigma$  reflects the inter-region  
 151 variability of the bias uncertainty. Suppose that  $\Theta$  is the vector of all uncertain parameters  
 152 required in the probabilistic modeling, resulting in  $\Theta = [\mu_\mu, \sigma_\mu, \sigma_c, \mu_{c1}, \mu_{c2}, \dots, \mu_{cr}]$  and  $[\mu_\mu, \sigma_\mu,$   
 153  $\mu_\sigma, \sigma_\sigma, \mu_{c1}, \mu_{c2}, \dots, \mu_{cr}, \sigma_{c1}, \sigma_{c2}, \dots, \sigma_{cr}]$  for HBM-2Hyp and HBM-4Hyp, respectively.

154 For given  $[\mu_\mu, \sigma_\mu]$  or  $[\mu_\sigma, \sigma_\sigma]$ ,  $\mu_{ci}$  or  $\sigma_{ci}$  can be described by the population-level distribution.  
 155 Three candidate distributions are considered in this study, including lognormal distribution,  
 156 normal distribution truncated to be positive (Trunc-normal), and Gumbel distribution truncated  
 157 to be positive (Trunc-Gumbel). For instance, the probability distribution function (PDF) of  $\mu_{ci}$   
 158 or  $\sigma_{ci}$  for lognormal distribution can be expressed as (e.g., Wang et al. 2020; Deng et al. 2022):

$$159 \quad f(x_{ci} | \mu_x, \sigma_x, Mdl) = \frac{1}{x_{ci} \sqrt{2\pi \ln(1 + \sigma_x^2 / \mu_x^2)}} \exp \left\{ - \frac{[\ln x_{ci} - \ln \mu_x + 0.5 \ln(1 + \sigma_x^2 / \mu_x^2)]^2}{2 \ln(1 + \sigma_x^2 / \mu_x^2)} \right\} \quad (5)$$

160 where  $x$  denotes ‘‘ $\mu$ ’’ for  $f(\mu_{ci} | \mu_\mu, \sigma_\mu, Mdl)$  or ‘‘ $\sigma$ ’’ for  $f(\sigma_{ci} | \mu_\sigma, \sigma_\sigma, Mdl)$ ;  $Mdl$  represents the  
 161 model configuration (e.g.,  $Mdl = [\text{CPT-4Hyp}, \text{lognormal}]$ ). The PDFs for the other candidate  
 162 distributions are listed in Table 1, where  $p$  and  $q$  denote distribution-specific parameters and

163 are transformed from  $[\mu_\mu, \sigma_\mu]$  or  $[\mu_\sigma, \sigma_\sigma]$ . The parameters  $\mu_\mu, \sigma_\mu, \mu_\sigma,$  and  $\sigma_\sigma$  are random variables,  
 164 and their samples are generated through MCS or MCMCS in Bayesian updating. The statistics  
 165 of these parameters can be estimated from posterior samples and will be explicitly provided in  
 166 the next section. Increasing  $\mu_\mu$  and  $\sigma_\mu$  would lead to lower overall  $P_{L0}$  and larger inter-region  
 167 difference in  $P_{L0}$ , respectively. Additionally,  $P_{L0}$  for each region would generally become  
 168 greater with increasing  $\mu_\sigma$  at  $F_{S0} > 1$  levels, whereas increasing  $\sigma_\sigma$  tends to enlarge the inter-  
 169 region difference of  $P_{L0}$ .

### 170 3.2. Hierarchical Bayesian modeling of the inter- and intra-region variabilities

171 Let  $\mathbf{Data} = [\mathbf{Dt}_1, \mathbf{Dt}_2, \dots, \mathbf{Dt}_r]$  denote case histories from  $r$  earthquake influencing regions and  
 172  $\mathbf{Dt}_i$  denote  $[Dt_{i1}, Dt_{i2}, \dots, Dt_{i,n_i}]$ , where  $Dt_{ij} = 1$  for liquefied case and 0 otherwise, and  $n_i$  is the  
 173 number of observations in the  $i$ -th region. For given  $Mdl$  and  $\mathbf{Data}$ ,  $\Theta$  can be estimated by  
 174 Bayesian updating, following the general idea below (e.g., Cao et al. 2019; Lo and Leung 2019;  
 175 Zhang et al. 2022):

$$176 \quad f(\Theta | \mathbf{Data}, Mdl) = KL(\mathbf{Data} | \Theta, Mdl) f(\Theta | Mdl) \quad (6)$$

177 where  $f(\Theta | \mathbf{Data}, Mdl)$  denotes the posterior joint PDF of  $\Theta$ ;

178  $K = \left[ \int_{\Theta} L(\mathbf{Data} | \Theta, Mdl) f(\Theta | Mdl) d\Theta \right]^{-1}$  is a normalization constant that is independent of  $\Theta$ ;

179  $f(\Theta | Mdl)$  is the prior joint PDF of  $\Theta$  and is expressed as:

$$180 \quad f(\Theta | Mdl) = \begin{cases} f(\mu_\mu) f(\sigma_\mu) f(\sigma_c) \prod_{i=1}^r f(\mu_{ci} | \mu_\mu, \sigma_\mu, Mdl) & \text{for 2Hyp} \\ f(\mu_\mu) f(\sigma_\mu) f(\mu_\sigma) f(\sigma_\sigma) \prod_{i=1}^r [f(\mu_{ci} | \mu_\mu, \sigma_\mu, Mdl) f(\sigma_{ci} | \mu_\sigma, \sigma_\sigma, Mdl)] & \text{for 4Hyp} \end{cases} \quad (7)$$

181 where  $f(\mu_{ci} | \mu_\mu, \sigma_\mu)$  and  $f(\sigma_{ci} | \mu_\sigma, \sigma_\sigma)$  are obtained from Eq. (5) or equations presented

183 in Table 1;  $f(\mu_\mu), f(\sigma_\mu), f(\mu_\sigma),$  and  $f(\sigma_\sigma)$  are the prior marginal PDFs of  $\mu_\mu, \sigma_\mu, \mu_\sigma,$  and  $\sigma_\sigma,$

184 respectively. According to the previous studies (Zhang et al. 2016, 2020), each prior PDF is  
 185 taken as 1/3 by assuming a uniform distribution within the [0, 3] interval.

186 The  $L(\mathbf{Data}|\Theta, Mdl)$  term in Eq. (6) represents the likelihood function for the observed data,  
 187 and is expressed as:

$$188 \quad L(\mathbf{Data}|\Theta, Mdl) = \prod_{i=1}^r \left[ \prod_{j=1}^{n_{Li}} P_{Lij}^{w_{Li}(Dt_{ij})} \prod_{j=n_{Li}+1}^{n_i} (1 - P_{Lij})^{w_{NLI}(1-Dt_{ij})} \right] \quad (8)$$

189 where  $n_{Li}$  is the number of liquefied cases in the  $i$ -th region;  $Dt_{ij}$  is taken as 1 for  $j = 1, 2, \dots,$   
 190  $n_{Li}$  and 0 for  $j \geq n_{Li}+1$ ;  $P_{Lij}$  is the predicted liquefaction probability for the  $j$ -th case history in  
 191 the  $i$ -th region, and is estimated by adjusting Eq. (4) as:

$$192 \quad P_{Lij} = \Phi \left( \frac{-\ln \mu_{ci} + 0.5 \ln(1 + \sigma_{ci}^2 / \mu_{ci}^2) - \ln F_{Sij}}{\sqrt{\ln(1 + \sigma_{ci}^2 / \mu_{ci}^2)}} \right) \quad (9)$$

193 in which  $\sigma_{ci}$  is simply taken as  $\sigma_c$  in the 2Hyp modeling;  $F_{Sij}$  is the calculated safety factor.

194 As parameters used to reduce the sampling bias effects,  $w_{Li}$  and  $w_{NLI}$  in Eq. (8) are the  
 195 weighting factors of the liquefaction and non-liquefaction cases in the  $i$ -th region, respectively,  
 196 which are defined as (e.g., Cetin et al. 2002; Zhang et al. 2016):

$$197 \quad w_{Li} = \frac{Q_p}{Q_{si}} = \frac{Q_p (n_{Li} + n_{NLI})}{n_{Li}} \quad (10)$$

$$198 \quad w_{NLI} = \frac{1 - Q_p}{1 - Q_{si}} = \frac{(1 - Q_p)(n_{Li} + n_{NLI})}{n_{NLI}} \quad (11)$$

199 where  $Q_{si}$  is the sample proportion of liquefied cases in the  $i$ -th region;  $Q_p$  denotes the true  
 200 proportion of liquefied cases in nature. Since  $Q_p$  is always unknown, a typical way of estimating  
 201  $w_{NLI}$  and  $w_{Li}$  is to specify a reasonable  $w_{NLI}/w_{Li}$  ratio. Cetin et al. (2002) consulted with a panel  
 202 of eight experts, suggesting that the most common estimate of the ratio should be between 1.5  
 203 and 2. They further conducted Bayesian updating analyses and found that  $w_{NLI}/w_{Li}=1.5$  yielded

204 the optimal result. This value has been consistently adopted in many subsequent studies (e.g.,  
 205 Boulanger and Idriss 2012), regardless of the specific databases and methods.

206 Using  $w_{NLi}/w_{Li} = 1.5$  and  $Q_{si} = Q_{s,total} = \sum_{i=1}^r n_{Li} / \sum_{i=1}^r (n_{Li} + n_{NLi})$ ,  $Q_p$  can thus be computed  
 207 as:

$$208 \quad Q_p = \frac{1}{1.5/Q_{s,total} - 0.5} = \frac{1}{1.5 \sum_{i=1}^r (n_{Li} + n_{NLi}) / \sum_{i=1}^r n_{Li} - 0.5} \quad (12)$$

209 where  $Q_{s,total}$  represents the total sample proportion of all liquefaction cases. Given  $Q_{s,total} =$   
 210 0.557 (Cetin et al. 2002), Zhang et al. (2016) derived  $Q_p = 0.456$  for their model development,  
 211 and this value was subsequently used for another dataset (Zhang et al. 2020). Yet,  $Q_p$  may be  
 212 considered as dataset-specific as long as the weighting ratio of 1.5 is satisfied. In this study,  
 213 both the fixed  $Q_p = 0.456$  and the calculated dataset-specific  $Q_p$  (from Eq. (12)) are taken as  
 214 candidate modeling choices, and their resulting models will be compared later.

### 215 3.3. Bayesian updating through the BUS approach with subset simulation

216 The BUS approach (Straub and Papaioannou 2015) is utilized herein to evaluate the posterior  
 217 PDF of  $\Theta$  in Eq. (7). This approach is powerful in dealing with high-dimensional Bayesian  
 218 inference (Betz et al. 2018a, 2018b; Jiang et al. 2018). The basic idea of BUS is to treat the  
 219 updating problem as a structural reliability problem such that the samples in an equivalent  
 220 failure domain  $\Omega_{\Theta}$  are taken as the posterior samples. The observed information described by  
 221  $L(\mathbf{Data}|\Theta, Mdl)$  is used to define  $\Omega_{\Theta}$  as:

$$222 \quad \Omega_{\Theta} = \{H(\Theta_+ | Mdl) \leq 0\} \quad (13)$$

223 where  $\Theta_+$  is the augmented outcome space  $[\Theta, \zeta]$ ;  $\zeta$  is the standard uniform variable in  $[0, 1]$

224 interval;  $H(\Theta_+)$  is the limit state function:

$$225 \quad H(\Theta_+ | Mdl) = \ln \zeta - \ln [\kappa L(\mathbf{Data} | \Theta, Mdl)] \quad (14)$$

226 where  $\kappa$  is the likelihood multiplier to ensure  $\kappa L(\mathbf{Data} | \Theta, Mdl) \leq 1$  and will be determined  
227 adaptively as discussed later.

228 The subset simulation (Au and Beck 2001) is coupled with the BUS approach (termed as  
229 BUS-SuS hereafter) to efficiently solve the reliability problem in obtaining the posterior  
230 samples. Consequently, the occurrence probability of the event  $Z = \{\Theta_+ \in \Omega_\Theta\}$  can be  
231 represented as a product of larger conditional probabilities of a set of intermediate events  $Z_l =$   
232  $\{H(\Theta_+ | Mdl) < h_l\}$  ( $l = 1, 2, \dots, m$ ), expressed as:

$$233 \quad P\{\Theta_+ \in \Omega_\Theta\} = P(Z) = P(Z_1) \prod_{l=2}^m P(Z_l | Z_{l-1}) \quad (15)$$

234 where these intermediate events satisfy  $Z_1 \supset Z_2 \supset \dots \supset Z_{m-1} \supset Z_m$ ;  $h_1 > h_2 > \dots > h_m = 0$  are a  
235 decreasing series of threshold values to be determined automatically. The intermediate  
236 probabilities  $P(Z_1)$  and  $P(Z_l | Z_{l-1})$  are equal to a predetermined value  $p_0$ , which is specified as  
237 0.1 (Betz et al. 2018b).

238 BUS-SuS starts with a direct Monte Carlo simulation (MCS) of  $N_l$  sets of  $\Theta_+$  samples,  
239 based on (i) the  $[0, 3]$  uniform distributions for  $[\mu_\mu, \sigma_\mu, \sigma_c]$  or  $[\mu_\mu, \sigma_\mu, \mu_\sigma, \sigma_\sigma]$ , (ii) the population-  
240 level distributions for  $\mu_{ci}$  or  $[\mu_{ci}, \sigma_{ci}]$ , and (iii) standard uniform distribution for  $\zeta$  in the first  
241 level ( $l = 1$ ). Specifically, the population-level sampling is performed with isoprobabilistic  
242 transformation expressed as:

$$243 \quad \mu_{ci} = F^{-1}(\Phi(\xi) | \mu_\mu, \sigma_\mu, Mdl) \quad (16a)$$

$$244 \quad \sigma_{ci} = F^{-1}(\Phi(\xi) | \mu_\sigma, \sigma_\sigma, Mdl) \quad (16b)$$

245 where  $\xi$  is a random realization of the standard normal variable;  $F^{-1}(\cdot|\mu_\mu, \sigma_\mu, Mdl)$  and  
 246  $F^{-1}(\cdot|\mu_\sigma, \sigma_\sigma, Mdl)$  represent the inverses of CDFs of  $\mu_{ci}$  and  $\sigma_{ci}$ , respectively, which are  
 247 included in Table 1.

248 The  $\Theta_+$  samples are substituted into Eq. (14) to calculate the corresponding  $H(\Theta_+|Mdl)$   
 249 values, which are ranked in an ascending order and the  $(p_0N_l+1)$ -th ranking value is taken as  
 250  $h_1$ . Consequently, the  $p_0N_s$  samples with the lowest  $H(\Theta_+|Mdl)$  values are regarded as seeds to  
 251 generate  $(1-p_0)N_s$  new samples of  $\Theta_+$  for the  $Z_2$  event via MCMCS (Papaioannou et al. 2015).  
 252 Both MCS and MCMCS are performed in the standard normal space. Similarly, the sample  
 253 ranking and MCMCS are conducted for the next SuS level ( $l = 2$ ), producing  $N_l$  samples of  $\Theta_+$ .  
 254 This process will be continued until the  $(p_0N_l+1)$ -th ranking value of  $H(\Theta_+)$  is equal to or less  
 255 than 0. During the process,  $\kappa$  in Eq. (14) is evaluated adaptively as the reciprocal of the  
 256 maximum of the  $L(\mathbf{Data}|\Theta, Mdl)$  values in the  $l$ -th SuS level, and is written as (Jiang et al. 2018):

$$257 \quad \kappa_l = \frac{1}{\max\left\{1/\kappa_{l-1}, [L_{l,1}, L_{l,2}, \dots, L_{l,N_l}]\right\}} \quad (17)$$

258 where  $\kappa_{l-1}$  and  $\kappa_l$  represent the constant values in the  $(l-1)$ -th and  $l$ -th levels, respectively. Once  
 259 the SuS process is finished, the failure samples in the final level are taken as seeds to generate  
 260  $N_l$  samples in the  $\Omega_\Theta$  domain, which are treated as the posterior samples of interest.

261 The resulting posterior statistics (i.e., mean  $E(\cdot)$  and standard deviation  $\text{Std}(\cdot)$ ) of  $\Theta = [\mu_\mu,$   
 262  $\sigma_\mu, \sigma_c, \mu_{c1}, \mu_{c2}, \dots, \mu_{cr}]$  or  $[\mu_\mu, \sigma_\mu, \mu_\sigma, \sigma_\sigma, \mu_{c1}, \mu_{c2}, \dots, \mu_{cr}, \sigma_{c1}, \sigma_{c2}, \dots, \sigma_{cr}]$  are used to estimate  $E(\cdot)$   
 263 and  $\text{Std}(\cdot)$  of the region-specific bias factors  $c_i$  ( $i = 1, 2, \dots, r$ ) through the laws of total  
 264 expectation and total variance, which can be expressed as:

$$265 \quad E(c_i) = E(\mu_{ci}) \quad (18)$$

$$\text{Std}(c_i) = \begin{cases} \sqrt{E^2(\sigma_c) + \text{Std}^2(\mu_{c_i}) + \text{Std}^2(\sigma_c)} & \text{for 2Hyp} \\ \sqrt{E^2(\sigma_{c_i}) + \text{Std}^2(\mu_{c_i}) + \text{Std}^2(\sigma_{c_i})} & \text{for 4Hyp} \end{cases} \quad (19)$$

For other regions outside the dataset, the bias factor is denoted as  $c_0$ , of which  $E(\cdot)$  and  $\text{Std}(\cdot)$  can be similarly estimated as:

$$E(c_0) = E(\mu_\mu) \quad (20)$$

$$\text{Std}(c_0) = \sqrt{E^2(\sigma_{c_0}) + \text{Std}^2(\mu_{c_0}) + \text{Std}^2(\sigma_{c_0})} \quad (21)$$

where  $E(\sigma_{c_0})$  is equal to  $E(\sigma_c)$  and  $E(\mu_\sigma)$  for 2Hyp and 4Hyp, respectively;  $\text{Std}(\sigma_{c_0})$  is equal to  $\text{Std}(\sigma_c)$  and  $\sqrt{E^2(\sigma_\sigma) + \text{Std}^2(\mu_\sigma) + \text{Std}^2(\sigma_\sigma)}$  for 2Hyp and 4Hyp, respectively;  $\text{Std}(\mu_{c_0}) = \sqrt{E^2(\sigma_\mu) + \text{Std}^2(\mu_\mu) + \text{Std}^2(\sigma_\mu)}$ .

### 3.4. Model selection through Bayesian evidence with Gaussian copula

When  $N_c$  candidate HBM configurations  $Mdl_k$  ( $k = 1, 2, \dots, N_c$ ) are available, it is desirable to select the most probable candidate through Bayesian model comparison techniques, such as residual analysis, leave-some-out cross-validation analysis, likelihood function evaluation, and deviance information criterion (e.g., Bozorgzadeh et al. 2019; Wu et al. 2022; Tao et al. 2023; Tabarrokhi et al. 2024). One of the most commonly used performance indexes is the evidence derived within Bayesian analysis (e.g., Yuen 2010; Cao and Wang 2014). For instance, Zhang et al. (2016) suggested an evidence calculation method for comparing HBM-based liquefaction-triggering models. In this study, a copula-based method is introduced to derive the evidence for HBM. Therefore, the general case where the posterior PDF may not be approximated by a specific distribution (e.g., normal distribution) can be addressed. Previous studies have shown that copula functions provide a flexible way of establishing the joint distribution in geotechnical applications (e.g., Tian et al. 2016; Tang et al. 2020).

287 The plausibility of a model class  $Mdl_k$  given  $\mathbf{Data}$  can be estimated by applying Bayes'  
 288 theorem (e.g., Cao and Wang 2014; Zhang et al. 2016):

$$289 \quad P(Mdl_k | \mathbf{Data}) = \frac{P(\mathbf{Data} | Mdl_k) P(Mdl_k)}{P(\mathbf{Data})} \quad (22)$$

290 where  $P(\mathbf{Data})$  is a normalization constant independent of  $Mdl_k$ ;  $P(Mdl_k)$  is the prior probability  
 291 of the  $k$ -th candidate and is typically set to  $1/N_c$  for equal weights. Therefore, the ranking of  
 292  $P(Mdl_k | \mathbf{Data})$  for different models is only dependent on  $P(\mathbf{Data} | Mdl_k)$ , which is usually termed  
 293 as *evidence*, and can be expressed as:

$$294 \quad P(\mathbf{Data} | Mdl_k) = \int_{\Theta} L(\mathbf{Data} | \Theta, Mdl_k) f(\Theta | Mdl_k) d\Theta \quad (23)$$

295 For a general case without requirement for normality of the posterior distribution, the logarithm  
 296 of evidence can be derived as (Yuen 2010):

$$297 \quad \begin{aligned} \ln [P(\mathbf{Data} | Mdl_k)] &= \int_{\Theta} \ln [L(\mathbf{Data} | \Theta, Mdl_k)] f(\Theta | \mathbf{Data}, Mdl_k) d\Theta \\ &\quad - \int_{\Theta} \ln \left[ \frac{f(\Theta | \mathbf{Data}, Mdl_k)}{f(\Theta | Mdl_k)} \right] f(\Theta | \mathbf{Data}, Mdl_k) d\Theta \end{aligned} \quad (24)$$

298 where the first integral represents the average logarithmic goodness-of-fit of a candidate model,  
 299 and the second integral measures the relative entropy between the prior and posterior PDFs.

300 Given  $N_c$  posterior samples of  $\Theta$  derived from the BUS-SuS approach, Eq. (24) can be  
 301 numerically evaluated through:

$$302 \quad \ln [P(\mathbf{Data} | Mdl_k)] = \frac{1}{N_l} \left\{ \sum_{j=1}^{N_l} \ln [L(\mathbf{Data} | \hat{\Theta}_j, Mdl_k)] - \sum_{j=1}^{N_l} \ln \left[ \frac{f(\hat{\Theta}_j | \mathbf{Data}, Mdl_k)}{f(\hat{\Theta}_j | Mdl_k)} \right] \right\} \quad (25)$$

303 where  $f(\hat{\Theta}_j | Mdl_k)$  and  $L(\mathbf{Data} | \hat{\Theta}_j, Mdl_k)$  are calculated from Eqs. (7) and (8) with the  $j$ -th  
 304 posterior sample  $\hat{\Theta}_j$ , respectively;  $f(\hat{\Theta}_j | \mathbf{Data}, Mdl_k)$  can be expressed through copula theory  
 305 (e.g., Tian et al. 2016; Tang et al. 2020), representing the joint posterior PDF of  $\Theta$  as the product

306 of the posterior marginal PDFs of individual parameters and a copula density function. A  
 307 feature of the copula method is its flexibility in characterizing the posterior PDF.

308 In this study, the Gaussian copula is selected such that  $f(\hat{\Theta}_j | \mathbf{Data}, Mdl_k)$  is estimated as:

$$309 \quad f(\hat{\Theta} | \mathbf{Data}, Mdl_k) = \begin{cases} D(\hat{F}(\hat{\mu}_\mu), \hat{F}(\hat{\sigma}_\mu), \hat{F}(\hat{\sigma}_c), \hat{F}(\hat{\mu}_{c1}), \hat{F}(\hat{\mu}_{c2}), \dots, \hat{F}(\hat{\mu}_{cr}); \boldsymbol{\theta}) \\ \quad \times \hat{f}(\hat{\mu}_\mu) \hat{f}(\hat{\sigma}_\mu) \hat{f}(\hat{\sigma}_c) \prod_{i=1}^r \hat{f}(\hat{\mu}_{ci}) & \text{for 2Hyp} \\ D(\hat{F}(\hat{\mu}_\mu), \hat{F}(\hat{\sigma}_\mu), \hat{F}(\hat{\mu}_\sigma), \hat{F}(\hat{\sigma}_\sigma), \\ \hat{F}(\hat{\mu}_{c1}), \hat{F}(\hat{\mu}_{c2}), \dots, \hat{F}(\hat{\mu}_{cr}), \hat{F}(\hat{\sigma}_{c1}), \hat{F}(\hat{\sigma}_{c2}), \dots, \hat{F}(\hat{\sigma}_{cr}); \boldsymbol{\theta}) \\ \quad \times \hat{f}(\hat{\mu}_\mu) \hat{f}(\hat{\sigma}_\mu) \hat{f}(\hat{\mu}_\sigma) \hat{f}(\hat{\sigma}_\sigma) \prod_{i=1}^r \hat{f}(\hat{\mu}_{ci}) \prod_{i=1}^r \hat{f}(\hat{\sigma}_{ci}) & \text{for 4Hyp} \end{cases} \quad (26)$$

310 where  $\hat{F}(\cdot)$  represents the empirical marginal CDFs for the posterior samples of individual  
 311 parameters;  $\hat{f}(\cdot)$  denotes the kernel smoothing-based marginal PDFs for the posterior  
 312 parameter samples;  $D(\cdot)$  is the copula density function and is written as:

$$313 \quad D(u_1, u_2, \dots, u_n; \boldsymbol{\theta}) = \det(\boldsymbol{\theta})^{-1/2} \exp\left[-\frac{1}{2} \boldsymbol{\eta}(\boldsymbol{\theta}^{-1} - \mathbf{I}) \boldsymbol{\eta}^T\right] \quad (27)$$

314 where  $n$  is the number of variables considered;  $\boldsymbol{\eta}$  is a 1-by- $n$  vector  $[\Phi^{-1}(u_1), \Phi^{-1}(u_2), \dots, \Phi^{-1}(u_n)]$ ,  
 315 in which  $\Phi^{-1}(\cdot)$  represents the inverse of CDF of the standard normal variable;  $\det(\cdot)$  is  
 316 determinant;  $T$  denotes matrix transpose;  $\mathbf{I}$  is an  $n$ -by- $n$  identity matrix;  $\boldsymbol{\theta}$  is a  $n$ -by- $n$   
 317 dependence parameter matrix. The maximum-likelihood estimation is adopted to derive  $\boldsymbol{\theta}$  as:

$$318 \quad \hat{\boldsymbol{\theta}} = \arg \max_{\boldsymbol{\theta}} \sum_{j=1}^{N_j} \ln D(u_{1j}, u_{2j}, \dots, u_{nj}; \boldsymbol{\theta}) \quad (28)$$

319 For each candidate model configuration, the posterior samples of  $\boldsymbol{\Theta}$  generated from the BUS-  
 320 SuS approach are substituted into Eq. (25) to calculate  $\ln[P(\mathbf{Data} | Mdl_k)]$ . Subsequently, the  
 321 configuration with the largest value of  $\ln[P(\mathbf{Data} | Mdl_k)]$  is identified as the best-fit one. The  
 322 performance of the evidence estimator is dependent on the quality of the fitting of posterior

323 PDFs, and an inappropriate fitting can result in biased estimates. It is worth noting that  
324 Bayesian evidence can also be obtained as a by-product of BUS using the algorithms proposed  
325 by Betz et al. (2018a, 2018b). Comparative results showed that there was generally a strong  
326 correlation between the evidences derived from the suggested method and the BUS estimator.  
327 The suggested estimator can be used regardless of the specific BUS algorithm employed, and  
328 thus may serve as an alternative or supplement to the BUS estimator. Fig. 3 summarizes the  
329 flowchart of the proposed Bayesian framework.

#### 330 **4. Bayesian updating for three liquefaction-triggering procedures**

##### 331 **4.1. Dataset preparation and model configurations**

332 Based on the Bayesian framework, the SPT- and CPT-based case history databases for sandy  
333 soils compiled in Boulanger and Idriss (2014) are used to develop the region-specific SPT-BI12  
334 and CPT-RW98 models, respectively, whereas the database documented by Rollins et al. (2022)  
335 is adopted to develop the region-specific Vs-R22 models for gravelly soils. After excluding  
336 some undesirable data (e.g., regions with inadequate data and case histories with unclear  
337 liquefaction classification or unrealistically large  $F_{s0}$ ), the final datasets for SPT-BI12, CPT-  
338 RW98, and Vs-R22 contain 217, 234, and 153 case histories, respectively. Detailed exclusion  
339 criteria and dataset information are given in the Supplementary material. For each liquefaction-  
340 triggering procedure, 12 candidate model configurations (2 hyperparameter choices  $\times$  3  
341 population-level distributions  $\times$  2  $Q_p$  values) are considered in this study.

##### 342 **4.2. Bayesian updating results**

343 Fig. 4a-c display the logarithmic evidence (i.e.,  $\ln[P(\mathbf{Data}|Mdl_k)]$ ) values of the 12 candidate

344 configurations for the SPT-BI12, CPT-RW98, and Vs-R22 models, respectively. It is observed  
345 that the conventional lognormal distribution produces lower evidence than the other  
346 distributions in many cases, and is therefore generally not the best-fit population-level  
347 distribution. In addition, the evidence for HBM-4Hyp is generally higher than that for HBM-  
348 2Hyp, due to the more complete statistical representation by HBM-4Hyp. Considering the best-  
349 fit population-level distribution, the posterior  $E(\cdot)$  and  $Std(\cdot)$  of the parameters  $[\mu_\mu, \sigma_\mu, \sigma_c, c_1,$   
350  $c_2, \dots, c_r]$  or  $[\mu_\mu, \sigma_\mu, \mu_\sigma, \sigma_\sigma, c_1, c_2, \dots, c_r]$  for the updated SPT-BI12, CPT-RW98, and Vs-R22  
351 models are listed in Tables 2-4, respectively. Several observations are made as follows:

352 (1) The observed  $E(\mu_\mu) > 1$  indicates that the deterministic model tends to, on average,  
353 underestimate  $F_S$ . Also, the positive  $E(\sigma_\mu)$  and  $E(\sigma_\sigma)$  for different models illustrate the existence  
354 of the inter-region variability. Specifically, the inter-region variability for the Vs-R22 models  
355 (with  $E(\sigma_\mu)=0.95\sim 0.99$  and  $E(\sigma_\sigma)=0.71\sim 0.72$ ) is generally larger than that for the other models,  
356 whereas the SPT-BI12 models correspond to the smallest inter-region variability (with  
357  $E(\sigma_\mu)=0.11\sim 0.15$  and  $E(\sigma_\sigma)=0.54\sim 0.57$ ). Meanwhile,  $E(\mu_\sigma)$  reflects the intra-region variability,  
358 which appears to be the largest for the CPT-RW98 models (with  $E(\mu_\sigma)=0.81\sim 0.87$ ).

359 (2)  $E(c_i)$  and  $Std(c_i)$  exhibit notable region-to-region differences, which are in accordance  
360 with the positive  $E(\sigma_\mu)$  and  $E(\sigma_\sigma)$ , respectively. For example, the influenced region by the 1989  
361 Loma Prieta earthquake corresponds to remarkably larger  $E(c_i)$  than the other regions for the  
362 Vs-R22 models. This is because the calculated  $F_{50}$  values are consistently smaller than 1, which  
363 is opposite to the observed nonexistence of liquefaction phenomena in this region.

364 (3) The  $E(c_i)$  values for HBM-2Hyp and HBM-4Hyp are generally comparable. Yet, HBM-  
365 4Hyp produces larger  $Std(c_i)$ ; the reason is that the inter-region variability for HBM-4Hyp is

366 represented by the inter-region differences of both  $\mu_{c_i}$  and  $\sigma_{c_i}$ , enlarging the total uncertainty of  
 367  $c_i$ . This is consistent with the larger evidence for HBM-4Hyp (see Fig. 4).

368 (4) The dataset-specific  $Q_p$  calculated from Eq. (12) generally yields smaller  $E(\mu_\mu)$  and  $E(c_i)$   
 369 compared to  $Q_p = 0.456$ . This is because the calculated  $Q_p$  is consistently larger than 0.456 and  
 370 is related to higher population proportion of liquefaction occurrence.

371 To display the posterior distributions of model parameters, Fig. 5 presents the histograms  
 372 of  $[\mu_\mu, \sigma_\mu, \sigma_c]$  or  $[\mu_\mu, \sigma_\mu, \mu_\sigma]$  in the updated CPT-RW98 and SPT-BI12 models considering  
 373 different combinations of  $Q_p$  and hyperparameters. The uniformly distributed prior relative  
 374 frequency is also shown for comparison. It is observed that the model parameters after Bayesian  
 375 updating are distributed in considerably narrower ranges. This is more evident for the SPT-  
 376 BI12 models due to their weaker inter- and intra-region variabilities. Note that the best-fit  
 377 population-level distribution will be mainly used in subsequent analyses, with focus on  
 378 evaluating the effects of hyperparameter and  $Q_p$  selections.

### 379 4.3. Relationship between liquefaction probability and nominal factor of safety

380 Based on the updating results in Tables 2-4, the liquefaction probability  $P_{L0}$  can be calculated  
 381 for a region of interest by substituting  $F_{s0}$  (from Eq. (1)) as well as  $E(c_i)$  and  $\text{Std}(c_i)$  into Eq.  
 382 (4), which is expressed as:

$$383 \quad P_{L0} = \Phi \left[ \frac{-\ln E(c_i) + 0.5 \ln(1 + \text{Std}^2(c_i)/E^2(c_i)) - \ln F_{s0}}{\sqrt{\ln(1 + \text{Std}^2(c_i)/E^2(c_i))}} \right] \quad (29)$$

384 Eq. (29) represents an approximate simplified expression for the liquefaction probability  
 385 accounting for the region-specific variability. Therefore, the  $E(c_i)$  and  $\text{Std}(c_i)$  values provided  
 386 in Tables 2-4, derived using the laws of total expectation and variance, can be directly employed

387 without the requirement for posterior samples. An alternative MCS-based approach of  
388 averaging the liquefaction probability values obtained for individual posterior samples of  $\mu_{ci}$   
389 and  $\sigma_{ci}$  may provide more consistent estimates.

390 To illustrate the performance of the region-specific  $P_{L0}$  models, the probability mass  
391 function (PMF) of the number of predicted liquefied cases in each region is obtained and  
392 compared with the actual number of liquefied cases. Specifically, the sum of independent  
393 indicators of liquefaction (= 1) or non-liquefaction (= 0) for individual case histories in each  
394 region is assumed to follow a Poisson binomial distribution, where the occurrence probability  
395 of each indicator is specified as a case-history-specific  $P_{L0}$  value. In this study, PMF is obtained  
396 by the ‘poibin’ package in the **R** software (Hong 2013).

397 Fig. 6a and b display the PMF results for the region-specific SPT-BI12 and CPT-RW98  
398 models, respectively. It is seen that the actual number of liquefied cases is generally within the  
399 range of relatively large PMF. The HBM-4Hyp, which tends to predict more liquefied cases  
400 than HBM-2Hyp, is observed to better match the actual liquefaction observations. Meanwhile,  
401 using the dataset-specific  $Q_p$  calculated from Eq. (12) appears to partially improve the fitting  
402 performance of the CPT-RW98 models (e.g., in Loma Prieta and Canterbury regions), for  
403 which the dataset-specific  $Q_p = 0.649$  is considerably larger than the conventionally used  $Q_p =$   
404  $0.456$ . For the SPT-BI12 models, the PMFs for the two  $Q_p$  choices (0.461 and 0.456) are not  
405 significantly different. These graphical results are consistent with the quantitative ranking of  
406 model evidence in Fig. 5. Note that the PMF-based comparison between the actual number of  
407 liquefied cases and the HBM predictions does not consider the sample imbalance, although the  
408 model-to-model difference can be illustrated. One may assign a reasonable weighting factor

409 for a fairer comparison.

#### 410 **4.4. Comparison of region-specific models for different HBM configurations**

411 Using  $Q_p = 0.456$ , Fig. 7a and b present the  $P_{L0}$  versus  $F_{S0}$  trends for the region-specific SPT-  
412 BI12 and CPT-RW98 models, respectively. Although not shown herein, the predicted  $P_{L0}$   
413 would generally become larger with increasing  $Q_p$ . Several findings can be obtained. First,  $P_{L0}$   
414 decreases with increasing  $F_{S0}$  (or safety level), and  $P_{L0} > 0.5$  can still be observed for  $F_{S0} > 1$   
415 due to the uncertainty involved. Second, there are evident region-to-region differences of  $P_{L0}$ ,  
416 though they are relatively small for the SPT-based models. It is thus necessary to account for  
417 the inter-region variability of model bias. Third, HBM-4Hyp results in more notable inter-  
418 region difference of  $P_{L0}$  than HBM-2Hyp, which is mainly attributed to the larger inter-region  
419 difference of  $\text{Std}(c_i)$  in HBM-4Hyp. In contrast, the lumped-region modeling that neglects  
420 region-specific features (see Fig. 2a) would lead to inaccurate estimations.

421 Furthermore, Fig. 7c compares the  $P_{L0}$ - $F_{S0}$  relationships derived from different HBM  
422 configurations. The comparison is based on the CPT-RW98 and Vs-R22 models for regions  
423 not included in the calibration databases. The 4Hyp-based models generally yield greater  $P_{L0}$   
424 than the 2Hyp-based models at a given  $F_{S0}$ , especially for CPT-RW98. Additionally, the  
425 predicted  $P_{L0}$  becomes greater with increasing  $Q_p$ , and is also notably influenced by the  
426 selection of population-level distribution.

427 A more detailed comparison of different HBM configurations is conducted. The influence  
428 of population-level distribution is illustrated in Fig. 8a using the fixed [2Hyp,  $Q_p=0.456$ ].  
429 Specifically, the ratios of  $P_{L0}$  for the three candidate distributions to  $P_{L0}$  for the best-fit  
430 distribution at multiple  $F_{S0}$  levels are displayed, where the boxplot represents the distribution

431 of ratios for all regions. It is observed that the conventional lognormal distribution tends to  
432 produce smaller  $P_{L0}$  compared to the generally identified best-fit Trunc-Gumbel distribution,  
433 and the difference is likely to be enlarged with increasing  $F_{S0}$ . Therefore, calibrating the  
434 population-level distribution is necessary to avoid biased liquefaction potential assessment.

435 Using the best-fit population-level distribution, Fig. 8b presents the ratios of  $P_{L0}$  for three  
436 parameter configurations to the conventional configuration [2Hyp,  $Q_p=0.456$ ]. The ratio values  
437 are observed to be as large as 10 at  $F_{S0} \geq 2$ . Thus, the influence of parameter configuration is  
438 more significant than that of the population-level distribution. The HBM-4Hyp generally  
439 results in larger  $P_{L0}$  than HBM-2Hyp, and adopting dataset-specific  $Q_p$  tends to yield slightly  
440 larger  $P_{L0}$  compared to  $Q_p=0.456$ . Thus, the [2Hyp,  $Q_p=0.456$ ] configuration can be  
441 unconservative for statistical description of the inter- and intra-region variabilities.

## 442 **5. Implication of region-variability for liquefaction-triggering hazard assessment**

443 The region-specific SPT-IB12 and CPT-RW98 models are applied to Yuanlin (YL-B-3 profile)  
444 and Zhangbin Industrial Park (LW-A10 profile), respectively, which are located in Taiwan and  
445 were influenced by the 1999 Chi-Chi earthquake. It should be noted that the two profiles are  
446 not included in our HBM-based model updating. The in-situ geotechnical investigations are  
447 obtained from the database compiled by Dr. C. H. Juang ([https://cecas.clemson.edu/chichi/TW-](https://cecas.clemson.edu/chichi/TW-LIQ/In-situ-Test.htm)  
448 [LIQ/In-situ-Test.htm](https://cecas.clemson.edu/chichi/TW-LIQ/In-situ-Test.htm)). For the YL-B-3 profile, the corrected SPT blow counts  $(N_1)_{60}$  were  
449 obtained from the raw SPT blow counts using the model of Juang et al. (2005). The fines  
450 content (FC) and  $(N_1)_{60}$  of this profile are displayed in Fig. 9a. Also, the measured cone tip  
451 penetration resistance ( $q_c$ ) and sleeve friction ( $f_s$ ) of the LW-A10 profile are shown in Fig. 9b.

## 452 **5.1. Scenario-based liquefaction potential assessment**

453 The deterministic SPT-BI12 and CPT-RW98 procedures are adopted to estimate  $F_{S0}$   
454 considering earthquake magnitude ( $M$ ) of 7.6, where the recorded  $a_{\max}$  values are 0.18 g and  
455 0.12 g for the YL-B-3 and LW-A10 profiles, respectively (Juang et al. 2005; Ku et al. 2012).  
456 The updated region-specific models are then used to estimate  $P_{L0}$  using Eq. (29). The results  
457 of  $F_{S0}$  and  $P_{L0}$  are included in Fig. 9. The largest liquefaction probabilities in the YL-B-3 and  
458 LW-A10 profiles are observed at the depths of about 5 m and 10.5 m, respectively, and are both  
459 larger than 0.5. The results are in accordance with the actual occurrence of liquefaction at the  
460 two sites. Specifically, HBM-4Hyp predicts higher  $P_{L0}$  than HBM-2Hyp.

## 461 **5.2. Probabilistic method for liquefaction-triggering hazard assessment**

462 The performance-based earthquake engineering design (e.g., Juang et al. 2010; Rathje et al.  
463 2014) usually requires an estimate of the total probability of liquefaction subjected to all  
464 possible ground motions in an exposure period (e.g., 50 years). Therefore, a discrete  
465 integration-based probabilistic method is suggested herein to assess liquefaction-triggering  
466 hazard accounting for multiple ground motion levels.

467 This method convolves ground motion hazard with a liquefaction-triggering procedure,  
468 following the idea of probabilistic seismic demand analysis (e.g., Wang et al. 2021). The mean  
469 annual rate of liquefaction ( $\lambda_L$ ) is calculated as:

$$470 \lambda_L = \sum_i \sum_j \sum_k \sum_l P(F_S < 1 | M_i, a_{\max,j}, r_{d,k}, MSF_l) P(MSF_l | M_i, a_{\max,j}) P(r_{d,k} | M_i, a_{\max,j}) P(M_i, a_{\max,j})$$

471 (30)

472 where  $i, j, k,$  and  $l$  identify the discretized bins of  $M, a_{\max}, r_d,$  and  $MSF$ , respectively;  $P(F_S <$   
473  $1|\cdot)$  is calculated from  $P_{L0}$  model through Eq. (29);  $P(M_i, a_{\max,j})$  is the joint annual probability

474 of occurrence of  $M = M_i$  and  $a_{\max} = a_{\max,j}$ , and is obtained as:

$$475 \quad P(M_i, a_{\max,j}) = \sum_m P(M_i, R_m | a_{\max,j}) P(a_{\max,j}) \quad (31)$$

476 where  $R_m$  represents the  $m$ -th bin of source-to-site distance  $R$ ;  $P(M_i, R_m | a_{\max,j})$  is the joint annual  
 477 probability of  $M = M_i$  and  $R = R_m$  given  $a_{\max} = a_{\max,j}$ , and can be computed from the  
 478 disaggregation of seismic hazard (Baker et al. 2021);  $P(a_{\max,j})$  represents the annual probability  
 479 of  $a_{\max} = a_{\max,j}$ , which can be derived as (Rathje et al. 2014):

$$480 \quad P(a_{\max,j}) = \frac{\lambda(a_{\max,j-1}) - \lambda(a_{\max,j+1})}{2} \quad (32)$$

481 where  $\lambda(a_{\max,j-1})$  and  $\lambda(a_{\max,j+1})$  denote the mean annual rates of  $a_{\max} > a_{\max,j-1}$  and  $a_{\max} > a_{\max,j+1}$ ,  
 482 respectively.

483 The  $P(r_{d,k} | M_i, a_{\max,j})$  term in Eq. (30) represents the occurrence probability of  $r_d = r_{d,k}$   
 484 conditioned on  $M = M_i$  and  $a_{\max} = a_{\max,j}$ , and is statistically described to account for the  
 485 uncertainty in  $r_d$  estimation. In this study,  $r_d$  is assumed to follow a normal distribution with  
 486 the mean of  $\mu_{rd}$  and standard deviation of  $\sigma_{rd}$  (Cetin and Seed 2004), which are often provided  
 487 by empirical models (e.g., Moss et al. 2006; Lasley et al. 2016). Given the uncertainty in  $MSF$   
 488 estimation as well (e.g., Green et al. 2020),  $\ln MSF$  is also assumed to be a normal variable with  
 489 the mean of  $\mu_{\ln MSF}$  and standard deviation of  $\sigma_{\ln MSF}$ . Specifically,  $\mu_{rd}$  and  $\mu_{\ln MSF}$  denote the  
 490 deterministic estimates of  $r_d$  and logarithmic  $MSF$ , respectively (see Supplementary material).  
 491 Both  $\sigma_{rd}$  and  $\sigma_{\ln MSF}$  are assumed to be 0.15 (Lasley et al. 2016; Green et al. 2020), unless stated  
 492 otherwise. Sensitivity analyses of the two parameters will be separately discussed later.

493 Under the assumption of Poisson's process (e.g., Kramer and Mayfield 2007; Wang et al.  
 494 2022),  $\lambda_L$  can be used to estimate the liquefaction probability ( $P_{L,T}$ ) in an exposure period  $T$ :

$$495 \quad P_{L,T} = 1 - \exp(-\lambda_L T) \quad (33)$$

496 In other words,  $P_{L,T}$  represents the total probability of liquefaction accounting for the uncertain  
497 occurrence of earthquake during a target period.

### 498 **5.3. Liquefaction-triggering hazard analysis results**

499 For the purpose of illustrating the proposed models, the soil profiles are assumed to be located  
500 at a distance of 10 km from the center of a 200 km-long strike-slip linear fault. The seismicity  
501 of the hypothetical source is described by a bounded Gutenberg-Richter relationship:

$$502 \log_{10}(\lambda_M) = 4.3 - M \quad (34)$$

503 where  $\lambda_M$  is the mean annual rate that earthquake magnitude exceeds  $M$ . The central values of  
504 discretized  $M$  bins range from 5 to 8 with increments of 0.1. For each bin of  $M$ , earthquake  
505 rupture occurs uniformly along the fault with the rupture length  $L$  as a function of  $M$  (Wells  
506 and Coppersmith 1994). The source-site relationship is schematically illustrated in Fig. 10.

507 The seismic hazard for  $a_{\max}$  is derived using the ground-motion model proposed by  
508 Campbell and Bozorgnia (2014). Since the shear-wave velocities of the YL-B-3 and LW-A10  
509 profiles are not available, the average upper 30 m shear-wave velocity ( $V_{S30}$ ) required by the  
510 ground-motion model is estimated as follows: (1) the raw SPT-N values are transformed into  
511  $V_s$  values using the empirical model developed by Xiao et al. (2021); (2) the depth-dependent  
512  $V_s$  values are used to estimate the average upper 15 m shear-wave velocity ( $V_{S15}$ ), which is  
513 then used in the  $V_{S30}$ - $V_{S15}$  relationship established by Boore et al. (2011).

514 First, the profiles are assumed to be in regions influenced by the 1999 Chi-Chi earthquake.  
515 Fig. 11a displays  $\lambda_L$  at different depths and the variation of  $P_{L,T}$  at the depth of 5.8 m with  
516 exposure period for the YL-B-3 profile. Similarly, the results for the LW-A10 profile are  
517 presented in Fig. 11b. Regarding the [2Hyp,  $Q_p=0.456$ ] configuration, largest  $\lambda_L$  values of

518 around 0.0055 and 0.0186 are observed at the 5.8 m and 10.5 m depths in the YL-B-3 and LW-  
519 A10 profiles, respectively. In contrast, the 4Hyp modeling results in considerably larger  $\lambda_L$  and  
520  $P_{L,T}$  because of the larger uncertainty considered in the model. Meanwhile,  $\lambda_L$  and  $P_{L,T}$  are  
521 higher for the larger calculated  $Q_p$ , though the influence on the SPT-BI12 models (for which  
522 the calculated  $Q_p$  is comparable with 0.456) is small. As also shown in the figures,  $P_{L,T}$   
523 increases with increasing exposure period; thus, a soil deposit with  $P_{L,T} < 0.5$  for relatively  
524 short  $T$  can still be likely to liquefy.

525 To compare the effects of choosing region-specific models, the profiles are then assumed  
526 to be located in different regions, subjected to the same earthquake source, path, and site  
527 conditions. This ensures that the difference of results only originates from the liquefaction-  
528 triggering models. Fig. 12a and b display the results of  $\lambda_L$  from different region-specific SPT-  
529 BI12 models (for the depth of 5.8 m in YL-B-3 profile) and CPT-RW98 models (for the depth  
530 of 10.5 m in LW-A10 profile), respectively. Besides, the lumped-region modeling is also  
531 included for comparison. It is seen that the  $\lambda_L$  values for different regions are notably different,  
532 and can considerably differ from the results for the lumped-region modeling.

533 Furthermore, the effects of region features on the liquefaction-triggering hazard analysis  
534 is studied by quantifying the ratio of  $P_{L,T}$  between region-specific and lumped-region models,  
535 namely  $[P_{L,T}(\text{region,para})/P_{L,T}(\text{lumped},Q_p=0.456)]$ . Fig. 13a and b present the boxplots of the  
536  $P_{L,T}$  ratios in different regions for the SPT- and CPT-based models, respectively, in which three  
537 exposure periods and four parameter configurations are considered. Except for the [CPT-2Hyp,  
538  $Q_p=0.456$ ] modeling, the  $P_{L,T}$  ratio values are found to be generally larger than 1, and exceed  
539 1.3 in a few cases. Thus, the lumped-region modeling could underestimate the liquefaction

540 probability by more than 30%. The distribution of the  $P_{L,T}$  ratio for the SPT-BI12 models  
541 appears to be insensitive to the exposure period, while the relative difference between the  
542 region-specific and lumped-region models is larger for shorter exposure period. In particular,  
543 the relative difference is more evident for HBM-4Hyp as compared to HBM-2Hyp. The above  
544 results thus indicate the importance of refined region-feature representation in liquefaction-  
545 triggering hazard assessment.

#### 546 **5.4. Effects of $\sigma_{rd}$ and $\sigma_{\ln MSF}$ on liquefaction-triggering hazard analysis**

547 As mentioned previously,  $\sigma_{rd}$  and  $\sigma_{\ln MSF}$  are greater than 0, due to the uncertainty in predictions  
548 of  $r_d$  and  $MSF$  (e.g., from ground-motion variability and parametric fitting error). The  $\sigma_{rd}$  versus  
549 depth trends reported in previous studies (Cetin and Seed 2004; Moss et al. 2006; Lasley et al.  
550 2016; Green et al. 2020) are plotted in Fig. 14a. It is seen that  $\sigma_{rd}$  increases with depth, reaching  
551 around 0.15-0.2 at depth larger than 15 m. In contrast,  $\sigma_{\ln MSF}$  has been reported to be about 0.15  
552 (Green et al. 2020). To investigate the effects of  $\sigma_{rd}$  and  $\sigma_{\ln MSF}$ , the  $P_{L,50}$  values along the LW-  
553 A10 profile for different  $\sigma_{rd}$  and  $\sigma_{\ln MSF}$  values (0, 0.1, 0.2, and 0.5) are displayed in Fig. 14b  
554 and c, respectively. The results are derived by the [CPT-2Hyp,  $Q_p=0.456$ ] model for the 1999  
555 Chi-Chi earthquake. It is observed that the effects of standard deviations on  $P_{L,T}$  are negligibly  
556 small when  $\sigma_{rd}$  and  $\sigma_{\ln MSF}$  are not larger than 0.2, yet become relatively large when increasing  
557  $\sigma_{rd}$  and  $\sigma_{\ln MSF}$  to 0.5.

## 558 **6. Discussion**

559 To examine the adequacy of the data size used to calibrate models, an additional set of CPT-  
560 4Hyp models is developed by increasing the current minimum data size for each region from

561 4 to 8. After removing two regions (i.e., 1983 Borah Peak and 2011 Tohoku) with less than 8  
562 case histories, 8 regions are retained for the model re-development. The resulting  $P_{L0}-F_{S0}$   
563 relationships are compared with those derived from the original database in Fig. 15a. The  
564 results for the two minimum data sizes (i.e., two datasets) are observed to be comparable,  
565 suggesting that the data size considered is generally adequate for the HBM-4Hyp calibration.

566 Regarding the SPT-BI12 procedure, Fig. 15b compares  $\text{Std}(c)$  versus  $R_{\text{FT}}$  trends for four  
567 Bayesian modeling choices, where  $R_{\text{FT}}$  is defined as the ratio of false to true predictions (i.e.,  
568  $(\text{FP}+\text{FN})/(\text{TP}+\text{TN})$  in Confusion Matrix) from a deterministic liquefaction-triggering  
569 procedure and data for each region. It can be seen that the  $\text{Std}(c)$  values generally increase with  
570 increasing  $R_{\text{FT}}$ , implying that the intra-region variability of bias factor can be roughly measured  
571 by  $R_{\text{FT}}$ . Additionally, HBM-4Hyp results in much smaller  $\text{Std}(c)$  than the independent-regions  
572 modeling. Thus, HBM-4Hyp differs significantly from the independent-regions modeling that  
573 fully neglects information borrowing across regions. In contrast, HBM-2Hyp leads to a  
574 negligible regional variation of  $\text{Std}(c)$ , which is similar to the lumped-regions modeling with a  
575 common  $\text{Std}(c)$ . The comparative results demonstrate the capability of HBM-4Hyp to capture  
576 the inter-region variability of bias variance.

577 In addition to the method suggested, another major contribution of this study is to  
578 demonstrate that there can be large region-to-region differences in the standard deviations of  
579 model bias factors and HBM-2Hyp with a common variance could underestimate liquefaction  
580 hazard. In this case, HBM-4Hyp is advantageous over HBM-2Hyp for characterizing the  
581 region-specific statistics. Yet, a notable challenge in implementing HBM-4Hyp is the increased  
582 data size requirement for calibration. Therefore, it is recommended that practitioners should

583 assess the plausibility of region-invariant variance and the adequacy of data size before making  
584 a choice between HBM-4Hyp and HBM-2Hyp, both of which are considered in this study. This  
585 work represents a step towards refining the representation of the inter-region variability of  
586 model bias. However, further efforts are necessary to continue advancing research in the  
587 regionalization of liquefaction-triggering models (e.g., Zhang et al. 2016, 2020; Green 2022),  
588 taking into account the continuously evolving global and regional liquefaction databases (e.g.,  
589 Brandenberg et al. 2020; Geyin et al. 2021).

## 590 **7. Conclusions**

591 This paper presented a Bayesian framework to incorporate the inter-region and intra-region  
592 variabilities of model bias into liquefaction-triggering procedures. The BUS approach with SuS  
593 is utilized for posterior sampling and the Gaussian copula-based Bayesian evidence is  
594 quantified for model selection. A new four-hyperparameter hierarchical Bayesian model  
595 (HBM-4Hyp) was suggested, which assumes the means ( $\mu_{ci}$ ) or standard deviations ( $\sigma_{ci}$ ) of the  
596 bias factor ( $c_i$ ) are region-specific, but are controlled by a population-level distribution. This  
597 model was systematically compared with HBM-2Hyp (with region-invariant  $\sigma_{ci}$ ). Meanwhile,  
598 three candidate distributions were adopted for the population-level statistical modeling.  
599 Scenario-based and probabilistic liquefaction-triggering hazard analyses were conducted to  
600 investigate the implication of region-feature modeling. The main findings drawn from this  
601 study are summarized as follows:

602 (1) The HBM-4Hyp requires around twice the number of parameters for HBM-2Hyp,  
603 leading to a high-dimensional Bayesian updating problem. This can be efficiently addressed by  
604 the BUS approach with SuS, and the benefits of such would be more notable with increasing

605 the number of regions compiled in future global databases. Meanwhile, the Bayesian evidence  
606 is shown to be effective for model comparison, and a HBM configuration with higher evidence  
607 is generally observed to better match liquefaction observations.

608 (2) The proposed Bayesian framework is utilized for incorporating the region-features into  
609 three liquefaction-triggering procedures (SPT-BI12, CPT-RW98, and Vs-R22), covering  
610 liquefaction-susceptible sandy and gravelly soil types. The coefficients of the updated region-  
611 specific probabilistic liquefaction models are provided for potential applications.

612 (3) There exist notable inter-region variabilities of  $\mu_{ci}$  and  $\sigma_{ci}$  especially for the CPT- and Vs-  
613 based models. Compared to HBM-2Hyp, HBM-4Hyp tends to yield higher Bayesian evidence  
614 and larger total standard deviation of  $c_i$ . When the calculated  $Q_p$  is considerably different from  
615 the conventional value of 0.456, this dataset-specific  $Q_p$  is likely to enlarge Bayesian evidence,  
616 and the predicted liquefaction probabilities are found to increase with increasing  $Q_p$ .

617 (4) The lognormal distribution may not be optimal for the population-level statistical  
618 descriptions in HBM, as illustrated by the lower evidence compared to the other candidate  
619 distributions, and could considerably underestimate liquefaction potential in some cases.

620 (5) A discrete integration-based probabilistic method is suggested to implement the  
621 developed models for liquefaction-triggering hazard assessment considering multiple ground-  
622 motion levels. Neglecting the region-variability is likely to underestimate liquefaction hazard,  
623 and HBM-4Hyp generally yields higher liquefaction hazard than HBM-2Hyp.

## 624 **Acknowledgment**

625 The work presented in this paper is financially supported by the Research Grants Council of  
626 Hong Kong Special Administrative Region (Project No. 15222021).

627 **Supplementary material**

628 Supplementary data related to the mathematical expressions and liquefaction datasets for the  
629 three liquefaction-triggering procedures are available with the article. Other data generated or  
630 analyzed during this study are available from the corresponding author upon reasonable request.

631 **Competing interests**

632 The authors declare there are no competing interests.

633 **References**

634 Au, S. K., Beck, J. L. (2001). Estimation of small failure probabilities in high dimensions by  
635 subset simulation. *Probabilistic Engineering Mechanics*, 16(4), 263-277.

636 Baker, J., Bradley, B., Stafford, P. (2021). *Seismic hazard and risk analysis*. Cambridge  
637 University Press.

638 Betz, W., Beck, J. L., Papaioannou, I., Straub, D. (2018a). Bayesian inference with reliability  
639 methods without knowing the maximum of the likelihood function. *Probabilistic  
640 Engineering Mechanics*, 53, 14-22.

641 Betz, W., Papaioannou, I., Beck, J. L., Straub, D. (2018b). Bayesian inference with subset  
642 simulation: strategies and improvements. *Computer Methods in Applied Mechanics and  
643 Engineering*, 331, 72-93.

644 Boore, D. M., Thompson, E. M., Cadet, H. (2011). Regional correlations of VS 30 and  
645 velocities averaged over depths less than and greater than 30 meters. *Bulletin of the  
646 Seismological Society of America*, 101(6), 3046-3059.

647 Boulanger, R. W., Idriss, I. M. (2012). Probabilistic standard penetration test-based  
648 liquefaction-triggering procedure. *Journal of Geotechnical and Geoenvironmental*

649 Engineering, 138(10), 1185-1195.

650 Boulanger, R. W., Idriss, I. M. (2014). CPT and SPT based liquefaction triggering  
651 procedures. Report No. UCD/CGM.-14, 1.

652 Bozorgzadeh, N., Harrison, J. P., Escobar, M. D. (2019). Hierarchical Bayesian modelling of  
653 geotechnical data: application to rock strength. *Géotechnique*, 69(12), 1056-1070.

654 Brandenberg, S. J., Zimmaro, P., Stewart, J. P., et al. (2020). Next-generation liquefaction  
655 database. *Earthquake Spectra*, 36(2), 939-959.

656 Campbell, K. W., Bozorgnia, Y. (2014). NGA-West2 ground motion model for the average  
657 horizontal components of PGA, PGV, and 5% damped linear acceleration response spectra.  
658 *Earthquake Spectra*, 30(3), 1087-1115.

659 Cao, Z., Wang, Y. (2014). Bayesian model comparison and selection of spatial correlation  
660 functions for soil parameters. *Structural Safety*, 49, 10-17.

661 Cao, Z. J., Zheng, S., Li, D. Q., Phoon, K. K. (2019). Bayesian identification of soil stratigraphy  
662 based on soil behaviour type index. *Canadian Geotechnical Journal*, 56(4), 570-586.

663 Cetin, K. O., Der Kiureghian, A., Seed, R. B. (2002). Probabilistic models for the initiation of  
664 seismic soil liquefaction. *Structural safety*, 24(1), 67-82.

665 Cetin, K. O., Seed, R. B. (2004). Nonlinear shear mass participation factor ( $r_d$ ) for cyclic shear  
666 stress ratio evaluation. *Soil Dynamics and Earthquake Engineering*, 24(2), 103-113.

667 Ching, J., Phoon, K. K., Ho, Y. H., Weng, M. C. (2021a). Quasi-site-specific prediction for  
668 deformation modulus of rock mass. *Canadian Geotechnical Journal*, 58(7), 936-951.

669 Ching, J., Wu, S., Phoon, K. K. (2021b). Constructing quasi-site-specific multivariate  
670 probability distribution using hierarchical Bayesian model. *Journal of Engineering*

671 Mechanics, 147(10), 04021069.

672 Deng, Z. P., Pan, M., Niu, J. T., Jiang, S. H. (2022). Full probability design of soil slopes  
673 considering both stratigraphic uncertainty and spatial variability of soil  
674 properties. *Bulletin of Engineering Geology and the Environment*, 81(5), 195.

675 Duan, W., Zhao, Z., Cai, G., Pu, S., Liu, S., Dong, X. (2022). Evaluating model uncertainty of  
676 an in situ state parameter-based simplified method for reliability analysis of liquefaction  
677 potential. *Computers and Geotechnics*, 151, 104957.

678 Geyin, M., Maurer, B. W., Bradley, B. A., Green, R. A., van Ballegooy, S. (2021). CPT-based  
679 liquefaction case histories compiled from three earthquakes in Canterbury, New Zealand.  
680 *Earthquake Spectra*, 37(4), 2920-2945.

681 Green, R. A. (2022). Regionalization of Liquefaction Triggering Models. In *Conference on  
682 Performance-based Design in Earthquake. Geotechnical Engineering* (pp. 437-451).  
683 Cham: Springer International Publishing.

684 Green, R. A., Bommer, J. J., Stafford, P. J., et al. (2020). Liquefaction hazard in the Groningen  
685 region of the Netherlands due to induced seismicity. *Journal of Geotechnical and  
686 Geoenvironmental Engineering*, 146(8), 04020068.

687 Guan, Z., Wang, Y. (2022). Quantifying reliability of liquefaction severity map developed from  
688 sparse cone penetration tests. *Canadian Geotechnical Journal*, 60(5), 623-641.

689 Hong, Y. (2013). On computing the distribution function for the Poisson binomial distribution.  
690 *Computational Statistics & Data Analysis*, 59, 41-51.

691 Hu, J., Liu, H. (2019). Bayesian network models for probabilistic evaluation of earthquake-  
692 induced liquefaction based on CPT and Vs databases. *Engineering Geology*, 254, 76-88.

693 Jiang, S. H., Papaioannou, I., Straub, D. (2018). Bayesian updating of slope reliability in  
694 spatially variable soils with in-situ measurements. *Engineering Geology*, 239, 310-320.

695 Juang, C. H., Jiang, T., Andrus, R. D. (2002). Assessing probability-based methods for  
696 liquefaction potential evaluation. *Journal of Geotechnical and Geoenvironmental*  
697 *Engineering*, 128(7), 580-589.

698 Juang, C. H., Ching, J., Luo, Z., Ku, C. S. (2012). New models for probability of liquefaction  
699 using standard penetration tests based on an updated database of case histories.  
700 *Engineering Geology*, 133, 85-93.

701 Juang, C. H., Ou, C. Y., Lu, C. C., Luo, Z. (2010). Probabilistic framework for assessing  
702 liquefaction hazard at a given site in a specified exposure time using standard penetration  
703 testing. *Canadian Geotechnical Journal*, 47(6), 674-687.

704 Juang, C. H., Yang, S. H., Yuan, H., Fang, S. Y. (2005). Liquefaction in the Chi-Chi earthquake-  
705 effect of fines and capping non-liquefiable layers. *Soils and foundations*, 45(6), 89-101.

706 Kramer, S. L., Mayfield, R. T. (2007). Return period of soil liquefaction. *Journal of*  
707 *Geotechnical and Geoenvironmental Engineering*, 133(7), 802-813.

708 Ku, C. S., Juang, C. H., Chang, C. W., Ching, J. (2012). Probabilistic version of the Robertson  
709 and Wride method for liquefaction evaluation: development and application. *Canadian*  
710 *Geotechnical Journal*, 49(1), 27-44.

711 Lasley, S. J., Green, R. A., Rodriguez-Marek, A. (2016). New stress reduction coefficient  
712 relationship for liquefaction triggering analyses. *Journal of Geotechnical and*  
713 *Geoenvironmental Engineering*, 142(11), 06016013.

714 Liu, W. F., Leung, Y. F. (2018). Spatial variability of saprolitic soil properties and relationship

715 with joint set orientation of parent rock: Insights from cases in Hong Kong. *Engineering*  
716 *Geology*, 246, 36-44.

717 Lo, M. K., Leung, Y. F. (2019). Bayesian updating of subsurface spatial variability for improved  
718 prediction of braced excavation response. *Canadian Geotechnical Journal*, 56(8), 1169-  
719 1183.

720 Maurer, B. W., Green, R. A., Van Ballegooy, S., Wotherspoon, L. (2019). Development of  
721 region-specific soil behavior type index correlations for evaluating liquefaction hazard in  
722 Christchurch, New Zealand. *Soil Dynamics and Earthquake Engineering*, 117, 96-105.

723 Moss, R. E., Seed, R. B., Kayen, R. E., Stewart, J. P., Der Kiureghian, A., Cetin, K. O. (2006).  
724 CPT-based probabilistic and deterministic assessment of in situ seismic soil liquefaction  
725 potential. *Journal of Geotechnical and Geoenvironmental Engineering*, 132(8), 1032-1051.

726 Papaioannou, I., Betz, W., Zwirgmaier, K., Straub, D. (2015). MCMC algorithms for subset  
727 simulation. *Probabilistic Engineering Mechanics*, 41, 89-103.

728 Rathje, E. M., Wang, Y., Stafford, P. J., Antonakos, G., Saygili, G. (2014). Probabilistic  
729 assessment of the seismic performance of earth slopes. *Bulletin of Earthquake*  
730 *Engineering*, 12, 1071-1090.

731 Robertson, P. K. (2022). Evaluation of flow liquefaction and liquefied strength using the cone  
732 penetration test: An update. *Canadian Geotechnical Journal*, 59(4), 620-624.

733 Robertson, P. K., Wride, C. E. (1998). Evaluating cyclic liquefaction potential using the cone  
734 penetration test. *Canadian Geotechnical Journal*, 35(3), 442-459.

735 Rollins, K. M., Roy, J., Athanasopoulos-Zekkos, A., et al. (2022). A new Vs-based liquefaction-  
736 triggering procedure for gravelly soils. *Journal of Geotechnical and Geoenvironmental*

737           Engineering, 148(6), 04022040.

738   Seed, H. B., Idriss, I. M. (1971). Simplified procedure for evaluating soil liquefaction potential.  
739           Journal of the Soil Mechanics and Foundations Division, 97(9), 1249-1273.

740   Straub, D., Papaioannou, I. (2015). Bayesian updating with structural reliability methods.  
741           Journal of Engineering Mechanics, 141(3), 04014134.

742   Tabarroki, M., Ching, J., Yuan, S. H., Phoon, K. K., Teng, F. (2024). Data-driven hierarchical  
743           Bayesian model for predicting wall deflections in deep excavations in clay. Computers  
744           and Geotechnics, 168, 106135.

745   Tang, X. S., Wang, M. X., Li, D. Q. (2020). Modeling multivariate cross-correlated  
746           geotechnical random fields using vine copulas for slope reliability analysis. Computers  
747           and Geotechnics, 127, 103784.

748   Tao, Y., Phoon, K. K., Sun, H., Cai, Y. (2023). Hierarchical Bayesian model for predicting  
749           small-strain stiffness of sand. Canadian Geotechnical Journal. [https://doi.org/10.1139/cgj-](https://doi.org/10.1139/cgj-2022-0598)  
750           2022-0598.

751   Tian, M., Li, D. Q., Cao, Z. J., Phoon, K. K., Wang, Y. (2016). Bayesian identification of  
752           random field model using indirect test data. Engineering Geology, 210, 197-211.

753   Wang, M. X., Tang, X. S., Li, D. Q., Qi, X. H. (2020). Subset simulation for efficient slope  
754           reliability analysis involving copula-based cross-correlated random fields. Computers and  
755           Geotechnics, 118, 103326.

756   Wang, M. X., Li, D. Q., Du, W. (2021). Probabilistic seismic displacement hazard assessment  
757           of earth slopes incorporating spatially random soil parameters. Journal of Geotechnical  
758           and Geoenvironmental Engineering, 147(11), 04021119.

759 Wang, M. X., Li, D. Q., Liu, Y., Du, W. (2022). Probabilistic decoupled approach to estimate  
760 seismic rotational displacements of flexible slopes considering depth-dependent soil  
761 variability. *Acta Geotechnica*, 17(4), 1551-1567.

762 Wells, D. L., Coppersmith, K. J. (1994). New empirical relationships among magnitude,  
763 rupture length, rupture width, rupture area, and surface displacement. *Bulletin of the*  
764 *seismological Society of America*, 84(4), 974-1002.

765 Wu, S., Ching, J., Phoon, K. K. (2022). Quasi-site-specific soil property prediction using a  
766 cluster-based hierarchical Bayesian model. *Structural Safety*, 99, 102253.

767 Xiao, S., Zhang, J., Ye, J., Zheng, J. (2021). Establishing region-specific N-Vs relationships  
768 through hierarchical Bayesian modeling. *Engineering Geology*, 287, 106105.

769 Youd, T. L., Idriss, I. M., Andrus, R. D., et al. (2001). Liquefaction resistance of soils: summary  
770 report from the 1996 NCEER and 1998 NCEER/NSF workshops on evaluation of  
771 liquefaction resistance of soils. *Journal of Geotechnical and Geoenvironmental*  
772 *Engineering*, 127(4), 297-313.

773 Yuen, K.V. (2010). *Bayesian Methods for Structural Dynamics and Civil Engineering*. John  
774 Wiley & Sons (Asia) Pte Ltd, Singapore.

775 Zhang, J., Li, J. P., Zhang, L. M., Huang, H. W. (2014). Calibrating cross-site variability for  
776 reliability-based design of pile foundations. *Computers and Geotechnics*, 62, 154-163.

777 Zhang, J., Juang, C. H., Martin, J. R., Huang, H. W. (2016). Inter-region variability of  
778 Robertson and Wride method for liquefaction hazard analysis. *Engineering Geology*, 203,  
779 191-203.

780 Zhang, J., Xiao, S., Huang, H., Zhou, J. (2020). Calibrating a standard penetration test based

781 method for region-specific liquefaction potential assessment. *Bulletin of Engineering*  
782 *Geology and the Environment*, 79, 5185-5204.

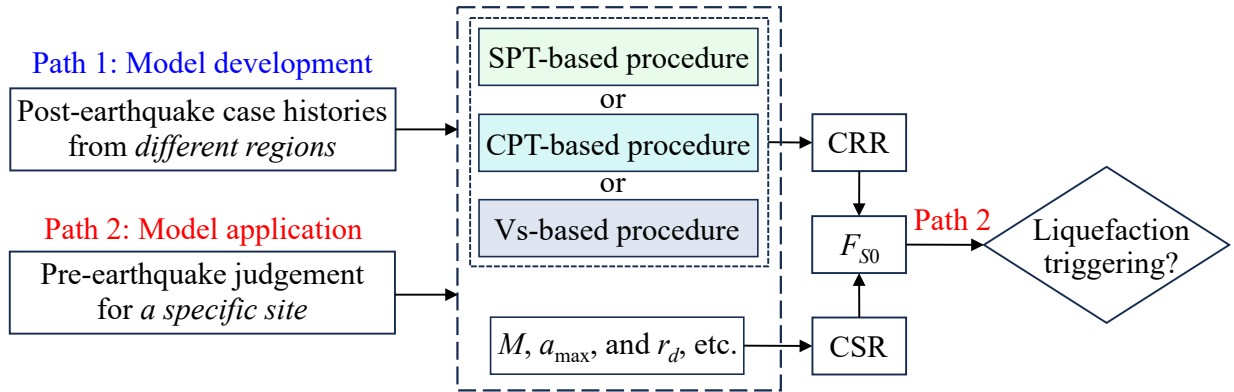
783 Zhang, J., Yang, S., Zhang, L. L., Zhou, M. L. (2022). Bayesian estimation of soil-water  
784 characteristic curves. *Canadian Geotechnical Journal*, 59(4), 569-582.

785

786

### List of Figures

787

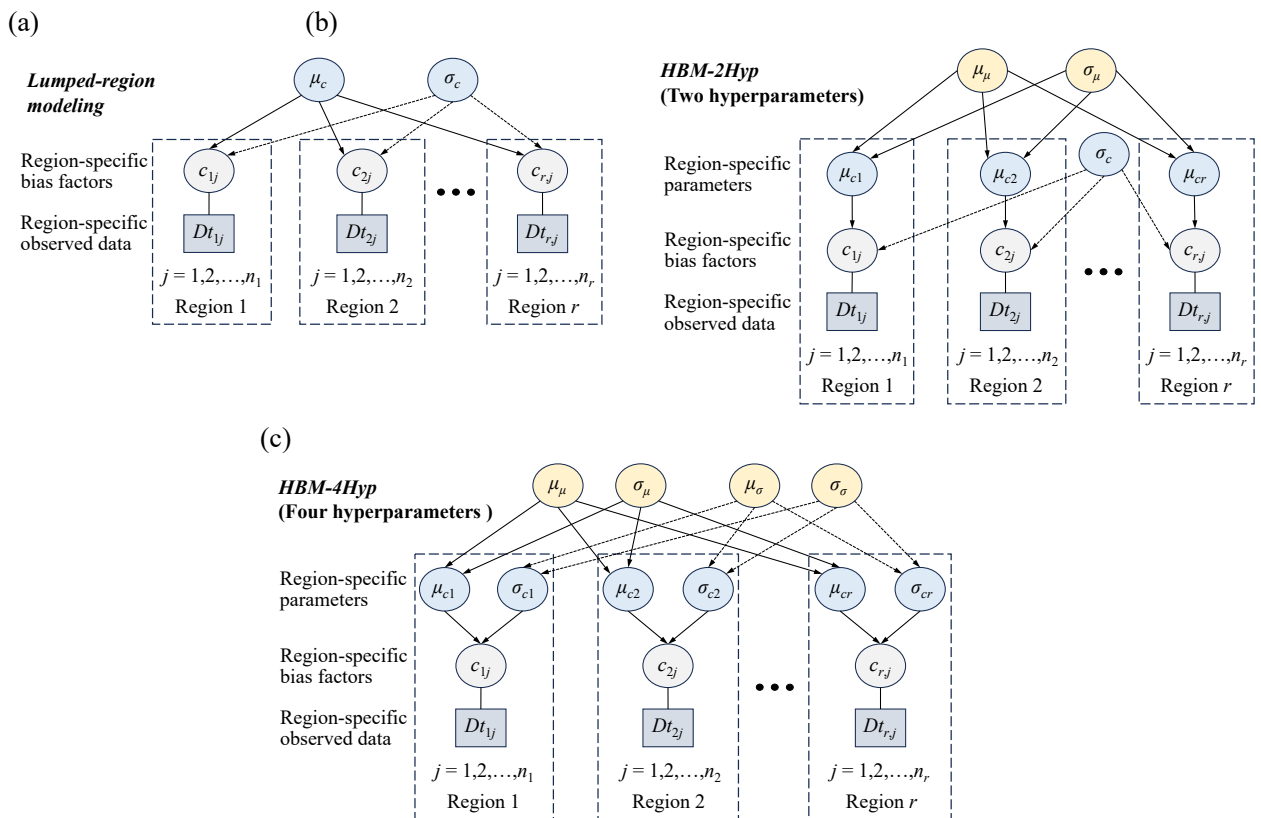


788

789 **Fig. 1.** Introduction to the development and application of conventional liquefaction-triggering assessment  
790 procedures.

791

792



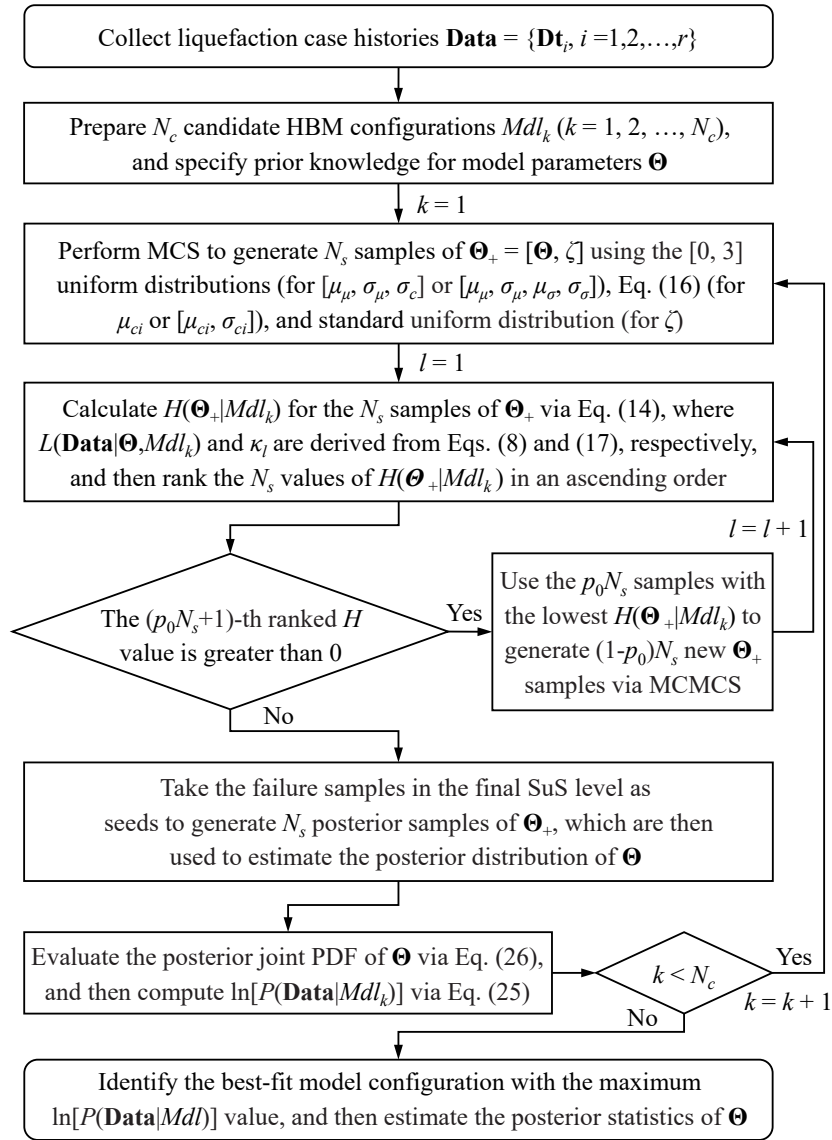
793

794

795

796 **Fig. 2.** Schematic illustrations of (a) lumped-region modeling, (b) HBM-2Hyp, and (c) HBM-4Hyp for  
797 updating liquefaction-triggering procedures.

798



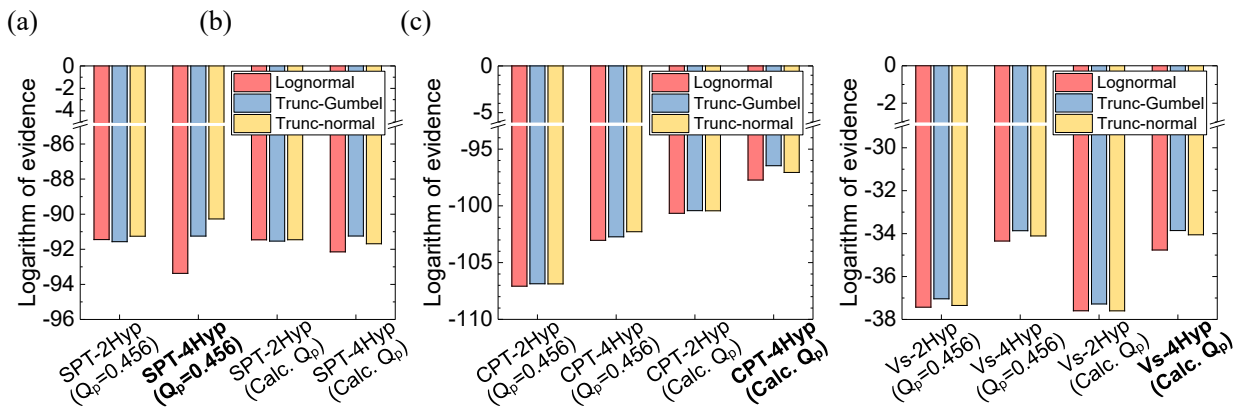
799

800

**Fig. 3.** Flowchart of the BUS-powered hierarchical Bayesian updating framework.

801

802



803

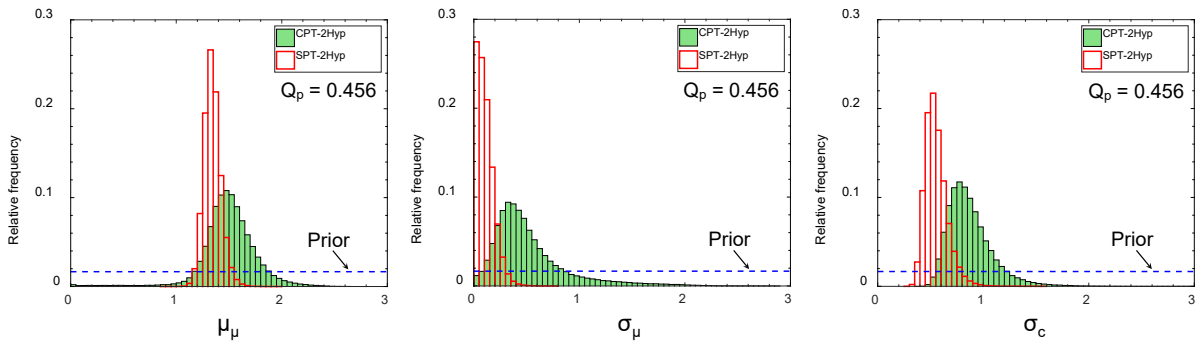
804

**Fig. 4.** Model evidence values for different HBM configurations of updated (a) SPT-BI12, (b) CPT-RW98,

805

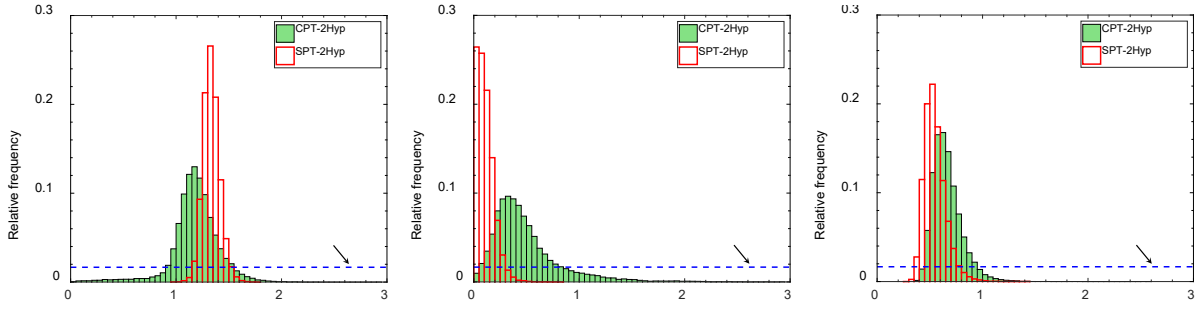
and (c) Vs-R22 models. The best-fit configuration is indicated as bold on the x-axis.

806 (a)



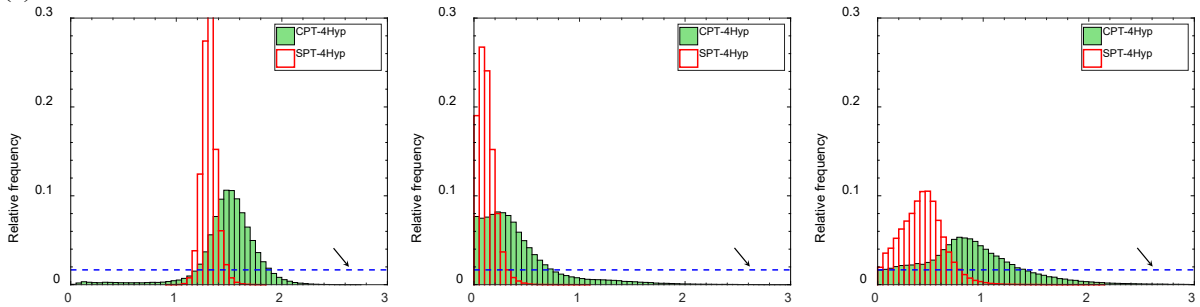
807

808 (b)



809

810 (c)

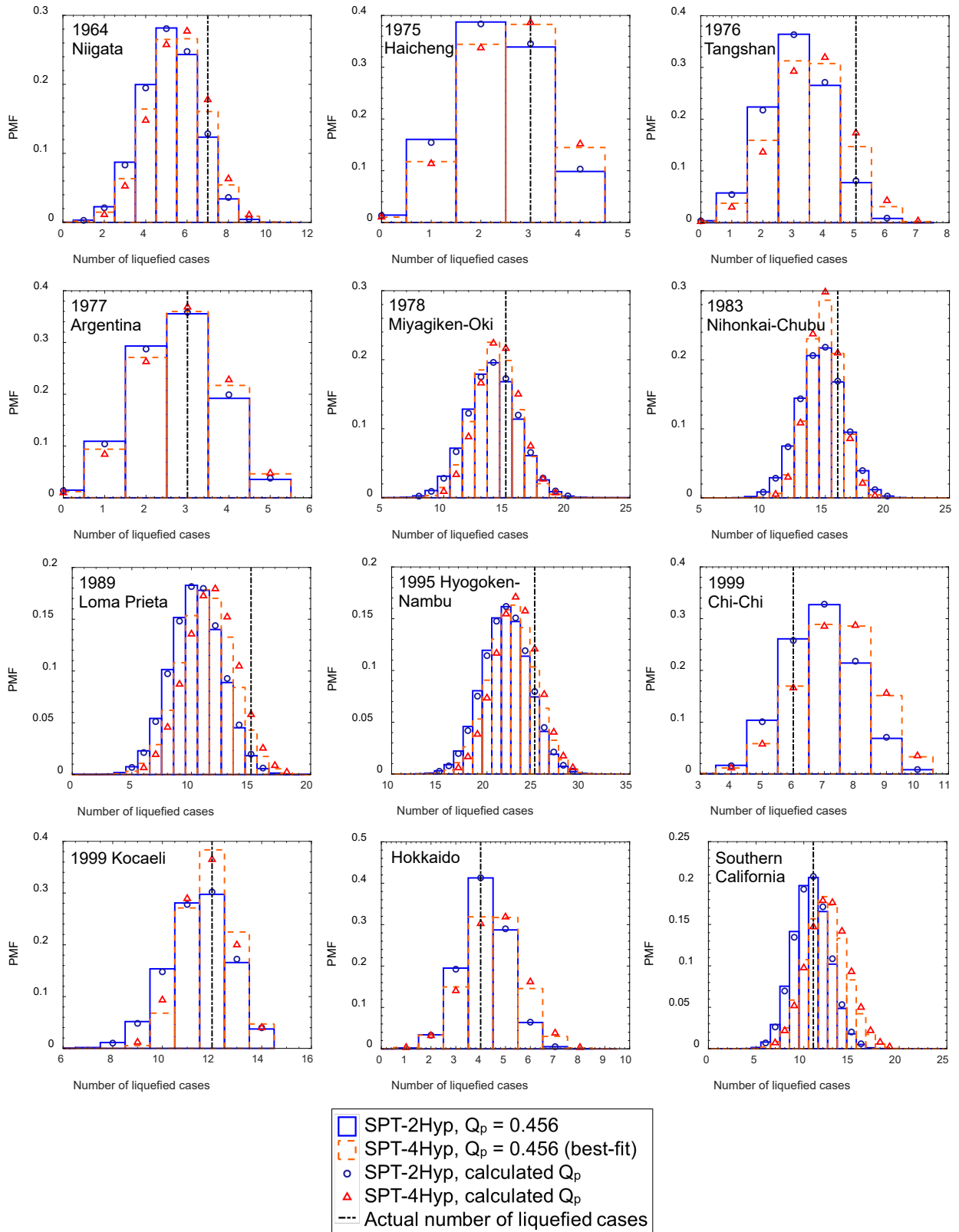


811

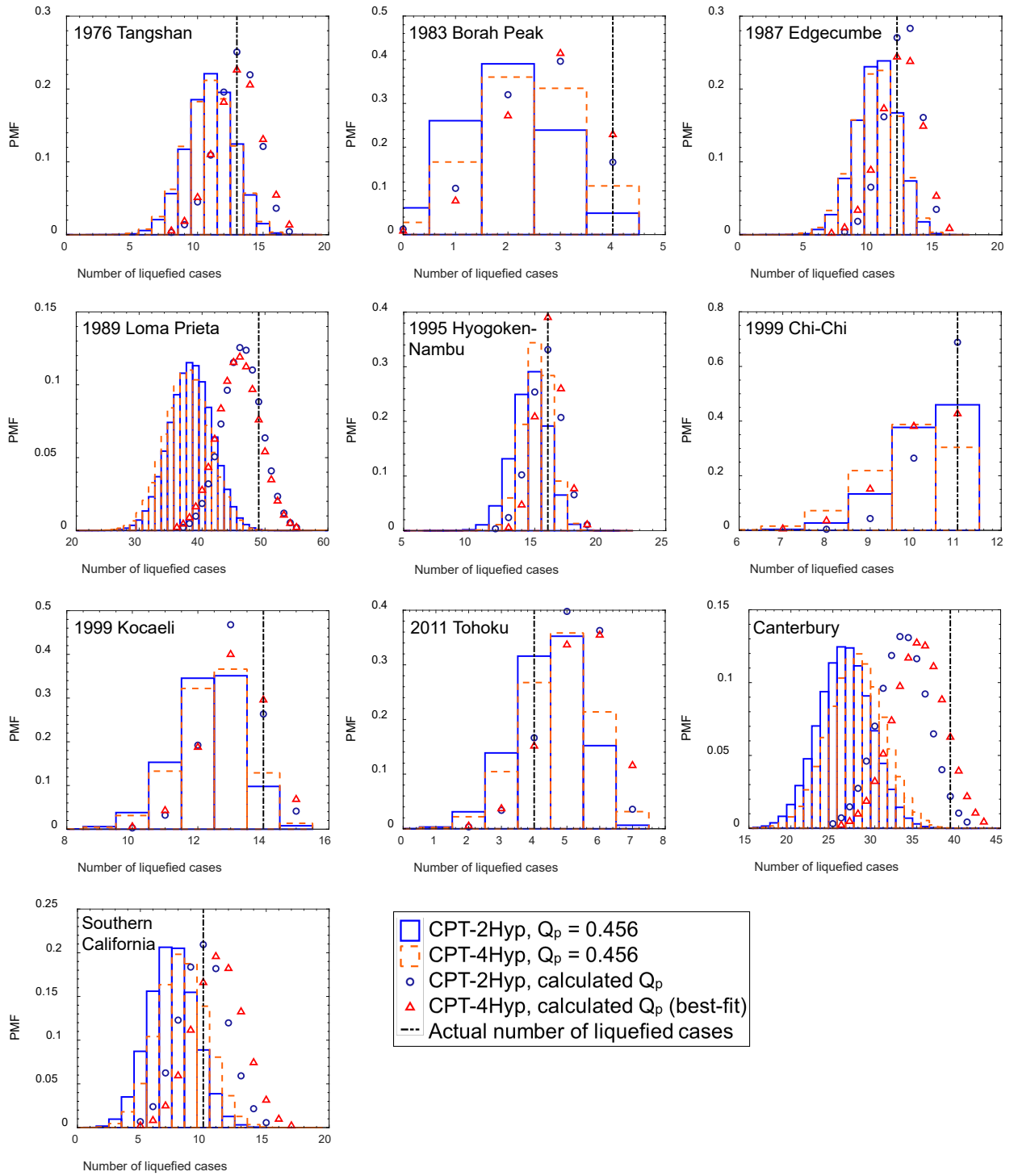
812 **Fig. 5.** Posterior sample distributions for (a)  $[\mu_\mu, \sigma_\mu, \sigma_c]$  in CPT-2Hyp and SPT-2Hyp models ( $Q_p=0.456$ ), (b)

813  $[\mu_\mu, \sigma_\mu, \sigma_c]$  in CPT-2Hyp and SPT-2Hyp models (calculated  $Q_p$ ), and (c)  $[\mu_\mu, \sigma_\mu, \mu_\sigma]$  in CPT-4Hyp and SPT-

814 4Hyp models ( $Q_p=0.456$ ). The prior relative frequency is derived considering the bin width of histogram.

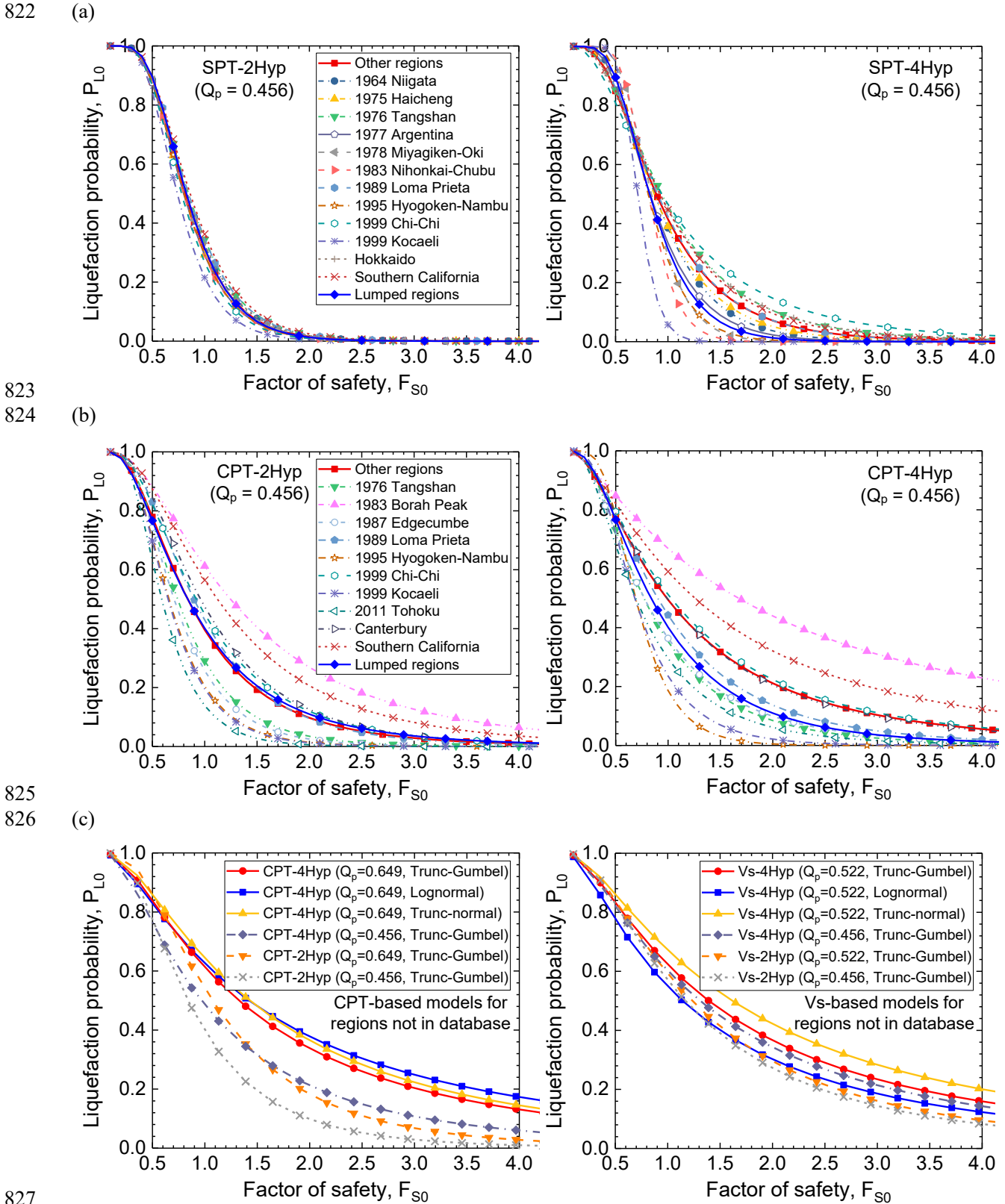


817 (b)



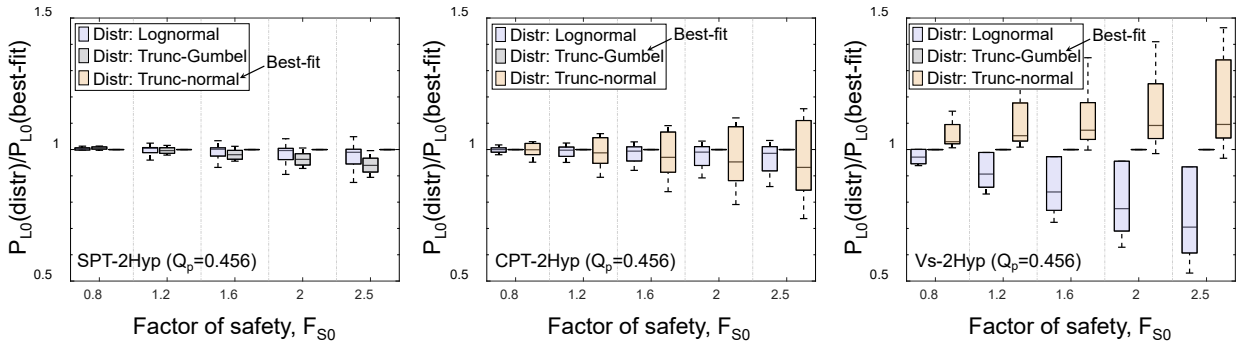
818

819 **Fig. 6.** Probability mass functions of predicted number of liquefied cases for region-specific (a) SPT-BI12  
 820 models and (b) CPT-RW98 models with four HBM configurations. The actual number of liquefied cases is  
 821 also shown.



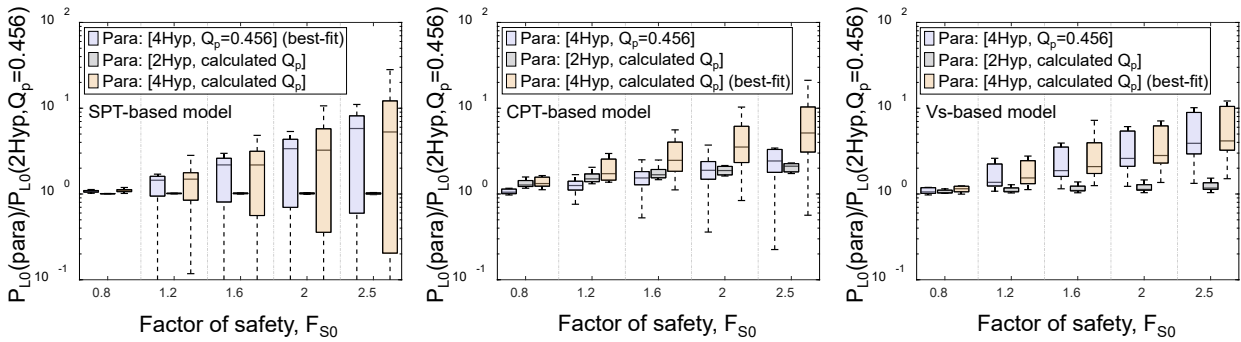
827  
 828 **Fig. 7.** Relationships between  $P_{L0}$  and  $F_{S0}$  for (a) region-specific SPT-BI12 models with  $Q_p=0.456$  and best-  
 829 fit distribution, (b) region-specific CPT-RW98 models with  $Q_p=0.456$  and best-fit distribution, and (c) CPT-  
 830 RW98 and Vs-R22 models with different HBM configurations for regions not in the calibration databases.

831 (a)



832

833 (b)

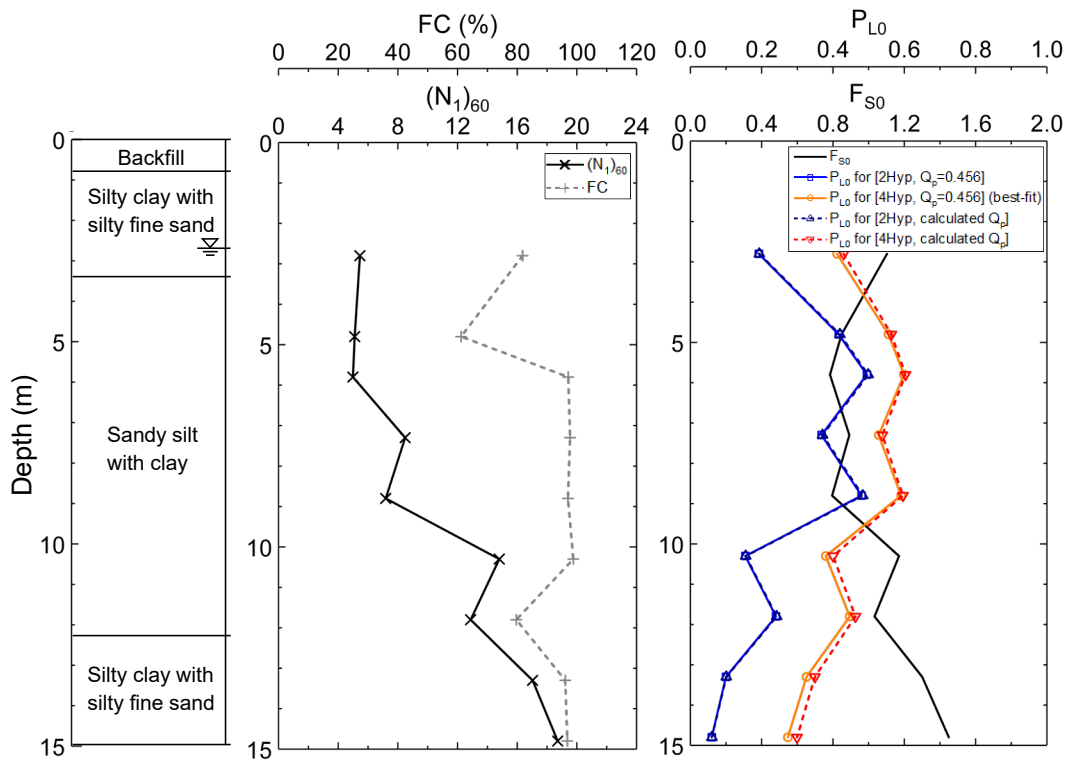


834

835 **Fig. 8.** (a) Ratios of  $P_{L0}$  for three candidate distributions to  $P_{L0}$  for best-fit distribution using [2Hyp,  
 836  $Q_p=0.456$ ]; (b) ratios of  $P_{L0}$  for three parameter configurations to  $P_{L0}$  for [2Hyp,  $Q_p=0.456$ ] using the best-  
 837 fit distribution. Each boxplot displays the distribution of ratio values for all region-specific models, in which  
 838 the central line marks the 50<sup>th</sup> percentile, the edges of box mark the 25<sup>th</sup> and 75<sup>th</sup> percentiles, and the ends of  
 839 whiskers represent the 0.35<sup>th</sup> and 99.65<sup>th</sup> percentiles.

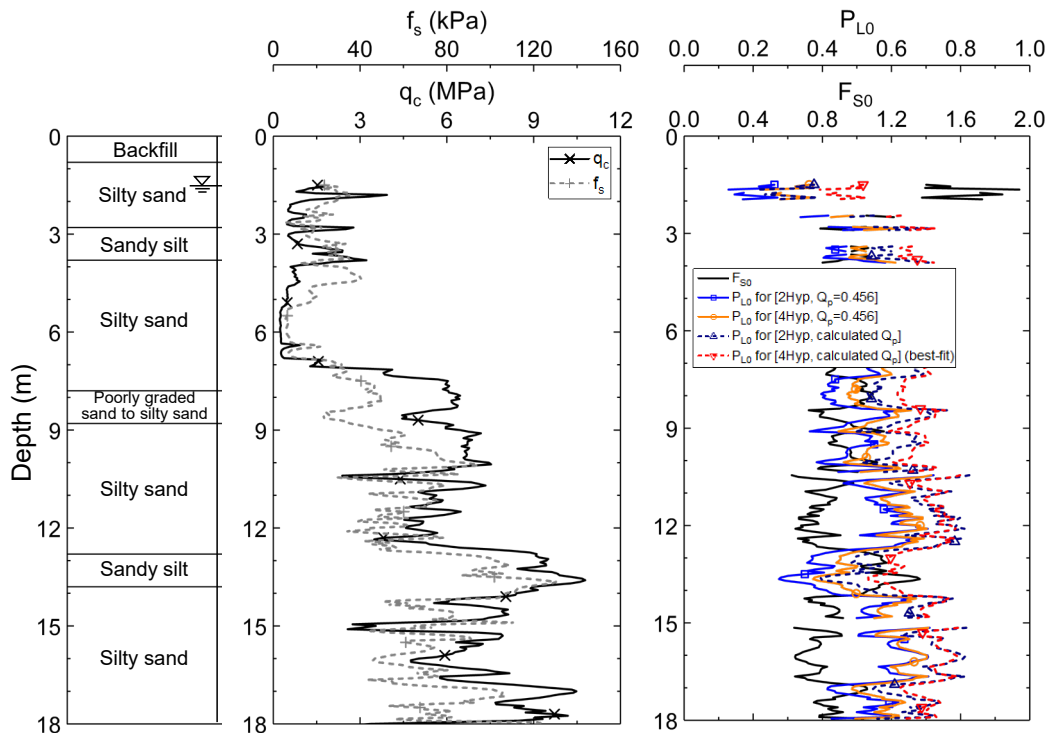
840

841 (a)



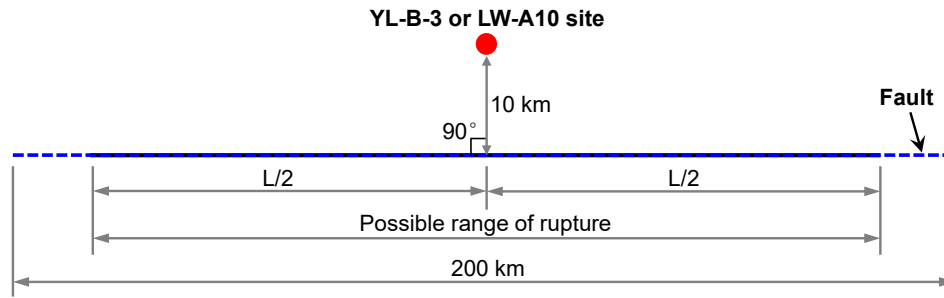
842

843 (b)

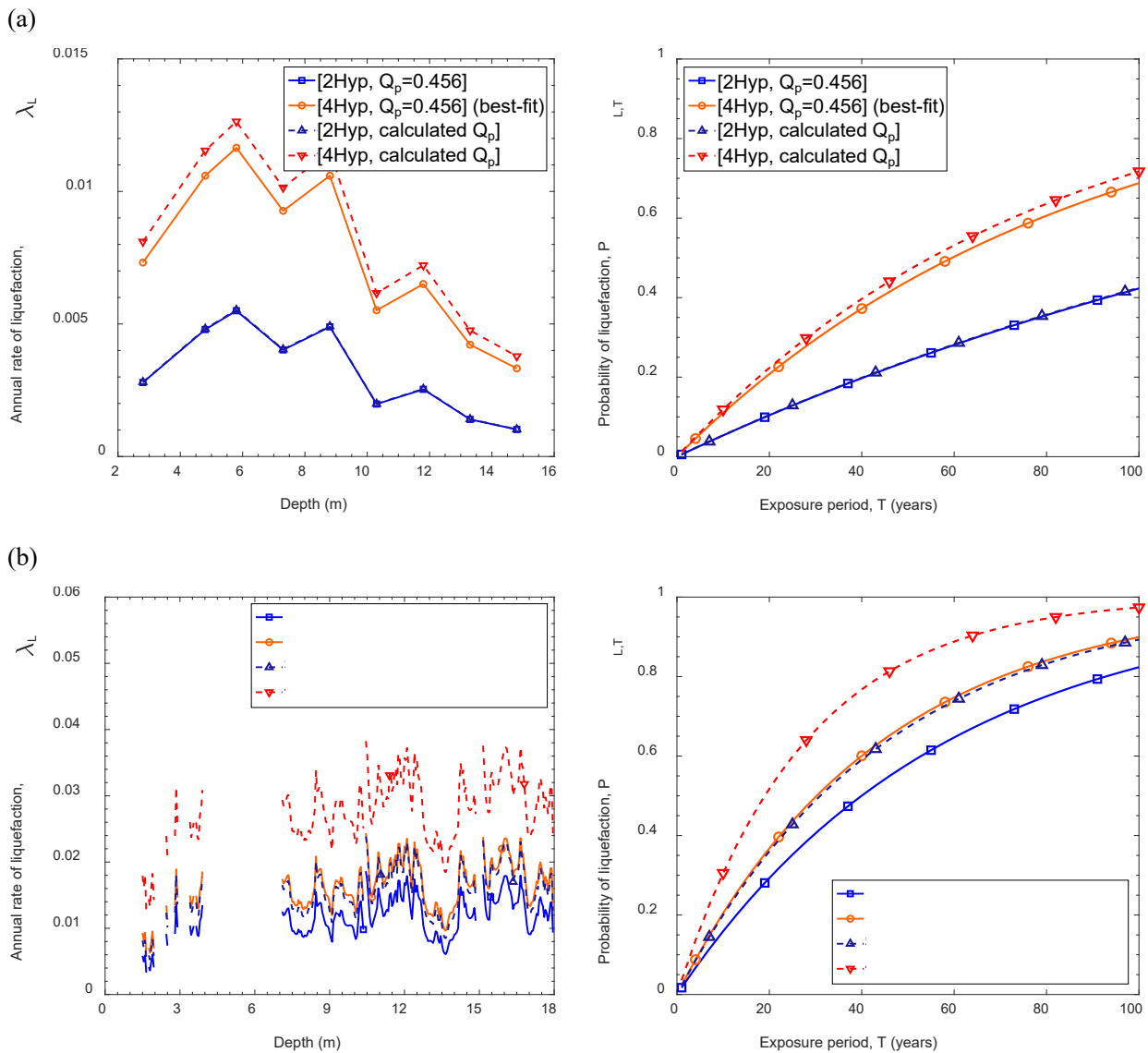


844

845 **Fig. 9.** Scenario-based liquefaction potential assessments for (a) YL-B-3 profile and (b) LW-A10 profile  
 846 using updated SPT-BI12 models and CPT-RW98 models, respectively. The 1999 Chi-Chi earthquake region  
 847 is considered.

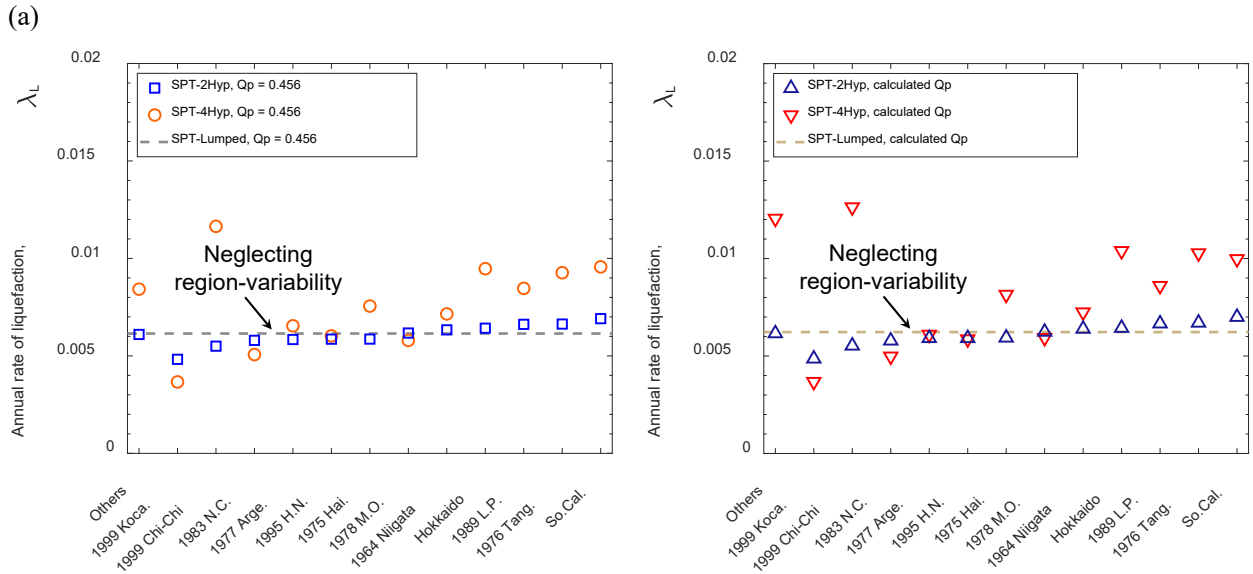


848  
849 **Fig. 10.** Source-site relationship in the illustrative example.



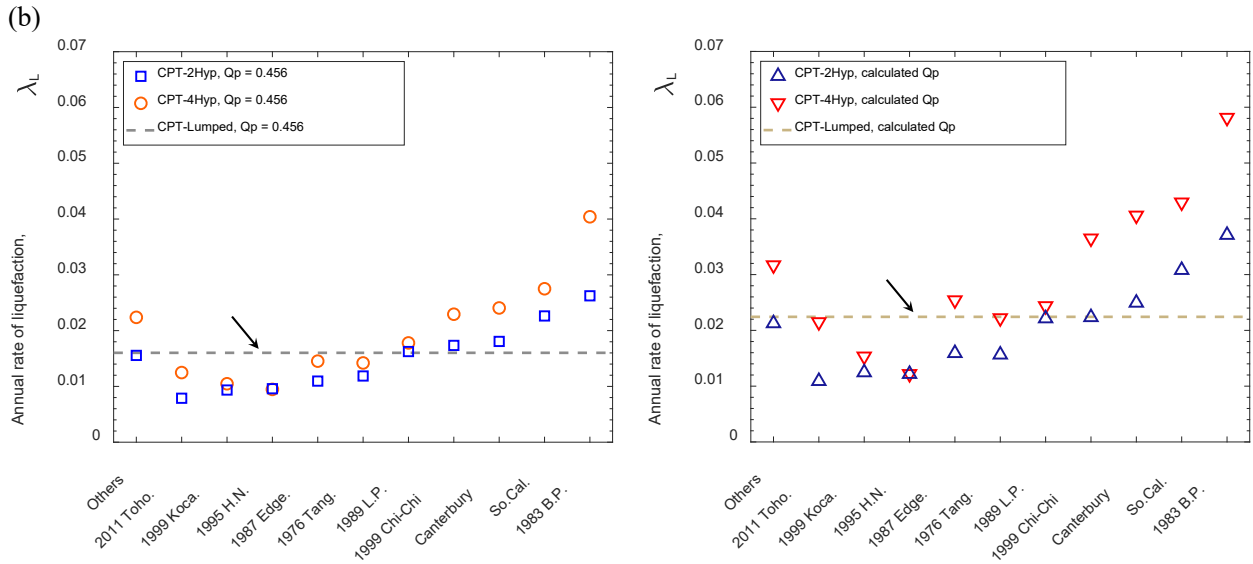
854  
855 **Fig. 11.** Probabilistic liquefaction-triggering hazard results: (a)  $\lambda_L$  versus depth and  $P_{L,T}$  at 5.8 m depth versus  
856 exposure period for YL-B-3 profile; (b)  $\lambda_L$  versus depth and  $P_{L,T}$  at 10.5 m depth versus exposure period for  
857 LW-A10 profile. The profiles are assumed to be located in regions influenced by the 1999 Chi-Chi  
858 earthquake.

860



861

862



863

864

**Fig. 12.** Comparison of annual liquefaction rates from region-specific (a) SPT-BI12 models for the 5.8 m

865

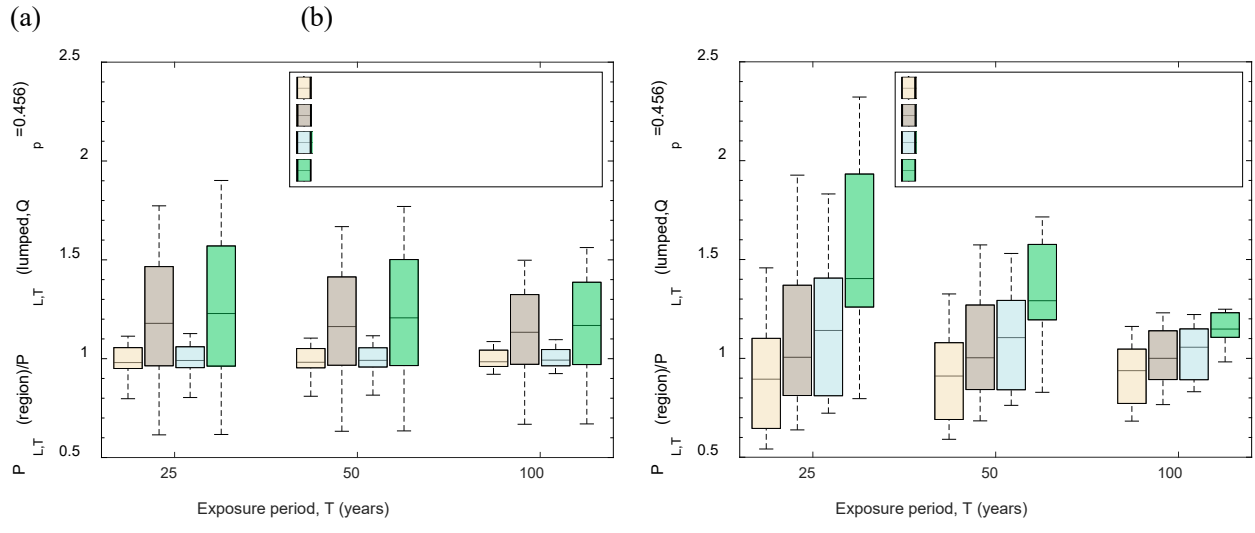
depth in YL-B-3 profile and (b) CPT-RW98 models for the 0.5 m depth in LW-A10 profile. The profiles are

866

assumed to be located in different regions.

867

868



869

870

871

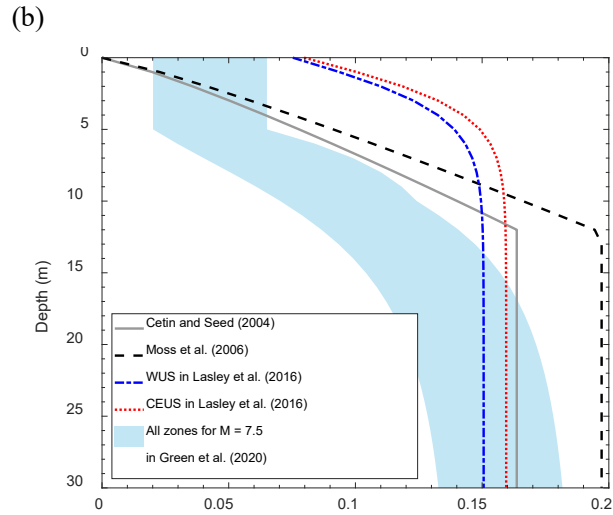
872

873

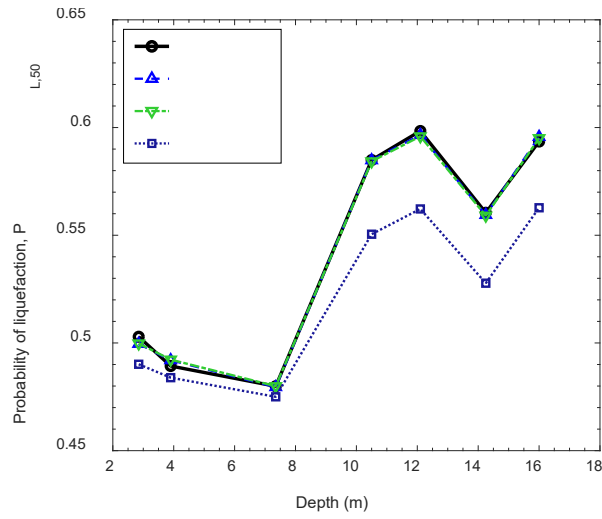
874

**Fig. 13.** Ratios of exposure period-consistent liquefaction probabilities (a) between region-specific and lumped-region SPT-BI12 models and (b) between region-specific and lumped-region CPT-RW98 models. The 5.8 m depth in YL-B-3 profile and 10.5 m depth in LW-A10 profile are considered for the SPT- and CPT-based models, respectively.

875 (a)

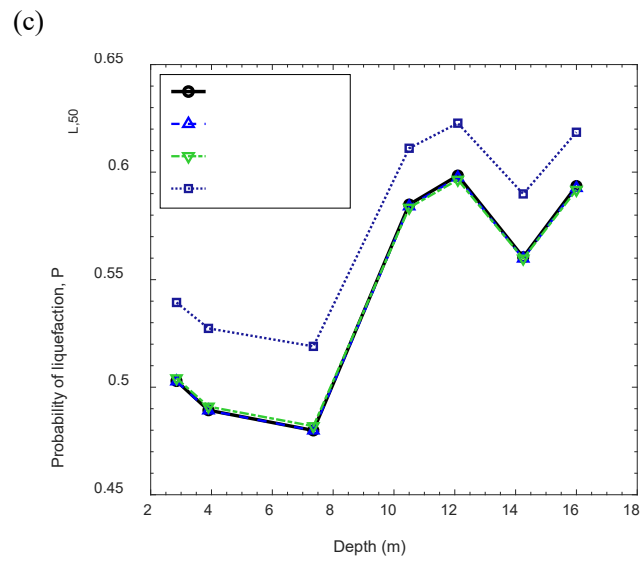


876



877

878



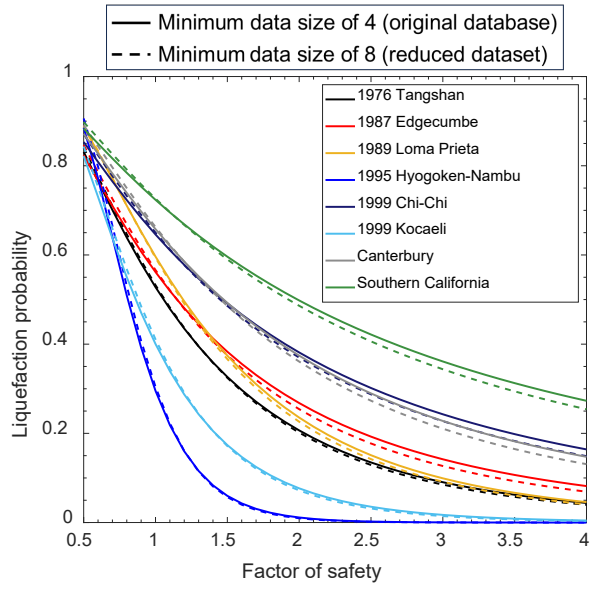
879

880 **Fig. 14.** (a) Variation of reported standard deviation of  $r_d$  with depth; (b) effect of  $\sigma_{rd}$  on  $P_{L,50}$  using [CPT-

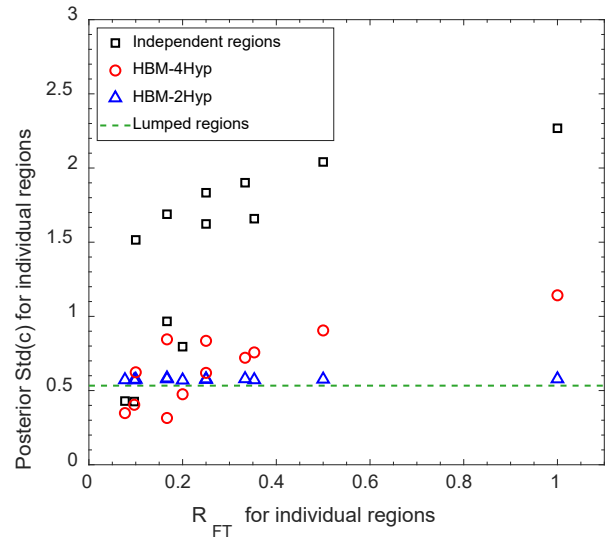
881 2Hyp,  $Q_p=0.456$ ]; (c) effect of  $\sigma_{inMSF}$  on  $P_{L,50}$  using [CPT-2Hyp,  $Q_p=0.456$ ].

882

(a)



(b)



883

884 **Fig. 15.** (a) Comparison between  $P_{L0}-F_{S0}$  relationships derived from the reduced dataset and the original

885 database; (b) comparison among Std( $c$ ) values derived from different Bayesian models.

886