

**Long-term temporal variations and source changes of halocarbons in the
Greater Pearl River Delta region, China**

Lewei Zeng^{1#}, Juan Dang^{1#}, Hai Guo^{1*}, Xiaopu Lyu¹, Isobel J. Simpson², Simone
Meinardi², Yu Wang¹, Luyao Zhang¹, Donald R. Blake²

¹Air Quality Studies, Department of Civil and Environmental Engineering, The Hong
Kong Polytechnic University, Hong Kong, China

²Department of Chemistry, University of California at Irvine, USA

*Corresponding author. ceguohai@polyu.edu.hk

These authors contributed equally to this work.

Highlights:

1. Mixing ratios of CFCs (CFC-11, CFC-12 and CFC-113) decreased over the past 18 years except for CFC-114, which increased in the Greater PRD region.
2. Refrigeration applications were the main sources of halocarbons in the Greater PRD region.
3. Refrigeration industry dominated by CFCs decreased during the past 18 years, while source of CFCs replacement increased.
4. Based on the measured ratios of halocarbons to CO, the estimated total halocarbon emissions in 2016 were 46.5 ± 16.7 Gg for the whole Greater PRD region.

27 **Abstract**

28 Halocarbons are widely used in the Greater Pearl River Delta (PRD) region of China.
29 To study the long-term trends, source changes and emissions of major halocarbons, a
30 total of 1505 canister air samples were collected in the Greater PRD region during 2001-
31 2018. Mixing ratios of CFCs decreased significantly over the past 18 years except for
32 CFC-114, which significantly increased in the Greater PRD, consistent with the recent
33 observations of significant CFC-114 emissions in East Asia. Declines in CFCs in the
34 Greater PRD region were faster than those at the background Mauna Loa (MLO) site,
35 indicating effective control regulations. Source apportionment simulations showed that
36 refrigeration applications, including refrigeration industry and CFCs replacement, were
37 the main sources of halocarbons. During the study period, refrigeration industry
38 experienced a progressive decline in both mixing ratio and percentage, while the
39 contribution of CFCs replacement remained increasing. Contribution of solvent use in
40 electronic industry traced by C₂Cl₄ dramatically decreased during the study period, and
41 stayed at a low level in recent years. Based on the measured ratios of halocarbons to
42 CO, the estimated total halocarbon emissions in 2016 were 46.5 ± 16.7 Gg for the
43 Greater PRD region. This study provides useful information for examining the evolving
44 emission status of halocarbons in the Greater PRD region in response to control
45 strategies and changing usage.

46 **Keywords:** Halocarbons; Long-term trends; Greater Pearl River Delta; Source
47 apportionment; Emission estimation

48

49

50

51

52

53

54

1. Introduction

Halocarbons are an important category of volatile organic compounds (VOCs), and are extensively used as foam blowing agents, refrigerants, aerosol propellants, dry cleaning solvents, degreasing agents, feedstock for chemical production, and fire extinguishing agents [Godish, 2003]. Among the halocarbons, chlorofluorocarbons (CFCs) and hydrochlorofluorocarbons (HCFCs) are ozone-depleting substances (ODSs) due to their long atmospheric lifetimes and catalytic impact on ozone decomposition in the stratosphere. Hydrofluorocarbons (HFCs) were then introduced as replacements of HCFCs and CFCs because of their near-zero ozone depletion potential (ODP) values [Fortems-Cheiney et al., 2013; Hundy et al., 2016]. However, because HFCs are powerful greenhouse gases, their emissions may make an increasing contribution to global radiative forcing [Ou-Yang et al., 2015; Simmonds et al., 2017; Velders et al., 2015]. Given their high ODP, production and emissions of many halocarbons have been regulated under the Montreal Protocol and its amendments [Grubb et al., 1999; Maione et al., 2011].

Mainland China belongs to the “Article 5” in the Montreal Protocol. According to the reduction plans, the production and consumption (usage) of CFCs was banned in mid-2007 with the exception of the exempted uses [WMO/UNEP, 2018]. The consumption of HCFCs was frozen in 2013 at the baseline of the average 2009-2010 level, with a goal of 10% reduction by 2015 and 35% by 2020. For Hong Kong (HK), as a “non-Article 5” party, the consumption of CFCs was phased out in 1996, and with a goal of 90% reduction of HCFC consumption by 2015 and 100% by 2030 [Fang et al., 2012; HKEPD, 2018a; Hurst et al., 2006]. As global production and consumption of CFCs was phased out in response to the Montreal Protocol, the leakage of CFCs in existing products and landfills, referred as “banks”, could become the important sources of CFCs emission in the future [Zhang et al., 2010a]. However, it is important to note that some ODSs (e.g. CFC-11) are not decreasing in the atmosphere as quickly as expected based on the Montreal Protocol, implying ongoing emissions that appear to be unreported new production rather than release from banks [Rigby et al., 2019; SPARC,

2017; Lunt *et al.*, 2018; Montzka *et al.*, 2018]. As temporary substitutes of CFCs and HCFCs, the consumption and production of HFCs are increasing steeply in developing countries since CFCs were banned and the consumption of HCFCs was frozen [Su *et al.*, 2015].

Besides CFCs and their replacements, other halocarbons also play important roles in ozone depletion processes in the stratosphere. CH_3CCl_3 is of particular interest since its atmospheric mass budget has been used to estimate ambient concentrations of the hydroxyl radical (OH) [Prinn *et al.*, 1987; Gentner *et al.*, 2010]. CCl_4 was used as a feedstock in the production of CFC-11 and CFC-12, and it is still used as a feedstock in certain countries [Hurst *et al.*, 2004; Harrison *et al.*, 2017; WMO/UNEP, 2018]. According to the terms of the Montreal Protocol and its Amendments, production and consumption of CH_3CCl_3 and CCl_4 were both banned in developed countries in 1996 and in developing countries in 2015 and 2010, respectively [Hurst *et al.*, 2006; Harrison *et al.*, 2017; WMO/UNEP, 2018]. Bromine also plays an important role in stratospheric ozone destruction. Halons, with solely anthropogenic sources, were primarily used as fire suppressants [Fraser *et al.*, 1999]. H-1211, H-1301 and H-2402 are major halons in the troposphere, which have been phased out in HK by 1994, and in mainland China by 2007. CH_3Br , another type of bromine, is released from both natural sources (oceans, biomass burning, rice paddies and other terrestrial ecosystems) and anthropogenic sources (soil fumigation and leaded gasoline) [Yokouchi *et al.*, 2002; Maione *et al.*, 2011], and its anthropogenic uses have been prohibited under the Montreal Protocol, except some consumption in quarantine and pre-shipment applications [WMO/UNEP, 2018; Yvon-Lewis *et al.*, 2009]. CH_3Cl , the most abundant halocarbon, is mainly emitted from biomass burning, oceans, tropical vegetation, coal combustion and waste incineration [Blake *et al.*, 1996; Li *et al.*, 2017; Maione *et al.*, 2011; Wang *et al.*, 2005; WMO/UNEP, 2018; Youkouchi *et al.*, 2000]. CH_3Cl is also commonly used as a solvent, processing agent and feedstock for chemical production in China [Li *et al.*, 2017], and Guo *et al.* [2009] speculated that about 40% of the CH_3Cl in the Greater Pearl River Delta (PRD) of China was emitted from solvent usage in the

113 production processes of refrigerants. C₂Cl₄ is primarily used as a dry cleaning solvent
114 and degreasing agent, and its predominant anthropogenic source makes it an excellent
115 marker for industrial pollution [*Simpson et al., 2004*].

116 Furthermore, some very short lived substances (VSLS), such as CH₂Br₂, CHBr₃ and
117 CH₃I, are largely produced by oceanic emissions [*Lennartz et al., 2015*]. CH₂Cl₂,
118 another VSLS, with an emission distinctly increased in the troposphere in the past
119 decade [*Oram et al., 2017*]. It is widely used as solvents in paints, paint strippers and
120 degreasing processes, and as an aerosol propellant and a blowing and cleaning agent in
121 foam production [*Borkar et al., 2010*]. Besides, CH₂Cl₂ was recently reported to be
122 used as a feedstock in the production of HFC-32 [*Leedham Elvidge et al., 2015*], which
123 is the main ingredient of the refrigerant R-410A increasingly used in new indoor air-
124 conditioning to replace HCFC-22 [*Fang et al., 2016*]. Although VSLS are not regulated
125 by the Montreal Protocol due to their short atmospheric lifetimes and low ODPs, a
126 recent study indicated their growing threat to the O₃ layer [*Oram et al., 2017*].

127 The inner PRD region is located in Guangdong province in southern China, and consists
128 of nine cities (Guangzhou, Shenzhen, Zhuhai, Dongguan, Zhongshan, Foshan,
129 Jiangmen, Huizhou, and Zhaoqing), while the Greater PRD region extends the inner
130 PRD region to Hong Kong and Macau (*Figure 1*). As a heavily populated and leading
131 manufacturing region, the atmospheric abundances and trends of halocarbons in this
132 region have attracted much attention and been the focus of several measurement studies
133 in the past two decades. *Chan and Chu* [2007] analyzed halocarbons in the inner PRD
134 region in 2000, while halocarbons were analyzed in 45 Chinese cities including the
135 PRD region in winter 2001 [*Barletta et al., 2006*]. *Guo et al.* [2009] analyzed
136 halocarbons from 2001 to 2002 and investigated their source origins in both inner PRD
137 and HK. *Shao et al.* [2011] reported the mixing ratios of halocarbons in 2004 in the
138 PRD region and estimated their emissions, and *Zhang et al.* [2014] investigated the
139 spatial variations of CFCs and HCFC-22 in 2008 and 2009 in the PRD region. However,
140 there are limited studies on the long-term variations of halocarbons as well as their
141 source changes since the 2010 phase-out in China [*Su et al., 2015*]. *Zhang et al.* [2010b]

142 studied long-term variations and emission patterns of halocarbons from 1998 to 2008
143 in the Greater PRD region. Therefore, more updated studies are needed to examine the
144 temporal variations and emission status of halocarbons in the Greater PRD region.

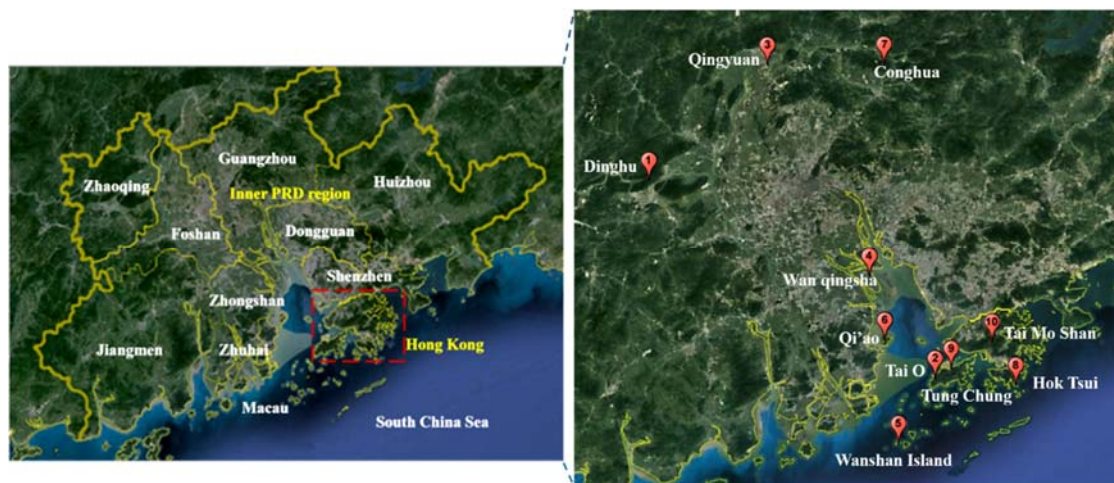
145 In this study, comprehensive field measurements were carried out from 2001 to 2018
146 at different sites in the Greater PRD region. The data provide an opportunity to
147 investigate the long-term variations, source changes and emissions for halocarbons, and
148 infer the implementation status of the Montreal Protocol in the Greater PRD region.

149 **2. Methodology**

150 **2.1 Description of the Sampling Sites**

151 In this study, the inner PRD sampling sites were Dinghu, Qingyuan, Wanqingsha,
152 Wanshan Island, Qi'ao Island, and Conghua. In addition, field campaigns were carried
153 out in HK at Hok Tsui, Tai O, Tung Chung and Tai Mo Shan ([Figure 1](#)). Some of the
154 sampling sites were ambient air quality monitoring stations established by the local
155 governments while others were set up by our group. Detailed information of the
156 sampling sites and corresponding sampling time are presented in [Table 1](#). Briefly, the
157 Dinghu site, located in Zhaoqing, was a rural site surrounded by tropical forest [[Chan
158 et al., 2006a](#)]. Qingyuan site was a rural area in northwestern Guangzhou [[Zhang et al.,
159 2008](#)]. Wanqingsha, surrounded by croplands and some clothing workshops, was a
160 suburban site in southern Guangzhou [[Zhang et al., 2010b](#)]. In addition, there was
161 another rural site (Conghua) in the northeastern Guangzhou surrounded by lakes and
162 forests. The remote Wanshan Island site was in the South China Sea and the distance
163 from this island to urban centers of Hong Kong, Macau and Zhuhai was about 64 km,
164 35 km and 40 km, respectively [[Wang et al., 2018](#)]. The rural Qi'ao Island site was
165 located in the northeastern Zhuhai, on the west bank of the Pearl River Estuary. Hok
166 Tsui was a background site situated at the south tip of HK with sparse anthropogenic
167 sources [[Lee et al., 2002](#)]. Tai O, a rural/coastal site in southwestern HK, was about 32
168 km from the urban center [[Wang et al., 2005](#)]. Tung Chung was situated on northern
169 Lantau Island and was a typical suburban site receiving both local and regional air
170 masses. Tai Mo Shan site was located in the summit of the highest mountain in HK,

171 surrounded by Country Parks to the east, south, and west, and defined as a suburban
172 site [Ling *et al.*, 2014]. On the whole, the air samples collected at these sites were
173 representative of typical local halocarbon levels in the atmosphere since all the
174 sampling sites were suburban and/or rural locations with similar geographical
175 proximity, distant from emission sources.



176

177 **Figure 1.** Locations of the Sampling Sites in the Greater Pearl River Delta Region

Table 1. Description of the Sampling Sites for the Field Campaigns

Site	City	Longitude (°)	Latitude (°)	Site Type	Sampling Time	Number of samples
Dinghu	Zhaoqing	112.56	23.15	Rural	3/3/2001-3/19/2001	39
Qingyuan	Qingyuan	113.05	23.69	Rural	2/29/2004-4/13/2004	33
Wanqingsha	Guangzhou	113.55	22.71	Rural	4/20/2004-6/29/2004; 10/26/2007-12/1/2007	88 102
Wanshan Island	Zhuhai	113.70	21.93	Marine Background	9/11/2013- 11/21/2013	170
Qi'ao	Zhuhai	113.65	22.42	Rural	10/14/2016-11/18/2016	118
Conghua	Guangzhou	113.62	23.65	Rural	1/18/2018- 1/23/2018; 6/29/2018-7/4-2018	36
Hok Tsui	Hong Kong	114.25	22.21	Rural	3/3/2001-4/22/2002	79
Tai O	Hong Kong	113.85	22.26	Rural	5/3/2002-12/31/2002	109
Tai O (Regional) ^a	Hong Kong	113.85	22.26	Rural	10/10/2001-3/6/2002	28
Tung Chung	Hong Kong	113.96	22.29	Suburban	1/6/2003-8/28/2003;	40
					9/12/2006-2/22/2008;	155
					9/11/2013-11/21/2013;	148
					11/1/2016-11/18/2016	187
Tai Mo Shan	Hong Kong	114.12	22.40	Rural	9/28/2010-11/21/2010	201

a. Regional air masses from inner PRD region obtained by backward Lagrangian particle release simulations [*Guo et al., 2009*].

180 2.2 Sampling and Chemical Analysis

181 Whole air samples were collected in 2-L stainless steel canisters with electro-polished
182 inner surface, which were conditioned beforehand. Prior to sampling, canisters were
183 repeatedly cleaned at least four times by filling and evacuating zero air. Then, all
184 cleaned canisters were evacuated to 180 mtorr for sampling. A total of 1505 ambient
185 air samples were collected on selected days from 2001-2018, and sometimes multiple
186 samples were collected in one day. The sampling flow rate was ~35 mL/min and
187 sampling duration was one hour. After collection, samples were shipped to the
188 laboratory for chemical analysis within two weeks. 250 mL of each sample were
189 transferred to a pre-concentrator, concentrated with liquid nitrogen, and injected into a
190 gas chromatography-mass selective detector / electron capture detector (GC-MSD /
191 ECD) system. Most target substances were analyzed using ECD, while HCFCs, HFCs,
192 CFC-113, H-1211, CH₃Cl, CH₂Cl₂ and CH₃Br were measured using MSD in selected
193 ion monitoring (SIM) mode (target ions are listed in [Table S1](#)). Calibration standards
194 were produced and regularly quantified by the Rowland/Blake laboratory in the
195 University of California at Irvine (UCI). They followed the same protocols to prepare
196 their own standards for decades in order to keep consistency across standards
197 maintained and during the course of this study. Calibration curves were made with at
198 least five points ranging from 0 to 4 ppbv for quantification of the target halocarbons.
199 Linear correlation coefficients (R^2) between the responses and standard concentrations
200 were all above 0.995. Accuracy, replicate precision and method detection limit for each
201 species were calculated according to TO-14A/15 methods issued by USEPA (see
202 detailed description in [Text S1](#)), which were 1.3 to 8.0 %, 0.2 to 7.9 % and 0.2 to
203 5.7 pptv, respectively ([Table S1](#)).

204 Samples collected in earlier years (prior to 2013) were analyzed in the UCI laboratory,
205 while more recent samples were mainly measured in our laboratory at the Hong Kong
206 Polytechnic University (HKPU) using the same analytical method as that in the UCI
207 laboratory (the analytical systems in our laboratory were established following the
208 prototype of the UCI laboratory). Inter-comparison experiments were undertaken

209 between these two analytical systems, and the chemical analysis results were
210 comparable (as described in [Text S2](#)).

211 In addition, the trace gas carbon monoxide (CO) was simultaneously monitored during
212 most sampling campaigns using a gas filter correlation trace level CO analyzer (API
213 model 300EU) coupled with a heated, platinum CO scrubber. The quality assurance
214 and quality control of CO measurement are described in detail in [Text S3](#).

215 **2.3 PMF Receptor Model**

216 The source changes of halocarbons between 2001 and 2018 in the Greater PRD region
217 were identified using the Positive Matrix Factorization (PMF) 5.0 model. In this model,
218 based on the composition or fingerprints of the sources, a mathematical method is
219 applied to determine the contribution of sources to samples. An air sampling data set
220 can be regarded as a data matrix x of i by j dimensions. The chemical mass balance
221 between measured species concentrations and source profiles can be described as
222 follows (Equation (1)):

$$223 \quad x_{ij} = \sum_{k=1}^p g_{ik} f_{kj} + e_{ij} \quad (1)$$

224 where x_{ij} is the mixing ratio of j th species in the i th sample, g_{ik} is the proportion of the
225 k th source to i th sample, f_{kj} is the fraction of the j th species in the k th source, e_{ij} is the
226 residual value for the j th species in the i th sample, and p is the total number of
227 independent sources [[Paatero, 1997](#)].

228 As a multivariable factor analysis program, PMF resolves a matrix of sample data into
229 two matrices: factor contributions (G) and factor profiles (F). The various sources that
230 may contribute to the samples are determined according to the extracted factor profiles.
231 For all chemical species (m) in all the samples (n), factor contributions and profiles are
232 obtained from the model minimizing the objective function Q (Equation (2)):

$$233 \quad Q = \sum_{i=1}^m \sum_{j=1}^n \left[\frac{x_{ij} - \sum_{k=1}^p g_{ik} f_{kj}}{u_{ij}} \right]^2 \quad (2)$$

234 where u_{ij} is an estimate of the “uncertainty” in the j th species in the i th sample, which
235 is determined by the equation $u = 5/6 \times$ method detection limit (MDL) when the

concentration is \leq the MDL value; or by $u = [(\text{error fraction} \times \text{mixing ratios})^2 + (\text{MDL})^2]^{1/2}$ when the concentration is $>$ MDL value [Paatero, 1997; Zhang et al., 2014].

To better understand the sources of halocarbons, the source contributions, and the source changes in the past 18 years in the Greater PRD region, source apportionment was conducted on the whole dataset using PMF v5.0. Ten main halocarbons (as shown in section 3.3) were used for the model simulation as they had relatively high ambient mixing ratios and are typical tracers of various emission sources. They were also available in most years except 2007 and 2010. In addition, as a combustion tracer, CO was an input for model simulation. In total, 934 valid samples were used as inputs for source apportionment, and the uncertainty of the input species was set as 10%.

2.4 Emission Estimation

Previous studies have indicated that the ratios between the enhancements of halocarbons and a tracer can demonstrate their emission strengths. The selection of the tracer should follow the following rules: (1) a long lifetime and low chemical reactivity occurred during the transport; (2) known emissions; and (3) correlation with halocarbons that is as good as possible. Carbon monoxide (CO) is often chosen as the tracer to determine the halocarbon emissions [Palmer et al., 2003; Guo et al., 2009]. Although the sources of halocarbons and CO are not exactly the same, their sources are often co-located, such as coexisted refrigeration and combustion in factories, vehicles and households. Moreover, similar to halocarbons, CO has a relatively long lifetime and is less reactive, and the emission inventory of CO is more easily accessible than other species. Hence, the ratio of halocarbon species to CO has been widely used to estimate the emissions of halocarbons [Palmer et al., 2003; Guo et al., 2009; Shao et al., 2011; Fang et al., 2012; Wang et al., 2014]. In this study, we used CO as the reference compound to estimate the halocarbon emissions in the Greater PRD region in 2016 as an example.

The emission of each halocarbon can be calculated by Equation (3) [Shao et al., 2011]:

$$E_X = E_{CO} \times R \times \frac{M_X}{M_{CO}} \quad (3)$$

264 where E_X and E_{CO} are the emissions of the halocarbon X and CO, respectively; R is the
 265 slope of the linear correlation between the enhancements of halocarbon and CO
 266 (Δ halocarbon and Δ CO), which are obtained by subtracting the background values,
 267 especially for long-term observations. In this study, background values were defined as
 268 the lowest 20th percentile of the complete datasets for each halocarbon. The lowest CO
 269 mixing ratio in South China Sea air (defined by backward trajectory simulation) was
 270 characterized as the CO background value [Shao *et al.*, 2011]. Halocarbon emissions in
 271 2016 were estimated, when samples were mainly collected in autumn (October and
 272 November). Under similar synoptic conditions within this short period, the background
 273 value was supposed to change insignificantly. Since the emission estimation was based
 274 on the slopes of correlations between the species, whether the background value was
 275 subtracted or not would not influence the results [Shao *et al.*, 2011; Fang *et al.*, 2012].
 276 M_X and M_{CO} are the molecular weights of the halocarbon X and CO, respectively.
 277 Using linear least squares fitting method, the uncertainties of halocarbon emissions can
 278 be estimated by Equation (4) [Shao *et al.*, 2011]:

$$\sigma_x = \sqrt{\sigma_{E_{CO}}^2 \times R^2 + E_{CO}^2 \times \sigma_R^2} \times \frac{M_x}{M_{CO}} \quad (4)$$

280 where σ_x is the uncertainty of the estimated halocarbon emission, $\sigma_{E_{CO}}$ and σ_R are
 281 the uncertainties of E_{CO} and R, respectively.

282 **3. Results and Discussion**

283 **3.1 General Features**

284 Table 2 summarizes the mean mixing ratios (with 95% confidence level) of 15
 285 halocarbons in the Greater PRD (subtropical climate in the Northern Hemisphere (NH)),
 286 which are compared with the levels at the Mauna Loa station (MLO), Hawaii (19.54°N,
 287 155.58°W) from 2001 to 2018 [ESRL, 2018]. The reason for choosing MLO site was
 288 that it provided a complete data set for the study period and is representative of NH
 289 subtropical background levels of halocarbons. Average mixing ratios of halocarbons in
 290 the Greater PRD were calculated in four sampling periods (I: 2001-2004, II: 2005-2008,

291 III: 2010-2013, and IV: 2015-2018), which mainly reflected the temporal variations of
292 halocarbons in this region. Moreover, backward trajectories were simulated on
293 individual sampling days and origins of air masses were clustered into four categories
294 (Figure S2). Categories 2 and 3 came from the northeast of the Greater PRD region,
295 passing by relatively clean coastal areas. Remarkably, under the influence of the
296 background of the South China continent, air masses in these two categories (82%)
297 stayed in Guangdong province for more than 24 hrs before reaching the study region,
298 representative of local characteristics. Hence, categories 2 and 3 were further analyzed
299 while categories 1 (super-regional) and 4 (oceanic) were excluded.

300 Among all measured halocarbons (Table 2), the average value of CH₃Cl was the highest,
301 followed by CH₂Cl₂, CFC-12 and HCFC-22, indicating large-scale emissions of these
302 species from various sources in the study region. In comparison, the averages of
303 halocarbons for the whole period in the Greater PRD region were all higher ($p < 0.05$)
304 than the tropical NH background values at the MLO site, consistent with our previous
305 findings [Guo *et al.*, 2009; Zhang *et al.*, 2010b]. Specifically, mixing ratios of HCFC-
306 22, H-1211, HFC-134a, CH₃Cl, CH₃CCl₃ and CH₃Br were approximately twice those
307 at the MLO site. Moreover, abundances of CH₂Cl₂ and C₂Cl₄ in the study region were
308 25 and 50 times those at MLO site. The results indicated significant contributions of
309 local emissions (including “bank” emissions) and perhaps unreported new production
310 [SPARC, 2017; Lunt *et al.*, 2018; Montzka *et al.*, 2018].

311 Inspection on the averages of CFC-113, H-1211 and CCl₄ in different periods found
312 their continuous decrease ($p < 0.005$), implying the effectiveness of the Montreal
313 Protocol regulations on these species. In addition, values of other CFCs, *i.e.* CFC-11
314 and CFC-12, declined ($p < 0.005$) from period III to period IV. Their levels in period III
315 were similar to the results of a study conducted in 2010 in this region [Wu *et al.*, 2013].
316 Similarly, C₂Cl₄, CH₂Cl₂ and CHCl₃ experienced obvious decline ($p < 0.005$) from
317 periods III to IV. In contrast, as substitutes of phased-out CFCs, HCFC-22 and HFC-
318 134a had the highest averages in period IV compared to previous periods, reflecting
319 their expanding productions in this region. Besides, similar levels of HCFC-22 (526

pptv, 69% as relative standard deviation) and HFC-134a (84 pptv, 48%) were reported in suburban Guangzhou, 2010 [Wu *et al.*, 2013]. As the most abundant species, CH₃Cl experienced an increase ($p<0.05$) from periods I to III, and remained unchanged in periods III and IV, reflecting a possible alleviation in its increase.

3.2 Long-term Temporal Variations of Halocarbons in the Greater PRD Region

To evaluate the implementation status of halocarbon control measures, it is crucial to know the long-term changes of halocarbons. Although the halocarbon data were not continuously measured at a fixed site in this study, it was still possible to infer the variation trends of halocarbons in these years if we integrated and analyzed a range of halocarbon levels at some rural or suburban sites with similar sources and geographical environments. Figure 2 presents the trends of monthly averages of fifteen key halocarbons in the Greater PRD region from 2001 to 2018, compared with the monthly average mixing ratios of eleven halocarbons at MLO with error bars (95% confidence interval) during the same period. It should be noted that error bars are provided for our measurement data whenever more than one sample was collected per month (Figure 2). The dashed lines are the linear fits to the halocarbon mixing ratios [Kock *et al.*, 2005; Qin *et al.*, 2007; Zhang *et al.*, 2010a; Zhang *et al.*, 2010c]. Please note: the variation rates of some species may not be exactly constant during the whole period due to different control regulations. However, the linear regression method can provide reliable average variation trends throughout the 18 years. Moderate to good R^2 of regressions were obtained for most halocarbons. The slope of each halocarbon represented the yearly variation rate of each halocarbon (Figure 2). For easy comparison, the yearly growth rates of target halocarbons in the Greater PRD region and at the MLO site are summarized in Table 3.

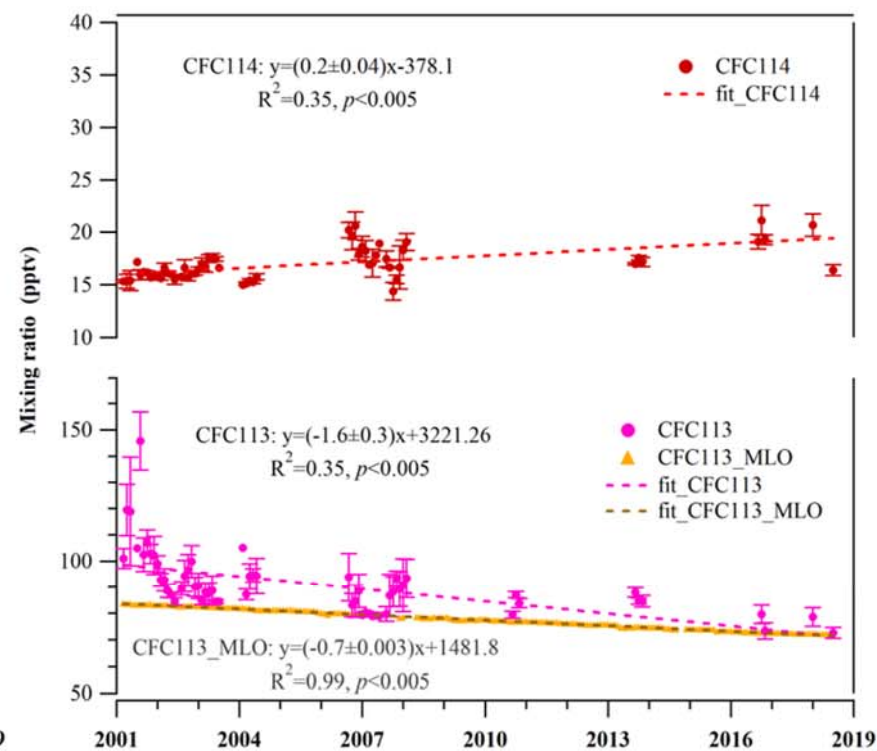
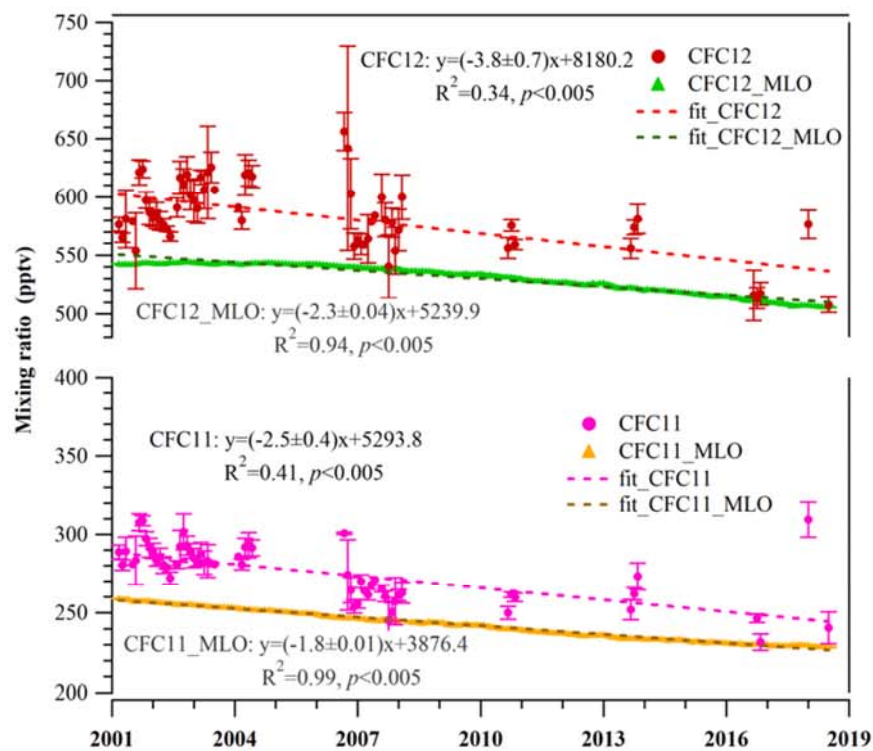
Overall, the monthly averages of all halocarbons in the Greater PRD region were higher than the corresponding NH background levels, revealing the contribution of anthropogenic activities to these halocarbons (Figure 2). However, the error bars of some species, *i.e.* CFC-11 and CFC-12, sometimes reached the levels lower than the

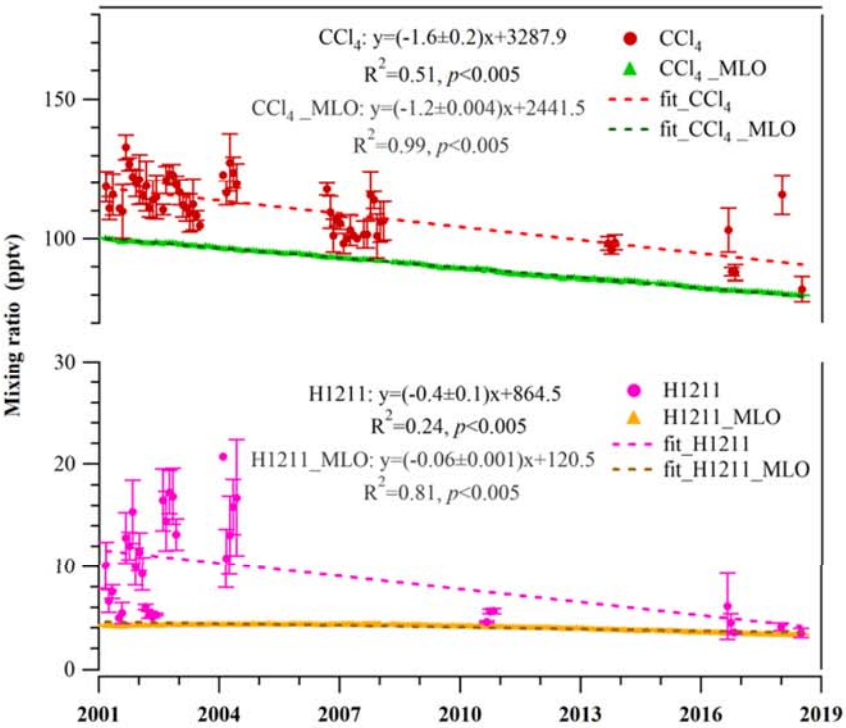
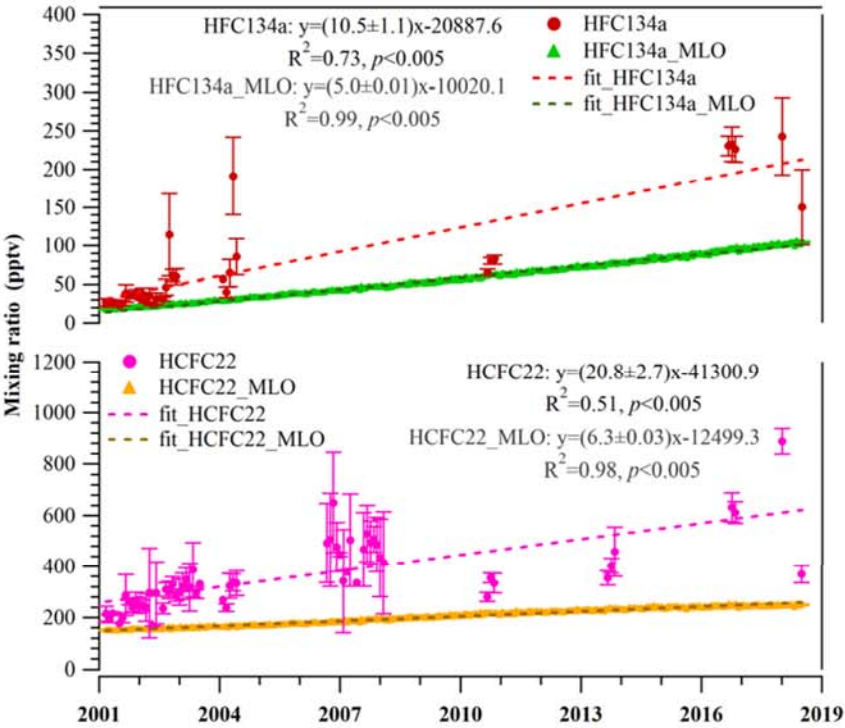
348 background values, perhaps caused by the discrepancy between our offline analytical
349 system and the online system at MLO site.

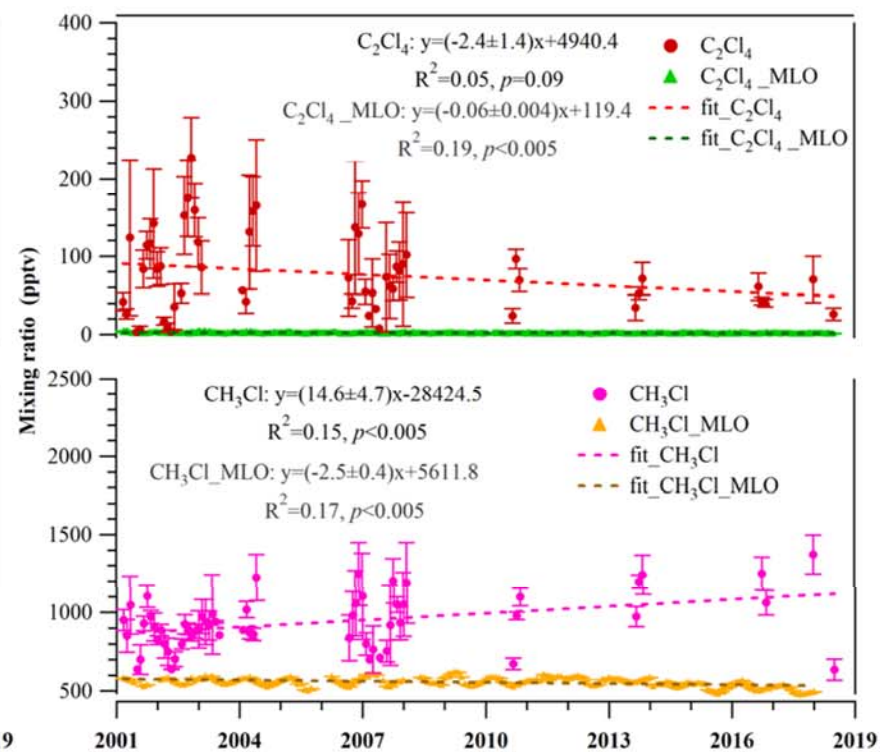
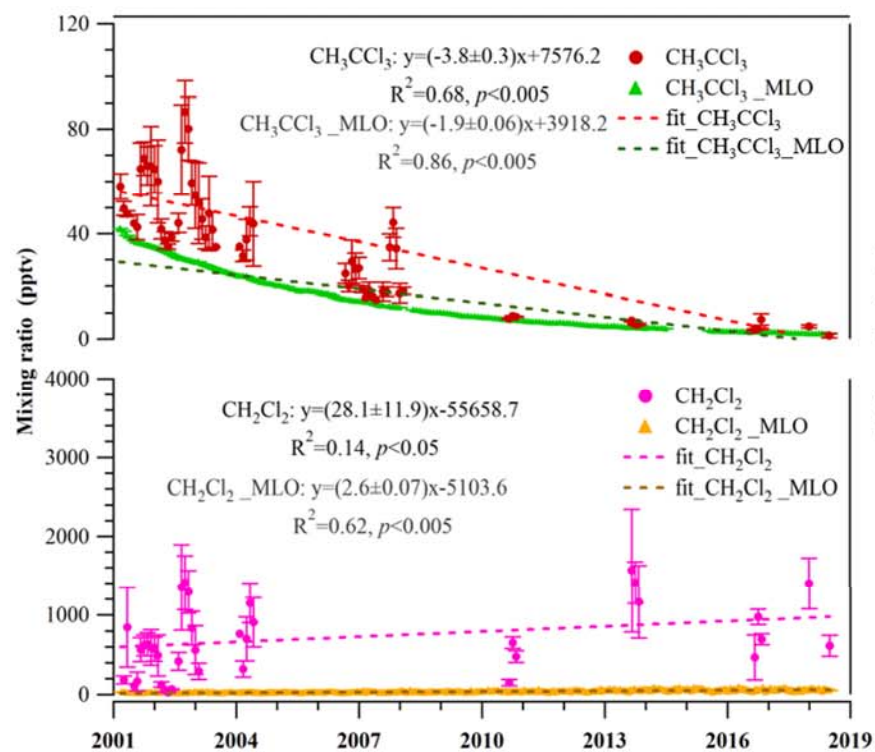
Table 2 Mixing Ratios of Halocarbons in the Greater PRD Region in Different Periods, Together With Subtropical Northern Hemisphere Background Levels from NOAA (Units: pptv)

Compounds (Chemical Formula)	MLO Background Mean ^a (95% CI) ^b		Average (\pm 95% CI) ^b in the Greater PRD			
	Whole period	Whole period	2001-2004	2005-2008	2010-2013	2015-2018
CFC-11 (CCl ₃ F)	242 \pm 1.3	270 \pm 1.7	292 \pm 1.7	249 \pm 4.8	261 \pm 2.0	245 \pm 5.3
CFC-12 (CCl ₂ F ₂)	530 \pm 1.6	571 \pm 42.9	603 \pm 3.2	569 \pm 9.7	571 \pm 3.3	508 \pm 6.6
CFC-113 (CCl ₂ FCClF ₂)	78 \pm 0.5	88 \pm 1.0	97 \pm 1.3	90 \pm 2.2	85 \pm 0.7	66 \pm 3.8
CFC-114 (CClF ₂ CClF ₂)		17 \pm 0.16	16 \pm 0.1	16 \pm 0.4	17 \pm 0.2	20 \pm 0.5
HCFC-22 (CHClF ₂)	204 \pm 2.1	398 \pm 11	281 \pm 8.8	502 \pm 36	379 \pm 17	625 \pm 33
H-1211 (CF ₂ ClBr)	4.1 \pm 0.04	9.2 \pm 0.47	13 \pm 0.7		5.6 \pm 0.2	3.9 \pm 0.3
HFC-134a (CH ₂ FCF ₃)	58 \pm 1.7	113 \pm 7.1	61 \pm 7.2		80 \pm 3.3	230 \pm 13
Methyl chloride (CH ₃ Cl)	555 \pm 4.2	1036 \pm 17	910 \pm 15	1059 \pm 45	1102 \pm 26	1171 \pm 63
Carbon tetrachloride (CCl ₄)	90 \pm 0.8	108 \pm 1.1	121 \pm 1.2	112 \pm 2.7	97 \pm 1.1	89 \pm 1.9
Methyl Chloroform(CH ₃ CCl ₃)	13 \pm 1.5	32 \pm 1.7	60 \pm 2.6	37 \pm 3.4	7.0 \pm 0.2	5.9 \pm 1.2
Tetrachloroethene(C ₂ Cl ₄)	1.7 \pm 0.04	85 \pm 5.2	121 \pm 11	79 \pm 12	65 \pm 5.9	43 \pm 3.9
Methylene chloride (CH ₂ Cl ₂)	39 \pm 1.1	960 \pm 53	826 \pm 73		1132 \pm 117	821 \pm 59
Methyl bromide (CH ₃ Br)	7.7 \pm 0.05	18 \pm 0.43	16 \pm 0.5	25 \pm 1.4	17 \pm 0.9	17 \pm 0.8
Chloroform (CHCl ₃)		59 \pm 2.4	45 \pm 2.4	178 \pm 16	78 \pm 4.3	43 \pm 3.3
Methyl iodide (CH ₃ I)		2.3 \pm 0.20	2.0 \pm 0.1		1.7 \pm 0.1	3.3 \pm 0.8

a. Subtropical NH background value from the NOAA/ESRL halocarbons program, <https://www.esrl.noaa.gov/gmd/dv/data/index.php?category=Halocompounds>. The monthly mean values at Mauna Loa, Hawaii from 2001 to 2018 are used. Measured in pptv; b. 95% CI represents 95% confidence interval.







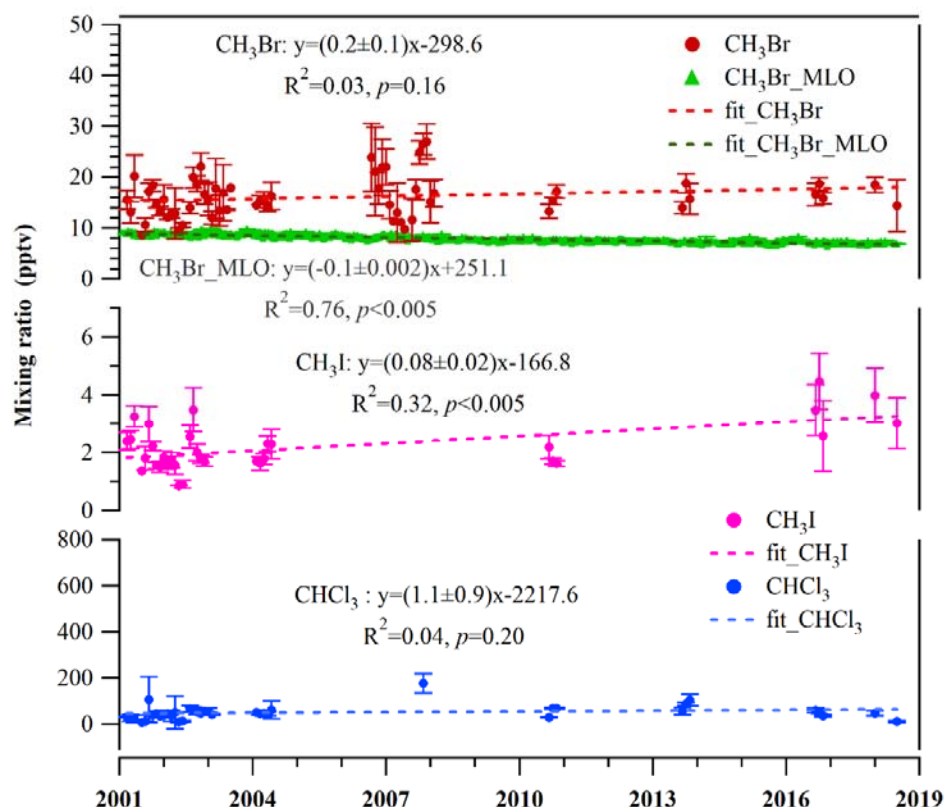


Figure 2 Monthly Variations of Halocarbons in the Greater PRD from 2001 to 2018.

Specifically, the mixing ratios of CFC-11, CFC-12, CFC-113 and H-1211 showed significant ($p < 0.005$) decreasing trends (-2.5 ± 0.4 , -3.8 ± 0.7 , -1.6 ± 0.3 , -0.4 ± 0.1 pptv/yr) while CFC-114 presented a small long-term increase (0.2 ± 0.04 pptv/yr, $p < 0.005$). The increasing pattern of CFC-114 was different from the global level which exhibited a slower declining trend in 2012-2016 compared to 2008-2012 [WMO/UNEP, 2018], but was similar to the findings of recent studies which indicated significant emissions of CFC-114 in East Asia (probably in Mainland China) in 2012-2016 [Laube et al., 2016; Vollmer et al., 2018]. Similar declining trends for CFC-11, CFC-112 and H-1211 from 1998 to 2008 were reported in this region (both HK and the inner PRD) in previous study [Zhang et al., 2010b]. Moreover, the decreasing rates of CFC-11, CFC-12, CFC-113 and H-1211 in the Greater PRD region were all higher than those at the MLO site (Table 3), indicating an efficient control under phase-out policies for CFCs and halons in South China. What's more, the mixing ratios of CFC-12, CFC-113 and H-1211

almost reached the NH background levels at MLO after 2016, while an increased emission of CFC-11 proposed in East Asia was likely responsible for the relatively slower decline of CFC-11 after 2012 [Montzka *et al.*, 2018], which also explained the larger difference in mixing ratio of CFC-11 between the Greater PRD region and the MLO site in recent years, compared to the other three species mentioned above.

In contrast, the levels of HCFC-22 and HFC-134a in the Greater PRD region increased ($p<0.05$) during the past 18 years. As shown in Figure 2 and Table 3, the annual growth rate of HCFC-22 (20.8 ± 2.7 pptv/yr) was triple that at MLO site (6.3 ± 0.03 pptv/yr), revealing continuous urban emissions. The production of HCFC-22 in 2010 and 2014 in China was 522 kt and 597.7 kt, respectively, while Hong Kong, as a “non-Article 5” party, has to reduce 90% HCFCs consumption by 2015 and 100% by 2030, indicating great challenges in managing its emissions over the next decade [CFOMIA, 2016]. The growth rate of HFC-134a (10.5 ± 1.1 pptv/yr) was twice that of the NH background (5.0 ± 0.01 pptv/yr). Because of similar thermodynamic properties to CFC-12, an insignificant ODP and a low global warming potential, HFC-134a is nowadays a substitute of CFC-12 [Su *et al.*, 2015]. In China, the HFC-134a production in 2014 almost doubled (162.3 kt) compared to that in 2010 (83.6 kt) [CFOMIA, 2016]. In summary, more reduction efforts are needed to mitigate the increase in the production and consumption of the alternative HCFCs and even HFCs after the ban of CFCs use in China.

Similar to CFC-11 and CFC-12, the CCl_4 mixing ratios decreased ($p<0.005$) at a rate of -1.6 ± 0.2 pptv/yr from 2001 to 2018, more rapidly than that at the MLO background site (-1.2 ± 0.004 pptv/yr). Moreover, CH_3CCl_3 declined from 2001 to 2018 at a rate of -3.8 ± 0.3 pptv/yr, which was twice that NH subtropical background level (-1.9 ± 0.06 pptv/yr). Xue *et al.* [2011] also reported the continuing declines in the ambient mixing ratios of CCl_4 and CH_3CCl_3 in northeast China. Furthermore, the monthly average mixing ratios of CH_3CCl_3 almost reached the NH subtropical background levels after 2016, suggesting an effective reduction of this solvent species in this region. However, the mixing ratio of C_2Cl_4 experienced a slight decrease but without significance ($p=0.09$) in the Greater PRD region despite its continuous reduction at MLO, indicating its continuous use, mainly as dry cleaning agent and degreasing fluid in the Greater PRD region.

407 Methyl chloride (CH_3Cl), the most abundant halocarbon in this study, increased
408 ($p < 0.005$) at a rate of 14.6 ± 4.7 pptv/yr from 2001 to 2018, though it decreased
409 significantly at MLO (-2.5 ± 0.4 pptv/yr), implying the un-controlled emissions of some
410 CH_3Cl sources, such as the use of solvent (including degreasant, oil paint and binding
411 agent), oceanic emissions and biomass burning activities [Guo *et al.*, 2009]. Our
412 previous study indicated that CH_3Cl was mainly emitted from the refrigeration industry
413 and biomass/biofuel burning in the PRD region [Guo *et al.*, 2009]. The exact causes for
414 its increased growth rate in these years remain unknown. Source apportionment in the
415 following section might give some hints.

416 Furthermore, CH_2Cl_2 was the fastest-growing halocarbon with a growth rate of
417 28.1 ± 11.9 pptv/yr, while its NH subtropical background growth rate was only 2.6 ± 0.07
418 pptv/yr. Similarly, growth in CH_2Cl_2 was captured in all NH regions between 1998 and
419 2012 [Leedham Elvidge *et al.*, 2015]. Increased CH_2Cl_2 levels were likely attributable
420 to the increased solvent consumption in the Greater PRD region. In addition, the mixing
421 ratio of CHCl_3 remained invariable in the past 18 years, different from the increase in
422 eastern China (increased by 49 (41-59) Gg between 2010 and 2015) but similar to the
423 insignificant change in southern China from 2007 to 2015 [Fang *et al.*, 2019]. CHCl_3
424 is mostly used as a feedstock to produce HCFC-22, tetrafluoroethene and Teflon, with
425 a limited use as an extractant for pharmaceutical products [Rossberg *et al.*, 2011].

426 The increase of CH_3Br from 2001 to 2018 was insignificant ($p = 0.16$), while CH_3I
427 showed an increasing trend at 0.08 ± 0.02 pptv/yr ($p < 0.005$). Sources of CH_3Br include
428 use as a fumigant and as a leaded fuel additive, oceanic production, biomass burning,
429 and plant and marsh emissions [Seinfeld and Pandis, 2016], but the anthropogenic uses
430 of CH_3Br have been prohibited in this region [CHP, 2018; HKEPD, 2018a; Vollmer *et al.*,
431 2009; Wang *et al.*, 2006]. Hence, natural sources, *i.e.* biomass burning and oceanic
432 emissions, are the most likely sources of CH_3Br . WMO/UNEP [2018] reported that
433 global natural sources of CH_3Br remained at a constant level from 1996 to 2016, in
434 agreement with the invariable trends in this study. CH_3I is known as an important tracer
435 of oceanic emissions [Lennartz *et al.*, 2015] and was found over the Tropical Western
436 Pacific [Fuhlbrugge *et al.*, 2016] and the Indian Ocean [Fiehn *et al.*, 2017]. Rasmussen
437 *et al.* [1982] also found that near-oceanic regions characterized by high biomass
438 productivity could yield 10 - 20 pptv of CH_3I . Besides, CH_3I is used as methylating
439 agents [Yin *et al.*, 2014]. In comparison, Yokouchi *et al.* [2012] claimed the varied

440 trends of CH₃I over the past decades, while the best estimate reported by *Carpenter and*
441 *Reimann et al.* [2014] indicated no changes in CH₃I from coastal and open-oceans
442 emissions. Thus, the growth of CH₃I was more possibly attributed to anthropogenic
443 activities. Further investigation is necessary.

444 Given that the long-term measurement data of halocarbons at a fixed site were not
445 available in this study, we integrated several field measurements of halocarbons
446 conducted in the past 18 years at rural and/or suburban sites with similar geographical
447 proximity to study the long-term variations of halocarbons. Even so, this method might
448 have some uncertainties, such as representative error, meteorological impact and
449 instrumental error. We selected the halocarbon datasets collected at geographically
450 similar sites to minimize the representative error. For the instrumental error, we used
451 the same protocols of air sampling, the same chemical standards for analysis, and the
452 same instrumental calibration and analytical method throughout the whole period. For
453 the samples analyzed in our laboratory, we followed the same procedures as those in
454 the Rowland/Blake laboratory, and frequently carried out inter-comparison between the
455 two laboratories to guarantee the data quality. Besides, the meteorological impact was
456 partially considered during the data analysis by removing samples under the influence
457 of oceanic and super-regional air masses from the whole dataset. Moreover, the
458 influence of data collected at each sampling site and in individual years, especially
459 those years with obvious step changes in mixing ratio, on the long-term trends was
460 assessed via comparison of scenarios of including and excluding data at a specific site
461 and in a specific year. The impact of various sites and different years on the variation
462 rates are listed in *Tables S2 and S3*, respectively (described in *Text S4*). In addition,
463 standard errors for all variation rates were presented to reflect the variability and
464 reliability of the results (*Figure 2 and Table 3*). Overall, despite some uncertainties, the
465 precious long-term halocarbon data could still provide useful information on temporal
466 variations of halocarbons to infer the implementation of the Montreal Protocol in the
467 Greater PRD region.

468

469

470

Table 3. Annual Growth Rate of Halocarbons in the Greater PRD region and at the MLO Site

Compounds (Chemical Formula)	Annual growth rate in the Greater PRD (2001-2018) (pptv/yr)	Annual growth rate at MLO (2001-2018) (pptv/yr)
CFC-11 (CCl ₃ F)	-2.5±0.4	-1.8±0.01
CFC-12 (CCl ₂ F ₂)	-3.8±0.7	-2.3±0.04
CFC-113 (CCl ₂ FCClF ₂)	-1.6±0.3	-0.7±0.003
CFC-114 (CClF ₂ CClF ₂)	0.2±0.04	-
HCFC-22 (CHClF ₂)	20.8±2.7	6.3±0.03
H-1211 (CF ₂ ClBr)	-0.4±0.1	-0.06±0.001
HFC-134a (CH ₂ FCF ₃)	10.5±1.1	5.0±0.01
Methyl chloride (CH ₃ Cl)	14.6±4.7	-2.5±0.4
Carbon tetrachloride (CCl ₄)	-1.6±0.2	-1.2±0.004
Methyl Chloroform(CH ₃ CCl ₃)	-3.8±0.3	-1.9±0.06
Tetrachloroethene(C ₂ Cl ₄)	-2.4±1.4 [#]	-0.06±0.004
Methylene chloride (CH ₂ Cl ₂)	28.1±11.9	2.6±0.07
Methyl bromide (CH ₃ Br)	0.2±0.1 [#]	-0.1±0.002
Chloroform (CHCl ₃)	1.1±0.8 [#]	-
Methyl iodide (CH ₃ I)	0.08±0.02	-

[#]. Non-significant, $p > 0.05$.

473

474 3.3 Emission Source Profiles in the Greater PRD region

475 [Figure 3](#) shows the source profiles resolved from the PMF modeling. Six sources were
476 identified in the Greater PRD region. The first source (F1) was distinguished by high
477 percentages of CFCs, such as CFC-11 (61.2±1.1 %), CFC-12 (63.7±2.0 %), CFC-113
478 (53.9±1.3 %) and CFC-114 (62.6±1.5 %). CFCs are mainly used as refrigerants in air
479 conditioning systems and cooling appliances. Unlike other CFCs, CFC-11 is not only a
480 refrigerant, but also extensively used as aerosol-spray propellants, foam-blowing agents
481 and solvents in the air conditioning systems in South China [[Chan et al., 2006b](#)].
482 Though CFC-11 has various applications, these uses are mostly related to the cooling
483 appliances. Thus these CFC species were grouped under the same factor. Moreover,
484 this factor explained 54.5±1.2 % of CCl₄. Prior to the Montreal Protocol, large
485 quantities of CCl₄ were used to produce CFC-11 and CFC-12 [[Simmonds et al., 1998](#)].
486 CH₃Cl also showed a high percentage (44.7±2.5 %), because it was once used as a
487 refrigerant (R-40) [[Sikdar, 2001](#)]. [Zhang et al. \[2010b\]](#) and [Guo et al. \[2009\]](#) also found
488 that CCl₄ and CH₃Cl had high correlations with CFCs in the PRD region. Interestingly,
489 37.9±0.8 % of HCFC-22 was quantified in this factor. HCFC-22 has been largely used
490 as refrigerants after the phase-out of CFCs, which might co-locate with CFCs. Overall,
491 the profile was consistent with previous studies in South China [[Guo et al., 2009](#); [Zhang](#)
492 [et al., 2010b](#)]. Hence, this source was categorized as the refrigeration industry.

493 A high mass percentage of HCFC-22 (60.4 ± 2.5 %) was observed in the second factor
494 (F2), with low loadings of other tracers. Being a replacement of CFCs, HCFC-22 is
495 mainly used in household and commercial refrigeration equipment and its emission
496 grew rapidly after the prohibition of CFCs. Thus we considered this source as the
497 replacement of CFCs.

498 The third source (F3) was inferred as solvent use in the electronic industry and dry
499 cleaning, characterized by high percentage of C_2Cl_4 (78.6 ± 2.2 %). C_2Cl_4 is primarily
500 used as an industrial cleaning solvent, metal degreasing agent and dry cleaning fluid
501 [Simmonds *et al.*, 2006]. In the study region, many dry cleaners spread in the urban
502 areas and some electronic workshops reside in industrial areas [Guo *et al.*, 2009; Zhang
503 *et al.*, 2010b]. This factor also accounted for 27.9 ± 1.1 % of CFC-113, since CFC-113
504 is often used as a solvent in electronic industry. However, the contribution of this factor
505 to CFC-113 was lower than that in 2001-2002 reported by Guo *et al.* [2009]. This was
506 attributed to the successful replacement of CFC-113 in electronic industry after the
507 implementation of Montreal Protocol [Chang *et al.*, 2008].

508 The fourth source (F4) was industrial and domestic solvent use, characterized by high
509 percentages of CH_2Cl_2 (86.0 ± 2.2 %). Because of its volatility and ability to dissolve a
510 wide range of organic compounds, CH_2Cl_2 is predominantly used as a solvent, such as
511 paint stripper, spray fluid, release agent, and cleaning solvent [Borkar *et al.*, 2010].

512 The fifth factor (F5) was considered to be feedstock in chemical manufacturing. High
513 contribution of $CHCl_3$ (85.7 ± 4.3 %) was found and other halocarbons made negligible
514 contributions to this source. As an important feedstock, the most important use of
515 $CHCl_3$ is the reaction with hydrogen fluoride to produce HCFC-22, largely for air
516 conditioning applications and the production of polytetrafluoroethylene (Teflon)
517 [Rossberg *et al.*, 2011].

518 The last source (F6) was characterized by high loading of CO (80.2 ± 2.5 %), indicating
519 its association with combustion sources. Further, 27.1 ± 2.3 % of CH_3Cl was found in
520 this source. Indeed, biomass/biofuel burning is one of the CH_3Cl sources and was
521 identified in previous studies including studies in the PRD region [Thompson *et al.*,
522 2002; Guo *et al.*, 2009; Zhang *et al.*, 2010b]. As such, this source was influenced by
523 biomass burning and/or coal burning.

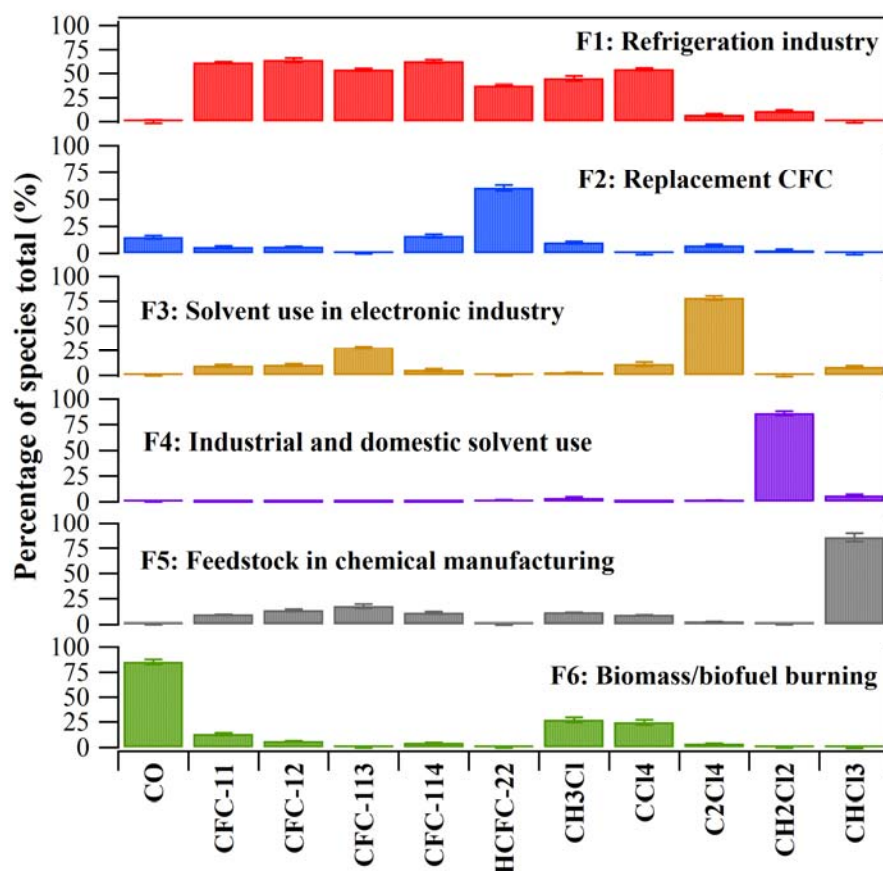


Figure 3. Source Profiles Resolved from PMF in the Greater PRD region

3.4 Temporal Variations of Halocarbon sources in the Greater PRD region

Temporal variations of source contributions to halocarbons in the Greater PRD region were extracted from the PMF results and merged into individual years, as shown in Figure 4 (in mixing ratio) and Figure S3 (in percentage). Two refrigeration application-related sources showed reverse trends in mixing ratios and in percentages during the study period. The refrigeration industry source dominated by CFCs experienced a progressive decline in the contribution to total halocarbons, from 1638.4 ± 35.4 pptv ($52.2 \pm 2.8\%$) in 2001 to 1167.1 ± 60.9 pptv ($34.5 \pm 2.4\%$) in 2018. CH_3Cl in this source decreased from ~ 540 pptv ($\sim 61\%$) in 2001 to ~ 400 pptv ($\sim 40\%$) in 2018. During the whole study period, source of refrigeration industry was always the largest contributor to the ten input halocarbons. In contrast, the contribution of replacement of CFCs, whether it is in mixing ratio or percentage, continued to increase from 197.1 ± 13.1 pptv ($6.3 \pm 0.5\%$) in 2001 to 848.1 ± 180.8 pptv ($25.1 \pm 2.2\%$) in 2018. The source of

540 replacement of CFCs contributed only $6.3\pm0.5 - 12.4\pm0.6\%$ during 2001-2013, while
541 it became the second largest contributor to total halocarbons after 2016 with
542 percentages of $22.7\pm1.0 - 25.1\pm2.2\%$, revealing successful replacement of CFCs with
543 HCFC-22 in the Greater PRD region. Since the production and consumption of CFCs
544 has been banned after mid-2007 in China (in 1996 in HK), the largest contribution of
545 refrigeration industry in recent years to halocarbons was possibly attributed to
546 background level and/or leakage from existing products or landfills.

547 Solvent use in electronic industry traced by C_2Cl_4 decreased dramatically from
548 $358.6\pm51.0 - 509.8\pm111.4$ pptv ($11.4\pm1.6 - 14.9\pm1.4\%$) during 2001-2004, to
549 209.3 ± 27.6 pptv ($5.7\pm0.4\%$) in 2013, and remained at a low level of
550 $162.6\pm58.5 - 167.5\pm20.3$ pptv ($4.8\pm1.1 - 5.3\pm0.5\%$) during 2016-2018. This result
551 confirmed the effective control of solvent use (mainly C_2Cl_4) in electronic industry in
552 this region. The other three sources, including industrial/domestic solvent use,
553 feedstock in chemical manufacturing and biomass/biofuel burning, were invariable
554 during the study period. In southern China, forest fires and cropland fires were two
555 main sources of biomass burning. [Yin et al. \[2019\]](#) estimated a decreased emission of
556 forest fire but an increase in the emission of cropland fire from 2003 to 2017. Total
557 biomass burning emission in China decreased remarkably in 2015-2016, but bounced
558 in 2017. The insignificant variation of biomass burning emission was similar to our
559 findings. [Feng et al. \[2018\]](#) reported that anthropogenically-emitted CH_2Cl_2 increased
560 from 2005 to 2016, contrary to our findings. There might be two causes. Firstly, the
561 estimation by [Feng et al. \[2018\]](#) was for the whole China while local situations in
562 different regions might be different. Secondly, the insignificant changes found in this
563 study might be owing to limited sampling years. Since CH_2Cl_2 , $CHCl_3$ and CH_3Cl did
564 not decline, their emissions should be controlled through more stringent measures.

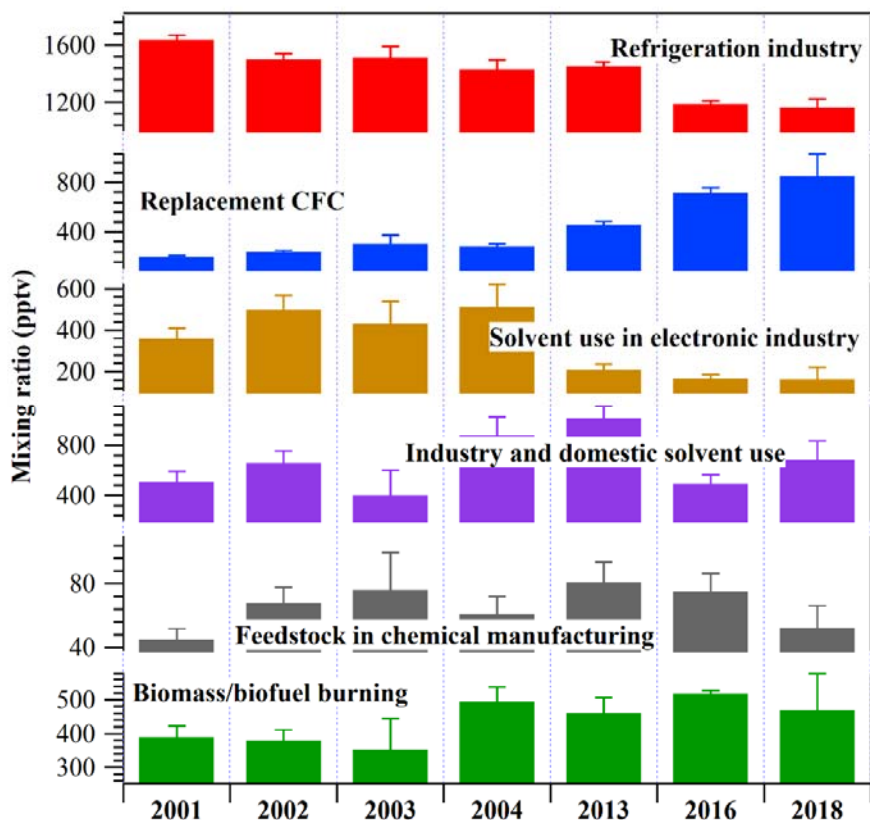


Figure 4. Source Contributions to the Total Halocarbons in the Greater PRD region (mean \pm 95% confidence interval) in 2001, 2002, 2003, 2004, 2013, 2016, 2018

3.5 Halocarbon Emission Estimates in the Greater PRD region

The CO emissions used for halocarbon emission estimations in the inner PRD and HK were 7305.40 Gg in 2012 (latest available CO emission inventory) with an uncertainty of $\pm 68\%$ and 58.52 Gg with an uncertainty of $\pm 29\%$ in 2016, respectively [Zhong *et al.*, 2018; HKEPD, 2018b]. Here we assumed that total CO emission in 2016 in the Greater PRD region was 7363.92 Gg. Scatter plots of mixing ratios of halocarbons versus CO values are shown in Figure S4. The slopes were applied for the emission estimation. As shown in Table 4, the total halocarbon emission in the Greater PRD region was estimated to be 46.5 ± 16.7 Gg. By comparison, our results were basically consistent with or lower than the estimated halocarbon emissions in 2001, 2004, 2009 and 2010 in previous studies [Guo *et al.*, 2009; Shao *et al.*, 2011; Zhang *et al.*, 2014], except for CFC-11, CFC-114, HCFC-22, CH_3Cl and CH_2Cl_2 . Specifically, our estimated

emissions of CFC-11 and CFC-114 in 2016 were higher than those in 2004, 2009 and 2010, suggesting possible new sources of these two species in recent years.

Among the target halocarbons in the Greater PRD region, the highest estimated emission was from CH_2Cl_2 (21 ± 7.8 Gg) accounting for 45% of the total emissions, implying significant contributions of solvent use and chemical manufacturing to ambient halocarbons. The estimated emission of CH_2Cl_2 was much lower than that in 2001 but higher than the emission in 2004. CH_3Cl had similar temporal variations in estimated emission to CH_2Cl_2 . Thus, tighter control measures are needed to reduce these two species. The estimated CFCs emission (2.6 ± 1.3 Gg) was much lower than those of HCFC-22 and HFC-134a (15 ± 3.9 Gg), in line with the successful replacement of CFCs with HCFCs and HFCs in this region. Moreover, the estimated emission of HCFC-22 was higher than previous levels, in agreement with its increasing ambient levels. *Saikawa et al.* [2012] also estimated the increase in HCFC-22 emissions from Article-5 Asia during 2005-2010. Since both HCFC-22 and HFC-134a are in the restriction list and need to be phased out in the next phase, controlling their emissions would pose a challenge to this region. The emissions of other halocarbons (mainly used in solvents and chemical manufacturing) in the Greater PRD region (29 ± 11 Gg) accounted for 62% of the total emissions, implying increased use of short-lived halocarbons. Among these short-lived species, the estimated emissions of CH_3CCl_3 and C_2Cl_4 decreased dramatically compared to previous estimations, mainly due to effective regulations. It is noteworthy that the technique used here is a very rough estimate of halocarbon emissions with some uncertainties. As the sources, lifetime and chemical properties of halocarbons and CO have some differences, the estimated halocarbon emissions would have some uncertainties using the ratio of the enhancement of halocarbon to CO. Furthermore, the CO emission inventory also has some uncertainties. Therefore, other approaches such as a better reference species for inter-species correlations [Wu et al., 2014] and/or inverse model simulations should be explored in future studies [Fang et al., 2015].

609

Table 4. Measured Halocarbons: CO Relationships and Estimated Halocarbon Emissions (in Gg) in 2016

Compounds	M (g·mol ⁻¹)	The Greater PRD region (n=218) ^a			Reference value in PRD region (Gg)			
		Ratio	Pearson correlation coefficient	Emission estimation (Gg)	2001 ^d	2004 ^e	2009 ^f	2010 ^g
CFC-11	137.36	0.05±0.02	0.18	1.7±0.9	1.0 ± 0.3	0.4±0.2	0.5±0.1	0.9
CFC-12	120.91	0.01±0.005	0.17	0.4±0.2	1.5 ± 0.4	1.6±1.0	1.2±0.3	1.6
CFC-113	187.38	0.004±0.002	0.11	0.2±0.1	0.9 ± 0.3	0.0±0.0	0.2±0.04	
CFC-114	170.92	0.006±0.001	0.48	0.3±0.05		0.0±0.0	0.1±0.01	
HCFC-22	86.47	0.54±0.13	0.41	12±3.0	2.2±1.2	3.5±2.2	2.5±0.7	10.7
HFC-134a	102.03	0.12±0.03	0.22	3.1±0.9				1.3
CH ₃ Cl	50.49	0.35±0.33	0.18	4.6±2.0	2.8 ± 0.5	0.6±0.4		
CH ₃ CCl ₃	133.40	0.006±0.0008	0.41	0.2±0.03		0.4±0.2		
CH ₂ Cl ₂	84.93	1.02±0.35	0.29	21±7.8	42.8 ± 7.2	7.0±4.6		
CCl ₄	153.81	0.02±0.004	0.24	0.8±0.2	0.7 ± 0.2	1.1±0.7		
CHCl ₃	119.37	0.03±0.02	0.34	0.9±0.6	2.4 ± 1.8	0.8±0.6		
C ₂ Cl ₄	165.82	0.03±0.02	0.38	1.3±0.9	7.3 ± 1.5	2.3±1.5		
Sum				46.5±16.7				
CO	28.01			7363.92 ^{b, c}				

610 a. The parameter n is the number of samples; the ratio is the regression slope of Δ halocarbon against Δ CO (pptv/ppbv); Pearson correlation coefficient is the parameter of
611 determination of the ratio; b. HKEPD public data, https://www.epd.gov.hk/epd/sc_chi/environmentinhk/air/data/emission_inve_co_C.html; c. [Zhong et al., 2018]; d. [Guo et
612 al., 2009]; e. [Shao et al., 2011]; f. [Zhang et al., 2014]; g. [Wu et al., 2014].

613

614

615 4. Conclusions

616 In this study, we investigated long-term temporal variations, source changes and
617 emissions of halocarbons in the Greater PRD region from 2001 to 2018 based on 1505
618 whole air samples. CH₃Cl was the most abundant halocarbon, followed by CH₂Cl₂,
619 HCFC-22 and CFC-12. Long-term analysis indicated significantly ($p<0.005$)
620 decreasing trends of CFC-11, CFC-12, CFC-113 and H-1211, and increases in HFC-
621 134a, HCFC-22 and CFC-114 in the past 18 years in the Greater PRD region. Compared
622 to the variation rates at the background MLO site, faster declines of CFCs (CFC-11,
623 CFC-12 and CFC-113) in the Greater PRD region suggested effectiveness of
624 previous/ongoing control measures under the Montreal Protocol. In addition, though
625 mixing ratios of CCl₄, CH₃CCl₃ and C₂Cl₄ decreased significantly ($p<0.005$), the values
626 of CH₂Cl₂, CH₃Cl, and CH₃I remarkably increased ($p<0.05$), indicating more stringent
627 control regulations are still needed. In particular, CH₂Cl₂ showed the highest growth
628 rate in the Greater PRD region.

629 Source apportionment results revealed that halocarbons in the Greater PRD were
630 mainly used in refrigerants applications, including refrigeration industry and
631 replacement of CFCs, in these years. Contribution of refrigeration industry showed a
632 decline during 2001-2018, while source of replacement of CFCs increased
633 progressively. Furthermore, solvent use in electronic industry made invariable
634 contributions during 2001-2004, but the contribution dramatically decreased in 2013
635 and remained at a low level in recent years. The contributions of industrial/domestic
636 solvent use, feedstock in chemical manufacturing and biomass/biofuel burning did not
637 change significantly during the study period and required further control measures.

638 According to the ratio of halocarbon to CO and the CO emissions, the total halocarbon
639 emissions in 2016 were estimated to be 46.5 ± 16.7 Gg for the Greater PRD. The
640 estimated emission of CH₂Cl₂ was the largest, suggesting significant contributions of
641 industrial and domestic solvent use to ambient halocarbons in the Greater PRD region.
642 The findings advanced our knowledge on halocarbons in the Greater PRD region.

643 Acknowledgements

644 The authors thank the financial support by the National Key R&D Program of China
645 via grant No. 2017YFC0212001. This study was supported by the Research Grants
646 Council of the Hong Kong Special Administrative Region Government via grants

647 PolyU152052/14E, PolyU152052/16E and CRF/C5004-15E, and the Public Policy
648 Research Funding Scheme from Policy Innovation and Co-ordination Office of the
649 Hong Kong Special Administrative Region Government (Project Number:
650 2017.A6.094.17D). The data at Mauna Loa station provided by the National Oceanic
651 and Atmospheric Administration (NOAA) is highly appreciated.

652 References

- 653 Barletta, B., Meinardi, S., Simpson, I. J., Rowland, F. S., Chan, C. Y., Wang, X. M.,
654 Zou, S. C., Chan, L. Y. and Blake, D. R. (2006), Ambient halocarbon mixing ratios in
655 45 Chinese cities, *Atmospheric Environment*, 40(40), 7706-7719.
- 656 Blake, N. J., Blake, D. R., Sive, B. C., Chen, T.-Y., Rowland, F. S., Collins Jr., J. E.,
657 Satche, G. W. and Anderson, B. E. (1996), Biomass burning emissions and vertical
658 distribution of atmospheric methyl halides and other reduced carbon gases in the South
659 Atlantic region, *Journal of Geophysical Research: Atmospheres*, 101, 24151-24164.
- 660 Borkar, C., Tomar, D. and Gumma S. (2010), Adsorption of Dichloromethane on
661 Activated Carbon, *Journal of Chemical & Engineering Data*, 55(4), 1640-1644.
- 662 Carpenter, L. J., Reimann, S., Burkholder, J. B., Clerbaux, C., Hall, B. D., Hossaini, R.,
663 Laube, J. C. and Yvon-Lewis, S. A. (2014), Update on ozone-depleting substances
664 (ODSs) and other gases of interest to the Montreal protocol, *Scientific assessment of*
665 *ozone depletion: 2014*, 1-1.
- 666 Chan, C. Y., Tang, J. H., Li, Y. S. and Chan, L. Y. (2006b), Mixing ratios and sources
667 of halocarbons in urban, semi-urban and rural sites of the Pearl River Delta, South
668 China, *Atmospheric Environment*, 40(38), 7331-7345.
- 669 Chan, L. Y. and Chu, K. W. (2007), Halocarbons in the atmosphere of the industrial-
670 related Pearl River Delta region of China, *Journal of Geophysical Research:*
671 *Atmospheres*, 112(D4), doi:doi:10.1029/2006JD007097.
- 672 Chan, J. Y., Tang, J. H., Li, Y. S. and Chan, L. Y. (2006a), Mixing ratios and sources
673 of halocarbons in urban, semi-urban and rural sites of the Pearl River Delta, South
674 China, *Atmospheric Environment*, 40(38), 7331-7345.
- 675 Chang, C. C., Lai, C. H., Wang, C. H., Liu, Y., Shao, M., Zhang, Y. and Wang, J. L.
676 (2008), Variability of ozone depleting substances as an indication of emissions in the
677 Pearl River Delta, China, *Atmospheric Environment*, 42(29), 6973-6981.
- 678 CFOMIA (2016), China Fluorosilicone Organic Materials Industry Association,
679 China's fluorine chemical industry "13th five-year" development plan,
680 <http://www.sif.org.cn/>.
- 681 CHP (2018), The Health Effects of Air Pollution,
682 <https://www.chp.gov.hk/en/healthtopics/content/460/3557.html>.
- 683 Devotta, S., Waghmare, A. V., Sawant, N. N. and Domkundwar, B. M. (2001),
684 Alternatives to HCFC-22 for air conditioners, *Applied Thermal Engineering*, 21(6),
685 703-715.
- 686 ESRL (2018), [https://www.esrl.noaa.gov/gmd/dv/data/index.php?category=](https://www.esrl.noaa.gov/gmd/dv/data/index.php?category=Halocompounds)
687 [Halocompounds](https://www.esrl.noaa.gov/gmd/dv/data/index.php?category=Halocompounds).

688 Fang, X. K., Wu, J., Su, S. S., Han, J. R., Wu, Y. S., Shi, Y. H., Wan, D., Sun, X. Z.,
689 Zhang, J. B. and Hu, J. X. (2012), Estimates of major anthropogenic halocarbon
690 emissions from China based on interspecies correlations, *Atmospheric Environment*, 62,
691 26-33.

692 Fang, X. K., Stohl, A., Yokouchi, Y., Kim, J., Li, S., Saito, T., Park, S. and Hu, J. (2015),
693 Multiannual top-down estimate of HFC-23 emissions in East Asia, *Environmental*
694 *Science & Technology*, 49(7), 4345-4353.

695 Fang X. K., Park, S., Saito, T., Tunnicliffe, R., Ganesan, A. L., Rigby, M., Li, S.
696 L., Yokouchi, Y., Fraser, P. J., Harth, C. M., Krummel, P. B., Mühle,
697 J., O'Doherty, S., Salameh, P. K., Simmonds, P. G., Weiss, R. F., Young,
698 D., Lunt, M. F., Manning, A. J., Gressent, A. and Prinn, R. G. (2019), Rapid
699 increase in ozone-depleting chloroform emissions from China, *Nature Geoscience*, 12,
700 89-93.

701 Fang, X., Velders, G. J. M., Ravishankara, A. R., Molina, M. J., Hu, J. and Prinn, R. G.
702 (2016), Hydrofluorocarbon (HFC) Emissions in China: An Inventory for 2005–2013
703 and Projections to 2050, *Environmental Science & Technology*, 50(4), 2027-2034.

704 Feng, Y., Bie, P., Wang, Z., Wang, L. and Zhang, J. (2018), Bottom-up anthropogenic
705 dichloromethane emission estimates from China for the period 2005–2016 and
706 predictions of future emissions, *Atmospheric Environment*, 186, 241-247.

707 Fiehn, A., Quack, B., Hepach, H., Fuhlbrügge, S., Tegtmeier, S., Toohey, M., Atlas, E.
708 and Krüger, K. (2017), Delivery of halogenated very short-lived substances from the
709 West Indian Ocean to the stratosphere during Asian summer monsoon, *Atmospheric*
710 *Chemistry and Physics*, 17, 6723-6741.

711 Fortems-Cheiney, A., Chevallier, F., Saunois, M., Pison, I., Bousquet, P., Cressot, C.,
712 Wang, H. J., Yokouchi, Y. and Artuso, F. (2013), HCFC-22 emissions at global and
713 regional scales between 1995 and 2010: Trends and variability, *Journal of Geophysical*
714 *Research: Atmospheres*, 118(13), 7379-7388.

715 Fraser, P. J., Oram, D. E., Reeves, C. E., Penkett, S. A. and McCulloch, A. (1999),
716 Southern Hemispheric halon trends (1978–1998) and global halon emissions, *Journal*
717 *of Geophysical Research: Atmospheres*, 104(D13), 15985-15999.

718 Fuhlbrügge, S., Quack, B., Tegtmeier, S., Atlas, E., Hepach, H., Shi, Q., Raimund, S.
719 and Krüger, K. (2016), The contribution of oceanic halocarbons to marine and free
720 troposphere air over the tropical West Pacific, *Atmospheric Chemistry and Physics*,
721 16(12), 7569-7585.

722 Gentner, D. R., Miller, A. M. and Goldstein, A. H. (2010), Seasonal Variability in
723 Anthropogenic Halocarbon Emissions, *Environmental Science & Technology*, 44(14),
724 5377-5382.

725 Godish, T. (2003), Air quality (Fourth edition), Lewis Books, Boca Raton, Florida.

726 Grubb, M., Vrolijk, C. and Brack, D. (1999), The Kyoto Protocol. A guide and
727 assessment, *Molecular Ecology*, 13(8), 2121-2133.

728 Guo, H., Ding, A.J., Wang, T., Simpson, I.J., Blake, D.R., Barletta, B., Meinardi, S.,
729 Rowland, F. S., Saunders, S.M., Fu, T.M., Hung, W.T. and Li, Y.S. (2009), Source
730 origins, modeled profiles, and apportionments of halogenated hydrocarbons in the
731 greater Pearl River Delta region, southern China, *Journal of Geophysical Research:*
732 *Atmospheres*, 114, 19, D11302.

733 Harrison, J. J., Boone, C. D. and Bernath, P. F. (2017), New and improved infra-red
 734 absorption cross sections and ACE-FTS retrievals of carbon tetrachloride (CCl₄),
 735 *Journal of Quantitative Spectroscopy and Radiative Transfer*, 186, 139-149.

736 HKEPD, (2018a), [https://www.epd.gov.hk/epd/english/environmentinhk/air/data/
 737 phase_out.html](https://www.epd.gov.hk/epd/english/environmentinhk/air/data/phase_out.html).

738 HKEPD, (2018b), [https://www.epd.gov.hk/epd/sc_chi/environmentinhk/air/data/
 739 emission_inve_co_C.html](https://www.epd.gov.hk/epd/sc_chi/environmentinhk/air/data/emission_inve_co_C.html).

740 Hundy, G.F., Trott, A.R. and Welch, T.C. (2016), Refrigeration, Air Conditioning and
 741 Heat Pumps. 5th Edition, Elsevier Ltd, London.

742 Hurst D. F., Romashkin, P. A., Elkins, J. W., Oberländer, E. A., Elansky, N. F., Belikov,
 743 I. B., Granberg, I. G., Golitsyn, G. S., Grisenko, A. M., Brenninkmeijer, C. A. M. and
 744 Crutzen, P. J. (2004), Emissions of ozone - depleting substances in Russia during 2001,
 745 *J. Geophys. Res.*, 109, D14303.

746 Hurst, D. F., Lin, J. C., Romashkin, P. A., Daube, B. C., Gerbig, C., Matross, D. M.,
 747 Wofsy, S. C., Hall, B. D. and Elkins, J. W. (2006), Continuing global significance of
 748 emissions of Montreal Protocol-restricted halocarbons in the United States and Canada,
 749 *Journal of Geophysical Research: Atmospheres*, 111(D15), 15.

750 Kock, H. H., Bieber, E., Ebinghaus, R., Spain, T. G., and Thees, B. (2005), Comparison
 751 of long-term trends and seasonal variations of atmospheric mercury concentrations at
 752 the two European coastal monitoring stations Mace Head, Ireland, and Zingst, Germany,
 753 *Atmospheric Environment*, 39(39), 7549-7556.

754 Laube, J. C., Mohd Hanif, N., Martinerie, P., Gallacher, E., Fraser, P.J., Langenfelds,
 755 R., Brenninkmeijer, C.A.M., Schwander, J., Witrant, E., Wang, J.-L., Ou-Yang, C.F.,
 756 Gooch, L.J., Reeves, C.E., Sturges, W.T. and Oram, D.E. (2016), Tropospheric
 757 observations of CFC-114 and CFC-114a with a focus on long-term trends and
 758 emissions, *Atmospheric Chemistry & Physics*, 16(23), 15347-15358.

759 Lee, S. C., Chiu, M. Y., Ho, K. F., Zou, S. C. and Wang, X. (2002), Volatile organic
 760 compounds (VOCs) in urban atmosphere of Hong Kong, *Chemosphere*, 48(3), 375-382.

761 Leedham Elvidge, E. C., Oram, D. E., Laube, J. C., Baker, A. K., Montzka, S. A.,
 762 Humphrey, S., O'Sullivan, D. A. and Brenninkmeijer, C. A. M. (2015), Increasing
 763 concentrations of dichloromethane, CH₂Cl₂, inferred from CARIBIC air samples
 764 collected 1998–2012, *Atmospheric Chemistry & Physics*, 15(4), 1939-1958.

765 Lennartz, S. T., Krysztofiak, G., Marandino, C. A., Sinnhuber, B.-M., Tegtmeier, S.,
 766 Ziskal, F., Hossaini, R., Krüger, K., Montzka, S. A., Atlas, E., Oram, D. E., Keber, T.,
 767 Bönisch, H. and Quack, B. (2015), Modelling marine emissions and atmospheric
 768 distributions of halocarbons and dimethyl sulfide: the influence of prescribed water
 769 concentration vs. prescribed emissions, *Atmospheric Chemistry & Physics*, 15(20),
 770 11753-11772.

771 Li, S., Park, M.-K., Jo, C. O. and Park, S. (2017), Emission estimates of methyl chloride
 772 from industrial sources in China based on high frequency atmospheric observations,
 773 *Journal of Atmospheric Chemistry*, 74(2), 227-243.

774 Ling, Z. H., Guo, H., Lam, S. H. M., Saunders, S. M. and Wang, T. (2014), Atmospheric
 775 photochemical reactivity and ozone production at two sites in Hong Kong: Application
 776 of a Master Chemical Mechanism–photochemical box model, *Journal of Geophysical
 777 Research: Atmospheres*, 119(17), 10567-10582.

778 Lunt, M. F., Park, S., Li, S., Henne, S., Manning, A. J., Ganesan, A. L., Simpson, I. J.,
 779 Blake, D. R., Liang, Q., O'Doherty, S., Harth, C. M., Mühle, J., Salameh, P. K., Weiss,
 780 R. F., Krummel, P. B., Fraser, P. J., Prinn, R. G., Reimann, S. and Rigby, M. (2018),
 781 Continued emissions of the ozone-depleting substance carbon tetrachloride from
 782 Eastern Asia, *Geophysical Research Letters*, 45(20), 11324-11430.

783 Maione, M., Giostra, U., Arduini, J., Furlani, F., Bonasoni, P., Cristofanelli, P., Laj, P.
 784 and Vuillermoz, E. (2011), Three-year observations of halocarbons at the Nepal
 785 Climate Observatory at Pyramid (NCO-P, 5079 m a.s.l.) on the Himalayan range,
 786 *Atmospheric Chemistry & Physics*, 11(7), 3431-3441.

787 Montzka, S. A., Dutton, G. S., Yu, P., Ray, E., Portmann, R. W., Daniel, J. S., Lambert,
 788 K., Hall, B. D., Mondeel, D., Siso, C., Nance, J. D., Rigby, M., Manning, A. J., Hu, L.,
 789 Moore, F., Miller, B. R. and Elkins, J. W. (2018), An unexpected and persistent increase
 790 in global emissions of ozone-depleting CFC-11, *Nature*, 557, 413-417.

791 Oram, D. E., Ashfold, M. J., Laube, J. C., Gooch, L. J., Humphrey, S., Sturges, W. T.,
 792 Leedham-Elvidge, E., Forster, G. L., Harris, N. R. P., Mead, M. I., Samah, A. A., Phang,
 793 S. M., Ou-Yang, C. F., Lin, N. H., Wang, J. L., Baker, A. K., Brenninkmeijer, C. A. M.
 794 and Sherry, D. (2017), A growing threat to the ozone layer from short-lived
 795 anthropogenic chlorocarbons, *Atmospheric Chemistry & Physics*, 17(19), 11929-11941.

796 Ou-Yang, C. F., Chang, C. C., Chen, S. P., Chew, C., Lee, B. R., Chang, C. Y., Montzka,
 797 S. A., Dutton, G. S., Butler, J. H. and Elkins, J. W. (2015), Changes in the levels and
 798 variability of halocarbons and the compliance with the Montreal Protocol from an urban
 799 view, *Chemosphere*, 138, 438-446.

800 Paatero, P. (1997), Least squares formulation of robust non-negative factor analysis,
 801 *Chemometrics & Intelligent Laboratory Systems*, 37(1), 23-35.

802 Palmer, P. I., Jacob, D. J., Mickley, L. J., Blake, D. R., Sachse, G. W., Fuelberg, H. E.
 803 and Kiley, C. M. (2003), Eastern Asian emissions of anthropogenic halocarbons
 804 deduced from aircraft concentration data, *Journal of Geophysical Research:*
 805 *Atmospheres*, 108(D24), 4753.

806 Prinn, R. G., Cunnold, D. M., Rasmussen, R., Simmonds, P. G., Alyea, F. N., Crawford,
 807 A., Fraser, P. J. and Rosen, R. (1987), Atmospheric trends in methylchloroform and the
 808 global average for the hydroxyl radical, *Science*, 238, 945-950.

809 Qin, D. (2007), Decline in the concentrations of chlorofluorocarbons (CFC-11, CFC-
 810 12 and CFC-113) in an urban area of Beijing, China, *Atmospheric Environment*, 41(38),
 811 8424-8430.

812 Rasmussen, R. A., Khalil, M. A. K., Gunawardena, R. and Hoyt, S. D. (1982),
 813 Atmospheric Methyl Iodide (CH₃I), *Journal of Geophysical Research: Oceans*, 87(C4),
 814 3086-3090.

815 Rigby, M., Park, S., Saito, T., Western, L. M., Redington, A. L., Fang, X., Henne, S.,
 816 Manning, A. J., Prinn, R. G., Dutton, G. S., Fraser, P. J., Ganesan, A. L., Hall, B. D.,
 817 Harth, C. M., Kim, J., Kim, K. –R., Krummel, P. B., Lee, T., Li, S., Liang, Q., Lunt, M.
 818 F., Montzka, S. A., Mühle, J., O'Doherty, S., Park, M. –K., Reimann, S., Salameh, P.
 819 K., Simmonds, P., Tunnicliffe, R. L., Weiss, R. F., Yokouchi, Y. and Young, D. (2019),
 820 Increase in CFC-11 emissions from eastern China based on atmospheric observations.,
 821 *Nature*, 569(7757), 546.

822 Rossberg, M., Lendle, W., Pfeleiderer, G., Tögel, A., Torkelson, T. R. and Beutel, K. K.
823 (2011), *Chloromethanes*, Wiley-VCH Verlag GmbH & Co. KGaA.

824 Saikawa, E., Rigby, M., Prinn, R. G., Montzka, S. A., Miller, B. R., Kuijpers, L. J.,
825 Fraser, P. J. B., Vollmer, M. K., Saito, T., Yokouchi, Y., Harth, C. M., Mühle, J., Weiss,
826 R. F., Salameh, P. K., Kim, J., Li, S., Park, S., Kim, K. -R., Young, S., O'Doherty, S.,
827 Simmonds, P. G., McCulloch, A., Krummel, P. B., Steele, L. P., Lunder, C., Hermansen,
828 O., Maione, M., Arduini, J., Yao, B., Zhou, L. X., Wang, H. J., Elkins, J. W. and Hall,
829 B. (2012), Global and regional emission estimates for HCFC-22, *Atmospheric*
830 *Chemistry and Physics*, 12(21), 10033-10050.

831 Scheeren, H. A., Lelieveld, J., De Gouw, J. A., Van der Veen, C. and Fischer, H. (2002),
832 Methyl chloride and other chlorocarbons in polluted air during INDOEX, *Journal of*
833 *Geophysical Research: Atmospheres*, 107(D19), INX2-14.

834 Seinfeld, J.H. and Pandis, S.N. (2016), *Atmospheric Chemistry and Physics: From Air*
835 *Pollution to Climate Change*. John Wiley & Sons, Hoboken.

836 Shao, M., Huang, D., Gu, D., Lu, S., Chang, C. and Wang, J. (2011), Estimate of
837 anthropogenic halocarbon emission based on measured ratio relative to CO in the Pearl
838 River Delta region, China, *Atmospheric Chemistry & Physics*, 11(10), 5011-5025.

839 Sikdar, S. K. (2001), Process design tools for the environment, *Crc Press*, 99(2), 223-
840 223.

841 Simmonds, P. G., Cunnold, D. M., Weiss, R. F., Prinn, R. G., Fraser, P. J., McCulloch,
842 A., Alyea, F. N. and O'Doherty, S. (1998), Global trends and emission estimates of
843 CCl₄ from in situ background observations from July 1978 to June 1996, *Journal of*
844 *Geophysical Research: Atmospheres*, 103(D13), 16017-16027.

845 Simmonds, P. G., Manning, A. J., Cunnold, D. M., McCulloch, A., O'Doherty, S.,
846 Derwent, R. G., Krummel, P. B., Fraser, P. J., Dunse, B., Porter, L. W., Wang, R. H. J.,
847 Greally, B. R., Miller, B. R., Salameh, P., Weiss, R. F. and Prinn, R. G. (2006), Global
848 trends, seasonal cycles, and European emissions of dichloromethane, trichloroethene,
849 and tetrachloroethene from the AGAGE observations at Mace Head, Ireland, and Cape
850 Grim, Tasmania, *Journal of Geophysical Research: Atmospheres*, 111(D18), D18304.

851 Simmonds, P. G., Rigby, M., McCulloch, A., O'Doherty, S., Young, D., Mühle, J.,
852 Krummel, P. B., Steele, P., Fraser, P. J. and Manning, A. J. (2017), Changing trends
853 and emissions of hydrochlorofluorocarbons (HCFCs) and their hydrofluorocarbon
854 (HFCs) replacements, *Atmospheric Chemistry & Physics*, 17(7), 4641-4655.

855 Simpson, I. J., Meinardi, S., Blake, N. J., Rowland, F. S. and Blake, D. R. (2004), Long-
856 term decrease in the global atmospheric burden of tetrachloroethene (C₂Cl₄),
857 *Geophysical Research Letters*, 31(8), L08108.

858 Stratosphere-troposphere Processes And their Role in Climate (SPARC) (2017),
859 SPARC Report on the Mystery of Carbon Tetrachloride, Q. Liang, P. A. Newman, and
860 S. Reimann (Eds.), *SPARC Report No. 7*, WCRP-13/2016.

861 Su, S. S., Fang, X. K., Li, L., Wu, J., Zhang, J. B., Xu, W. G. and Hu, J. X. (2015),
862 HFC-134a emissions from mobile air conditioning in China from 1995 to 2030,
863 *Atmospheric Environment*, 102, 122-129.

864 Thompson, A. E., Anderson, R. S., Rudolph, J. and Huang, L. (2002), Stable carbon
865 isotope signatures of background tropospheric chloromethane and CFC113,
866 *Biogeochemistry*, 60(2), 191-211.

867 Varner, R. K., Crill, P. M., Talbot, R. W. and Shorter, J. H. (1999), An estimate of the
 868 uptake of atmospheric methyl bromide by agricultural soils, *Geophysical research*
 869 *letters*, 26(6), 727-730.

870 Velders, G. J. M., Fahey, D. W., Daniel, J. S., Andersen, S. O. and McFarland, M.
 871 (2015), Future atmospheric abundances and climate forcings from scenarios of global
 872 and regional hydrofluorocarbon (HFC) emissions, *Atmospheric Environment*, 123, 200-
 873 209.

874 Vollmer, M. K., Zhou, L. X., Grealley, B. R., Henne, S., Yao, B., Reimann, S., Stordal,
 875 F., Cunnold, D. M., Zhang, X. C., Maione, M., Zhang, F., Huang, J. and Simmonds, P.
 876 G. (2009), Emissions of ozone-depleting halocarbons from China. *Geophysical*
 877 *research letters*, 36 (15), L15823.

878 Vollmer, M. K., Young, D., Trudinger, C. M., Mühle, J., Henne, S., Rigby, M., Park,
 879 S., Li, S., Guillevis, M., Mitrevski, B., Harth, C. M., Miller, B. R., Reimann, S., Yao,
 880 B., Steele, L. P., Wyss, S. A., Lunder, C. R., Arduini, J., McCulloch, A., Wu, S. H.,
 881 Rhee, T. S., Wang, R. H. J., Salameh, P. K., Hermansen, O., Hill, M., Langenfelds, R.
 882 L., Ivy, D., O'Doherty, S., Krummel, P. B., Maione, M., Etheridge, D. M., Zhou, L. X.,
 883 Fraser, P. J., Prinn, R. G., Weiss, R. F. and Simmonds, P. G. (2018), Atmospheric
 884 histories and emissions of chlorofluorocarbons CFC-13 (CClF₃), Σ CFC-114 (C₂Cl₂F₄),
 885 and CFC-115 (C₂ClF₅), *Atmospheric Chemistry & Physics*, 18(2), 979-1002.

886 Wang, C., Shao, M., Huang, D. K., Lu, S. H., Zeng, L. M., Hu, M. and Zhang, Q. (2014),
 887 Estimating halocarbon emissions using measured ratio relative to tracers in China,
 888 *Atmospheric Environment*, 89, 816-826.

889 Wang, T., Guo, H., Blake, D. R., Kwok, Y. H., Simpson, I. J. and Li, Y. S. (2005),
 890 Measurements of Trace Gases in the Inflow of South China Sea Background Air and
 891 Outflow of Regional Pollution at Tai O, Southern China, *Journal of Atmospheric*
 892 *Chemistry*, 52(3), 295.

893 Wang, W., Liu, X., Zhao, L., Guo, D., Tian, X., & Adams, F. (2006), Effectiveness of
 894 leaded petrol phase-out in Tianjin, China based on the aerosol lead concentration and
 895 isotope abundance ratio, *Science of the Total Environment*, 364 (1-3), 175-187.

896 Wang, Y., Guo, H., Zou, S. C., Lyu, X. P., Ling, Z. H., Cheng, H. R. and Zeren Y. Z.
 897 (2018), Surface O₃ photochemistry over the South China Sea: Application of a near-
 898 explicit chemical mechanism box model, *Environmental Pollution*, 234, 155-166.

899 WMO/UNEP (2018), Scientific Assessment of Ozone Depletion: 2018, Global Ozone
 900 Research and Monitoring Project. Report no. 58.88.
 901 <https://www.esrl.noaa.gov/csd/assessments/ozone/2018/>.

902 Wu, J., Fang, X. K., Xu, W. Y., Wan, D., Shi, Y. H., Su, S. S., Hu, J. X. and Zhang, J.
 903 B. (2013), Chlorofluorocarbons, hydrochlorofluorocarbons, and hydrofluorocarbons in
 904 the atmosphere of four Chinese cities, *Atmospheric Environment*, 75, 83-91.

905 Wu, J., Fang, X.K., Martin, J. W., Zhai, Z.H., Su, S.S., Hu, X., Han, J.R., Lu, S. H.,
 906 Wang, C., Zhang, J. B. and Hu, J. X. (2014), Estimated emissions of
 907 chlorofluorocarbons, hydrochlorofluorocarbons, and hydrofluorocarbons based on an
 908 interspecies correlation method in the Pearl River Delta region, China, *Science of the*
 909 *Total Environment*, 470–471, 829–834.

910 Xue, L. K., Wang, T., Simpson, I. J., Ding, A. J., Gao, J., Blake, D. R., Wang, X. Z.,
 911 Wang, W. X., Lei, H. C. and Jin, D. Z. (2011), Vertical distributions of non-methane

hydrocarbons and halocarbons in the lower troposphere over northeast China, *Atmospheric environment*, 45(36), 6501-6509.

Yin, L., Du, P., Zhang, M., Xu, T. and Song, Y. (2019), Estimation of emissions from biomass burning in China (2003–2017) based on MODIS fire radiative energy data, *Biogeosciences*, 16(7), 1629-1640.

Yin, Y., Li, Y., Tai, C., Cai, Y. and Jiang, G. (2014), Fumigant methyl iodide can methylate inorganic mercury species in natural waters, *Nature Communications*, 5, 4633.

Yokouchi, Y., Noijiri, Y., Barrie, L. A., Toom-Sauntry, D., Machida, T., Inuzuka, Y., Akimoto, H., Li, H. -J., Fujinuma, Y., and Aoki, S. (2000), A strong source of methyl chloride to the atmosphere from tropical coastal land, *Nature*, 403(6767), 295-298.

Yokouchi, Y., Toom-Sauntry, D., Yazawa, K., Inagaki, T. and Tamaru, T. (2002), Recent decline of methyl bromide in the troposphere, *Atmospheric Environment*, 36(32), 4985-4989.

Yokouchi, Y., Nojiri, Y., Toom-Sauntry, D., Fraser, P., Inuzuka, Y., Tanimoto, H., Nara, H., Murakami, R. and Mukai, H. (2012), Long-term variation of atmospheric methyl iodide and its link to global environmental change, *Geophysical research letters*, 39(23), L23805, doi: 10.1029/2012GL053695.

Yvon-Lewis, S. A., Saltzman, E. S. and Montzka, S. A. (2009), Recent trends in atmospheric methyl bromide: analysis of post-Montreal Protocol variability, *Atmospheric Chemistry and Physics*, 9(16), 5963-5974.

Zhang, F., Zhou, L.X., Yao, B., Vollmer, M.K., Grealley, B.R., Simmonds, P.G., Frode Stordal, S.R., Maione, M., Xu, L. and Zhang, X. C. (2010a), Analysis of 3-year observations of CFC-11, CFC-12 and CFC-113 from a semi-rural site in China, *Atmospheric Environment*, 44(35), 4454-4462.

Zhang, F., Zhou, L. X., Yao, B., Vollmer, M. K., Grealley, B. R., Simmonds, P. G., Reimann, S., Stordal, F., Maione, M., Xu, L., and Zhang, X. C. (2010c), Analysis of 3-year observations of CFC-11, CFC-12 and CFC-113 from a semi-rural site in China, *Atmospheric Environment*, 44(35), 4454-4462.

Zhang, Y. H., Hu, M., Zhong, L. J., Wiedensohler, A., Liu, S. C., Andreae, M. O., Wang, W., and Fan, S. J. (2008), Regional Integrated Experiments on Air Quality over Pearl River Delta 2004 (PRIDE-PRD2004): Overview, *Atmospheric Environment*, 42(25), 6157-6173.

Zhang Y. L., Guo H., Wang X. M., Simpson I. J., Barletta, B., Blake, D. R., Meinardi, S., Rowland, F. S., Cheng, H. R., Saunders, S. M. and Lam, S. H. M. (2010b), Emission patterns and spatiotemporal variations of halocarbons in the Pearl River Delta region, southern China, *Journal of Geophysical Research: Atmospheres*, 115, 16.

Zhang Y. L., Wang, X. M., Simpson, I. J., Barletta, B., Blake, D. R., Meinardi, S., Louie, P. K. K., Zhao, X. Y., Shao, M., Zhong, L. J., Wang, B. G. and Wu, D. (2014), Ambient CFCs and HCFC-22 observed concurrently at 84 sites in the Pearl River Delta region during the 2008-2009 grid studies, *Journal of Geophysical Research: Atmospheres*, 119(12), 7699-7717.

Zhong, Z. M., Zheng, J. Y., Zhu, M. N., Huang, Z. J., Zhang, Z. W., Jia, G. L., Wang, X. L., Bian, Y. L., Wang Y. L. and Li, N. (2018). Recent developments of

956 anthropogenic air pollutant emission inventories in Guangdong province, China.
957 *Science of the Total Environment*, 627, 1080-1092.
958
959
960
961