# RELIABILITY ASSESSMENT OF SLOPES CONSIDERING SAMPLING INFLUENCE AND SPATIAL VARIABILITY BY SOBOL' SENSITIVITY INDEX

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#### 5 ABSTRACT

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This paper presents an extended formulation of the Sobol' sensitivity index for geotech-6 nical reliability assessments involving spatially variable soil properties. It incorporates the 7 subsurface spatial correlation structure with the response surface method, which is then 8 assimilated into the context of the Sobol' index approach. A Sobol' index map can be gen-9 erated for the entire subsurface domain, identifying the sensitive zones which also represent 10 the optimal sampling locations. In addition, the approach allows the derivation of the mean 11 and variance of system response conditional to any sample value, without the needs to con-12 duct separate conditional random field simulations. This is adopted for the assessment of 13 reliability of slopes, where design charts are established for cases where a single sample is 14 obtained within slopes of  $c_u$  or  $c - \phi$  soils, with various conditions of geometries and spatial 15 variability. The approach can also be applied to multiple sampling points, thereby facilitat-16 ing a feedback mechanism where the planning of geotechnical investigation and evaluation 17 of performance uncertainty can be considered in a holistic manner. 18

<sup>19</sup> Keywords: Slope stability, probabilistic analyses, conditional random field, sampling loca-

tion, Sobol' sensitivity index

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#### 21 INTRODUCTION

Reliability of geotechnical systems depends heavily on the uncertainty of soil and rock 22 Recently, reliability approaches such as first-order second-moment (FOSM) properties. 23 method (Duncan 2000) and the Hasofer-Lind approach (first-order reliability method) (Ha-24 sofer and Lind 1974) have gained popularity in the field of geotechnical engineering. However, 25 many of these methods do not explicitly consider the spatial variations in material properties, 26 which is an important source of uncertainty in geotechnical engineering. For example, the 27 significance of such has been discussed by Christian and Baecher (2011), who also suggested 28 that traditional assumptions of perfect (or no) spatial correlations in geotechnical proper-29 ties do not always yield conservative estimates. The characterization of spatial variability 30 in geotechnical properties has been discussed in Phoon and Kulhawy (1999), Baecher and 31 Christian (2003), Liu et al. (2017b), Liu and Leung (2017), etc. A spatially variable soil 32 profile can be modeled by the random field theory (Vanmarcke 1984), which considers the 33 geotechnical property as a set of spatially correlated random variables, and its applications 34 include shallow foundations (Fenton and Griffiths 2003; Al-Bittar and Soubra 2014), slope 35 stability (Cho 2010; Jiang et al. 2015), soil liquefaction (Popescu et al. 2005), etc. Mean-36 while, the geotechnical profession has long recognized the importance of obtaining samples 37 or conducting *in situ* tests at 'representative' locations, and the spatial uncertainty can be 38 reduced considerably when such information is incorporated into the probabilistic analyses 39 (Lloret-Cabot et al. 2012; Li et al. 2016a). Accordingly, conditional random field modelling 40 is a numerical approach that can assess the influence of known sample values at designated 41 locations. In particular, some recent studies quantified the reductions in performance un-42 certainty considering the available sample values and their locations. These include the 43 probabilistic assessments of footing settlements by Lo and Leung (2017a), and those for 44 slope stability by Li et al. (2016b) and Liu et al. (2017a). 45

<sup>46</sup> Despite being a powerful approach to utilize existing field data, conditional random
 <sup>47</sup> field modelling suffers from some key limitations from the risk management perspective.

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Essentially, the approach requires sample values and locations to be specified prior to the analyses, and each probabilistic assessment involves hundreds to thousands of realizations, based on one set of such samples. This exercise becomes impractical if the objectives are to explore multiple options for future sampling locations, and to evaluate how the additional, currently unknown, sample values may affect the performance uncertainty.

A complete assessment of the performance uncertainty should be based on all possible 53 random field realizations. Also, it is desirable to develop an efficient strategy to obtain the 54 optimal sampling pattern tailored to the specific application, allowing assessments of the 55 influence of future sample values to the overall system reliability. To this end, this paper 56 extends the formulation of Sobol' sensitivity index for geotechnical reliability assessments 57 involving spatial variability in soil properties. It facilitates a feedback mechanism so that 58 the planning of geotechnical investigation and evaluation of performance uncertainty can 59 be considered in a holistic manner. While the approach is potentially applicable to various 60 geotechnical problems, this study focuses on slope stability and presents design charts to 61 illustrate the relationships between sample values and associated uncertainty in factor of 62 safety. 63

# RELIABILITY ASSESSMENT THROUGH SENSITIVITY ANALYSIS BY SOBOL' INDEX

Sobol' sensitivity index is a probabilistic tool developed by Sobol' (2001) to assess the 66 influence of each input parameter in a physical model. In general, each model parameter 67 can be regarded as a random variable. A Sobol' index can then be associated with each 68 parameter, and used to quantify its contribution to the variance of the model response. This 69 concept has been applied to a number of geotechnical applications, including investigations 70 on the parameters affecting the consolidation process (Houmadi et al. 2012) and ground 71 settlements induced by tunneling (Miro et al. 2014). Others have combined the Sobol' 72 index approach with the response surface (or surrogate modeling) method to enhance the 73 computational efficiency, and studied the parameters associated with footing displacements 74

<sup>75</sup> and stability of pressurized tunnels (Mollon et al. 2011).

These previous studies mainly focused on the influence of parameters in 'deterministic' 76 models, where spatial variability of the concerned parameters was not considered. In fact, the 77 Sobol' index was originally developed for independent input variables, and cannot be directly 78 applied to spatially correlated fields of soil properties. Therefore, this study extends the 79 original theory for applications to correlated random variables, which represent the correlated 80 properties at various locations. The Sobol' index can be evaluated for each location, resulting 81 in a 'Sobol' index map' that reveals the relative importance of soil properties at all location to 82 the system response. The maximum index value corresponds to the most influential location, 83 which is equivalent to the optimal sampling location. 84

The proposed Sobol' index approach enables site-specific geotechnical reliability assess-85 ments to be performed with considerations of spatial variability and information from soil 86 samples. Details of the assessment strategy are shown in a flowchart in Fig. S1, while the key 87 elements are briefly described herein. Based on existing soil samples or experience at the site, 88 the mean, variance and spatial correlation features of the parameters can be estimated. A 89 set of unconditional (or conditional if prior samples are available) random field analyses can 90 be performed for system response q. The Sobol' indices throughout the subsurface domain 91 are then evaluated to (1) identify optimal locations for additional samples; and (2) derive the 92 relationships between system reliability (mean and variance of g) and potential values of the 93 future samples (referred to as 'sensitivity functions' herein). These can provide guidance on 94 the adequacy of the design, and whether more samples should be planned to further reduce 95 the uncertainty or risk levels. In this process, only one set of random field simulations is 96 required unless major revisions of the original design are involved. In the following sections, 97 the formulation of Sobol' index approach for spatially variable soils are presented, together 98 with its application to single and multiple sample points in slope stability assessments. 99

#### FORMULATION FOR GEOTECHNICAL RELIABILITY ASSESSMENT

#### <sup>101</sup> Single sample point

In general, the properties of a spatially variable soil can be represented by a trend component and the residuals, or deviations from the trend:

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$$\boldsymbol{z} = \boldsymbol{\mu} + \boldsymbol{\varepsilon} \tag{1}$$

<sup>105</sup> where  $\boldsymbol{z} = \{z_1, z_2, \dots, z_d\}^{\mathrm{T}}$  is a vector of soil properties at d different locations.  $\boldsymbol{\varepsilon}$  represents <sup>106</sup> the random deviation of the property at each location, and has a constant variance  $\sigma_z^2$ .  $\boldsymbol{\mu}$ <sup>107</sup> represents the trend, or expected value, at each location. In this study,  $\boldsymbol{\mu}$  is assumed to be <sup>108</sup> a constant vector, i.e., the random field is statistically stationary.  $\boldsymbol{z}$  represents lognormal <sup>109</sup> random fields of soil properties in this study, i.e.  $z_i = \ln(z_i')$ , where  $z_i'$  is the original value of <sup>110</sup> the concerned property. The  $z_i$  components are correlated with each other, and the spatial <sup>111</sup> correlation structure is described herein by a squared exponential function:

$$R_{ij} = \exp\left[-\left(\frac{x_i - x_j}{\theta_{\ln,x}}\right)^2 - \left(\frac{y_i - y_j}{\theta_{\ln,y}}\right)^2\right]$$
(2)

where  $R_{ij}$  is the spatial correlation (autocorrelation) between  $z_i$  and  $z_j$ , which also represents the ij-th element of the correlation matrix **R**. The spatial coordinates of  $z_i$  and  $z_j$  are  $(x_i, y_i)$ and  $(x_j, y_j)$ , respectively, while  $\theta_{\ln,x}$  and  $\theta_{\ln,y}$  are the autocorrelation distances in x and ydirections. The system response of a geotechnical model can be represented as g(z), which is also a random variable since the model input z is random. Alternatively, to facilitate the formulation of the proposed approach, g can be treated as a function of the standardized residual instead, i.e. g(e), where  $e = \varepsilon/\sigma_z$ .

The Sobol' index can be adopted to assess the influence of property at each soil location,  $e_i$ , to the variance of g. By the law of total variance, the first order Sobol' index can be defined as:

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$$S(e_i) = \frac{\operatorname{Var}_{e_i} \left[ \operatorname{E}_{\boldsymbol{e}_{-i}}(g|e_i) \right]}{\operatorname{Var}(g)} = 1 - \frac{\operatorname{E}_{e_i} \left[ \operatorname{Var}_{\boldsymbol{e}_{-i}}(g|e_i) \right]}{\operatorname{Var}(g)}$$
(3)

In the context of spatially variable subsurface domain,  $E_{e_i}[Var_{e_i}(g|e_i)]$  is the expected 124 value of system variance if sample value  $(e_i)$  is known for location i. Such variance arises from 125 the fact that properties are unknown at other locations (-i).  $\operatorname{Var}_{e_i}[\operatorname{E}_{e_{-i}}(g|e_i)]$  represents the 126 variance in expected system response, due to all possible values of  $e_i$ . These two terms are 127 denoted as  $E[Var(g|e_i)]$  and  $Var[E(g|e_i)]$  hereafter. In other words, the Sobol' index can 128 be interpreted as the variance reduction in model response, when a soil sample is obtained 129 at the *i*-th location. To facilitate the calculation of the Sobol' index, the response surface 130 method may be used to reduce the number of simulations of the geotechnical model. In 131 this study, the second order Polynomial Chaos Expansion (PCE) is adopted (Ghanem and 132 Spanos 1991), which is given by: 133

$$g = a_0 + \sum_{j=1}^{M} a_j \xi_j + \sum_{j_1=1}^{M} \sum_{j_2=j_1}^{M} a_{j_1 j_2} (\xi_{j_1} \xi_{j_2} - \delta_{j_1 j_2})$$
(4)

where  $\xi_1, \ldots, \xi_M$  are independent standard normal variables, which may be grouped as a vec-135 tor  $\pmb{\xi}$  with M principal components, representing a realization of soil profile.  $\delta_{j_1 j_2}$  is the Kro-136 necker delta. The unknown PCE coefficients  $(a_0, a_j, a_{j_1j_2})$  are determined by a non-intrusive 137 approach, which involves generating and simulating N realizations of the geotechnical model 138 for q, followed by a linear regression analysis. The prediction accuracy of the coefficients 139 is measured by an indicator known as  $Q^2$ . Meanwhile, an adaptive algorithm proposed by 140 Blatman and Sudret (2010) is adopted, whereby only the PCE coefficients that can increase 141 the  $Q^2$  are kept in the expansion (Eq. (4)). Once the PCE coefficients are determined, the 142 expectation and variance of g can be evaluated as below: 143

$$E(g) = a_0 \tag{5a}$$

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$$\operatorname{Var}(g) = \sum_{j=1}^{M} a_j^2 + \sum_{j_1=1}^{M} \sum_{j_2=j_1}^{M} a_{j_1j_2}^2 (1+\delta_{j_1j_2})$$
(5b)

To utilize the PCE for evaluation of  $S(e_i)$ , it should be noted that although the vectors 48  $\boldsymbol{\xi}$  and  $\boldsymbol{e}$  are not equivalent, they are related by the following transformation:

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$$\boldsymbol{e} = \mathbf{H}\boldsymbol{\Lambda}^{\frac{1}{2}}\boldsymbol{\xi} = \mathbf{C}\boldsymbol{\xi} \tag{6}$$

where **H** and **A** are obtained from spectral decomposition of the correlation matrix ( $\mathbf{R} = \mathbf{H}\mathbf{A}\mathbf{H}^{\mathrm{T}}$ ), with  $\mathbf{H} = [\mathbf{h}_1, \mathbf{h}_2, \dots, \mathbf{h}_d]$  being a matrix containing d orthonormal eigenvectors, and **A** is a diagonal matrix with d descending eigenvalues  $(\lambda_1, \lambda_2, \dots, \lambda_d)$ . The residual at the *i*-th location is therefore  $e_i = \mathbf{C}_i \boldsymbol{\xi}$ .  $\mathbf{C}_i$  is the *i*-th row of **C**, and is abbreviated as  $\boldsymbol{c}$  in the subsequent formulation. Taking a conditional expectation on both sides of Eq. (4), the mean response of g conditioned on sample value  $e_i$  can be represented by:

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$$E(g|e_i) = a_0 + a_j \sum_{j=1}^M E(\xi_j|e_i) + a_{j_1j_2} \sum_{j_1=1}^M \sum_{j_2=j_1}^M \left[E(\xi_{j_1}\xi_{j_2}|e_i) - \delta_{j_1j_2}\right]$$
(7)

Since  $\boldsymbol{\xi}$  is multivariate normal and e is standard normal, the conditional distribution of  $\boldsymbol{\xi}$ given  $e_i$  is multivariate normal with mean of  $\boldsymbol{c}^{\mathrm{T}}e_i$  and covariance of  $\mathbf{I} - \boldsymbol{c}^{\mathrm{T}}\boldsymbol{c}$  (I is the identity matrix). Therefore,

$$\mathbf{E}(\xi_j|e_i) = c_j e_i \tag{8a}$$

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$$E(\xi_{j_1}\xi_{j_2}|e_i) = \delta_{j_1j_2} - c_{j_1}c_{j_2} + c_{j_1}c_{j_2}e_i^2$$
(8b)

Substituting (8) into (7),  $E(g|e_i)$  can be rewritten as:

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$$E(g|e_i) = r_0 + r_1 e_i + r_2 e_i^2$$

W

here 
$$r_1 = \sum_{j=1}^{M} a_j c_j$$
  
 $r_2 = \sum_{j_1=1}^{M} \sum_{j_2=j_1}^{M} a_{j_1 j_2} c_{j_1} c_{j_2}$ 

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$$r_0 = a_0 - r_2 \tag{9}$$

Using Eq. (9), the mean response can be evaluated efficiently for any sample value, without resorting to additional conditional random field simulations. To evaluate  $\operatorname{Var}[\mathrm{E}(g|e_i)]$ , it should be noted that  $e_i$  is a standard normal variable, and  $\operatorname{Var}(e_i) = 1$ ;  $\operatorname{Var}(e_i^2) = 2$ ;  $\operatorname{Cov}(e_i, e_i^2) = 0$ . Therefore,

$$\operatorname{Var}\left[\mathrm{E}(g|e_i)\right] = r_1^2 + 2r_2^2 \tag{10}$$

With Eqs. (9) and (10), the Sobol' index  $S(e_i)$  can be computed by Eq. (3). A Sobol' index map can be generated once  $S(e_i)$  are calculated for all locations, and can be used to identify the most influential location, where the  $S(e_i)$  value is maximum.

While Eq. (9) represents the conditional mean of g for normal random variable e, the soil property in this study is modeled as lognomal random variable. In other words, e in Eq. (9) represents the residual in log-transformed space:  $e = (\ln z' - \mu_{\ln z'})/\sigma_{\ln z'}$ , but the residual in original space is  $e' = (z' - \mu_{z'})/\sigma_{z'}$ . Conversion between the original residual e'and transformed residual e is given by:

$$e = \frac{\ln(e' \operatorname{CV}(z') + 1)}{\sqrt{\ln(1 + \operatorname{CV}(z')^2)}} + \frac{1}{2}\sqrt{\ln(1 + \operatorname{CV}(z')^2)}$$
(11)

where  $CV(z') = \sigma_{z'}/\mu_{z'}$ . Substituting (11) into (9), the conditional mean in terms of e' is:

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$$E(g|e'_i) = s_0 + s_1 \ln [e'_i CV(z') + 1] + s_2 (\ln [e'_i CV(z') + 1])^2$$

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where 
$$s_0 = r_0 + \frac{r_1}{2}\sqrt{\ln(1 + CV(z')^2)} + \frac{r_2}{4}\ln(1 + CV(z')^2)$$
  
 $s_1 = \frac{r_1}{1 + CV(z')^2} + r_2$ 

s

$$r_{1} = \sqrt{\ln(1 + CV(z')^{2})} + r_{2}$$

$$r_{2} = \frac{r_{2}}{\ln(1 + CV(z')^{2})}$$
(12)

From the risk assessment perspective, a key concern is the remaining uncertainty in system response after a sample value has been obtained. This is represented by the conditional variance,  $E[Var(g|e_i)]$ , or the conditional standard deviation (SD). There are two methods to estimate the conditional SD. The first method is to estimate directly from the Sobol' index:

<sup>191</sup> 
$$\operatorname{E}\left[\operatorname{SD}(g|e_i)\right] \approx \sqrt{\operatorname{E}\left[\operatorname{Var}(g|e_i)\right]} = \sqrt{\left[1 - S(e_i)\right]\operatorname{Var}(g)} = \sqrt{1 - S(e_i)}\operatorname{SD}(g)$$
(13)

This method has been employed by Lo and Leung (2017b) to investigate the uncertainty 192 reduction in footing displacement, and is accurate if the conditional SD is insensitive to the 193 sample value. If the sample value heavily influences the conditional SD, an explicit function 194 for conditional SD is required. Therefore, a conditional variance function  $\operatorname{Var}(g|e_i) = V(\boldsymbol{\theta}, e_i)$ 195 is proposed, with  $\boldsymbol{\theta}$  being a vector of parameters describing the variance function. During 196 the construction of PCE, N combinations of model response and sample values  $(g_i, e_i)$  are 197 obtained.  $\theta$  can then be obtained by maximizing the following log-likelihood function (Da-198 vidian and Carroll 1987): 199

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$$L(\boldsymbol{\theta}) = -\sum_{i=1}^{N} \ln \left[ V(\boldsymbol{\theta}, e_i) \right] - \sum_{i=1}^{N} \frac{\left[ g_i - \mathcal{E}(g|e_i) \right]^2}{V(\boldsymbol{\theta}, e_i)}$$
(14)

where  $E(g|e_i)$  is evaluated from Eq. (9). The maximization is performed in this study by an evolutionary searching algorithm known as Differential Evolution (Storn and Price 1997), which is not prone to converging at local maxima. Once  $\operatorname{Var}(g|e_i)$  is obtained, it can be converted to  $\operatorname{Var}(g|e'_i)$  through Eq. (11).

#### <sup>205</sup> Multiple sample points

Geotechnical investigation usually involves multiple samples being retrieved from bore-206 hole(s). It is therefore beneficial to extend the sensitivity analysis framework, for determi-207 nation of multiple sample locations and the corresponding reduction in system uncertainty. 208 This can be achieved by evaluating the *n*-th order Sobol' sensitivity index S(e). The defi-209 nition of the *n*-th order Sobol' index is similar to the 1st order index in Eq. (3), with the 210 standardized residual  $e_i$  being replaced by a residual vector,  $\boldsymbol{e} = \{e_1, e_2, \dots, e_n\}^{\mathrm{T}}$ , represent-211 ing samples from n locations. S(e) can be interpreted as the averaged variance reduction of 212 the system response, when soil samples are available from n locations. 213

From a second order PCE, S(e) can be formulated by first considering the joint distribution of vectors  $\boldsymbol{\xi}$  and  $\boldsymbol{e}$ :

$$\begin{pmatrix} \xi_{1} \\ \xi_{2} \\ \vdots \\ \xi_{M} \\ e_{1} \\ e_{2} \\ \vdots \\ e_{n} \end{pmatrix} \sim N \begin{bmatrix} \begin{pmatrix} 0 \\ 0 \\ \vdots \\ 0 \\ 0 \\ 0 \\ \vdots \\ 0 \end{pmatrix}, \begin{pmatrix} 1 & 0 & \cdots & 0 & C_{11} & C_{21} & \cdots & C_{n1} \\ 0 & 1 & \cdots & 0 & C_{12} & C_{22} & \cdots & C_{n2} \\ \vdots & \vdots & \ddots & \vdots & \vdots & \vdots & \ddots & \vdots \\ 0 & 0 & \cdots & 1 & C_{1M} & C_{2M} & \cdots & C_{nM} \\ C_{11} & C_{12} & \cdots & C_{1M} & 1 & R_{12} & \cdots & R_{1n} \\ C_{21} & C_{22} & \cdots & C_{2M} & R_{21} & 1 & \cdots & R_{2n} \\ \vdots & \vdots & \ddots & \vdots & \vdots & \vdots & \ddots & \vdots \\ C_{n1} & C_{n2} & \cdots & C_{nM} & R_{n1} & R_{n2} & \cdots & 1 \end{pmatrix} \end{bmatrix} = N \begin{bmatrix} \begin{pmatrix} \mathbf{0} \\ \mathbf{0} \end{pmatrix}, \begin{pmatrix} \mathbf{I} & \mathbf{C}_{s}^{\mathrm{T}} \\ \mathbf{0} \end{pmatrix} \end{bmatrix}$$

$$(15)$$

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with covariance between  $\boldsymbol{\xi}$  and  $\boldsymbol{e}$  being  $\operatorname{Cov}(\boldsymbol{e}_i, \boldsymbol{\xi}_j) = \operatorname{Cov}(\mathbf{C}_i, \boldsymbol{\xi}, \boldsymbol{\xi}_j) = C_{ij}$ .  $\mathbf{C}_s$  and  $\mathbf{R}_s$  in Eq. (15) is a subset of the full  $\mathbf{C}$  (Eq. (6)) and  $\mathbf{R}$  matrices, and they are not identical. By multivariate normal theory, the conditional distribution of  $\boldsymbol{\xi}$  given  $\boldsymbol{e}$  is also multivariate normal with mean of  $\mathbf{C}_s^{\mathrm{T}} \mathbf{R}_s^{-1} \boldsymbol{e}$  and covariance of  $\mathbf{I} - \mathbf{C}_s^{\mathrm{T}} \mathbf{R}_s^{-1} \mathbf{C}_s$ . Hereafter, they are expressed as  $\mathbf{K} \boldsymbol{e}$  and  $\mathbf{I} - \mathbf{Q}$ , respectively. Based on this conditional distribution, the following conditional expectations can be obtained:

$$\mathbf{E}(\xi_j|\boldsymbol{e}) = \mathbf{K}_j \boldsymbol{e} \tag{16a}$$

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$$E(\xi_{j_1}\xi_{j_2}|\boldsymbol{e}) = \delta_{j_1j_2} - \mathbf{Q}_{j_1j_2} + (\mathbf{K}_{j_1}\boldsymbol{e})(\mathbf{K}_{j_2}\boldsymbol{e})$$
(16b)

Substituting Eq. (16) into the conditional mean of a 2nd order PCE, the conditional 226 mean of g can be written as: 227

E(g|e) = 
$$r_0 + \sum_{i=1}^n r_i e_i + \sum_{i_1=1}^n \sum_{i_2=i_1}^n r_{i_1 i_2} e_{i_1} e_{i_2}$$
  
where  $r_0 = a_0 - \sum_{j_1=1}^M \sum_{j_2=j_1}^M a_{j_1 j_2} \mathbf{Q}_{j_1 j_2}$ 

$$r_{i} = \left[\sum_{j=1}^{M} a_{j}\mathbf{K}_{j}\right]_{i}$$

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$$r_{i_1i_2} = \begin{cases} \mathbf{P}_{ii} & \text{if } i_1 = i_2 = i \\ \mathbf{P}_{i_1i_2} + \mathbf{P}_{i_2i_1} & \text{if } i_1 \neq i_2 \end{cases}$$

$$\mathbf{P} = \left[\sum_{j_1=1}^{M} \sum_{j_2=j_1}^{M} a_{j_1j_2} \mathbf{K}_{j_1}^{\mathrm{T}} \mathbf{K}_{j_2}\right]$$
(17)

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It is also possible to represent the conditional mean in terms of residuals in original space (e'), 233 and derive the s coefficients similar to Eq. (12). However, the corresponding formulation is 234 complex, and it may be easier to directly apply Eq. (17) with log-transformed e. Meanwhile, 235  $\operatorname{Var}[\operatorname{E}(q|\boldsymbol{e})]$  is given by: 236

Var 
$$[E(g|\boldsymbol{e})] = \sum_{i=1}^{n} \sum_{k=1}^{n} r_i r_k Cov(e_i, e_k) + \sum_{i_1=1}^{n} \sum_{i_2=i_1}^{n} \sum_{k_1=1}^{n} \sum_{k_2=k_1}^{n} r_{i_1i_2} r_{k_1k_2} Cov(e_{i_1}e_{i_2}, e_{k_1}e_{k_2})$$
 (18)

To evaluate the components in Eq. (18), it should be noted that the components in e are 238 spatially correlated. The following covariance can be obtained by the theory of characteristic 239 functions (with details in the Appendix): 240

$$Cov(e_i, e_k) = R_{ik} \tag{19a}$$

$$\operatorname{Cov}(e_i, e_{k_1} e_{k_2}) = 0 \tag{19b}$$

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$$Cov(e_{i_1}e_{i_2}, e_{k_1}e_{k_2}) = R_{i_1k_1}R_{i_2k_2} + R_{i_1k_2}R_{i_2k_1}$$
(19c)

and  $\operatorname{Var}[\operatorname{E}(g|\boldsymbol{e})]$  can then be used to calculate the *n*-order Sobol' index.

To obtain the maximum index S(e), the spatial configuration of the sample vector e has 247 to be optimized. The n locations for vector e may be searched simultaneously using various 248 optimization algorithms. Alternatively, the multiple sample locations can be determined 249 sequentially. For example, starting with i = 1, the Sobol' index is evaluated across the 250 domain to identify the location that corresponds to the maximum value of  $S(e_1)$ . Once 251 this location is decided for the first sample, the Sobol' index is evaluated again with i = 2, 252 based on the selected location of sample 1, to identify the second sample location leading to 253 the maximum value of  $S(e_1, e_2)$ . This stepwise procedure is then repeated until the target 254 number of samples is reached. It is also possible to impose various constraints when searching 255 for the optimal sample locations. For example, when multiple samples are retrieved from a 256 single borehole, they need to share the same horizontal coordinates. 257

The proposed Sobol' index approach is essentially a post-processing technique for ran-258 dom field analyses, performed in this study through the PCE. The benefits of the proposed 259 approach are multifold. Contrary to previous conditional random field modeling techniques, 260 the Sobol' index is a global sensitivity index, which means it encompasses all possible values 261 of an uncertain input parameter. In addition, by constructing a Sobol' index map, all po-262 tential sample locations are assessed simultaneously for the formulation of optimal sampling 263 strategy. There are no needs for pre-determined sample values or sampling patterns, and it 264 is not necessary to perform separate conditional random field simulations to investigate mul-265 tiple scenarios, thereby reducing the computational demands substantially. In the following 266 sections, the Sobol' index approach will be validated through comparisons with conditional 267 random field simulations, and then applied to reliability assessments involving slope stability. 268

#### APPLICATIONS IN RELIABILITY ASSESSMENTS OF SLOPES

#### 270 Comparisons with conditional random field simulations

The proposed Sobol' index approach can be validated by comparisons with conditional 271 random field simulations, through which the advantages of the approach are also illustrated. 272 Essentially, it allows efficient evaluation of the mean and standard deviation of factor of 273 safety (FS) from any sample value e', which can then be used to obtain the reliability index 274 or failure probability of the slope. This section considers a 35° slope with a height of 5 m, 275 simulated in FLAC with a model boundary 15 m below the top of the slope. The soil is 276 idealized as Tresca material, with mean undrained shear strength  $(c_u)$  of 36.3 kPa, which 277 would correspond to a 'deterministic' FS of 2.0, if the soil profile was assumed to be uniform. 278 Instead, the  $c_u$  profile is simulated as lognormally distributed in the model, with coefficient 279 of variation of 0.4, horizontal autocorrelation distances of  $\theta_{\ln,x} = 14.3$  m and  $\theta_{\ln,y} = 2.5$  m. 280 Throughout this study, the element size in the FLAC model is about 0.5 m (vertical)  $\times$ 281 1.0 m (horizontal), which are always smaller than half of the autocorrelation distances in the 282 corresponding directions. The influence of element size on FS values of slopes was studied by 283 Dawson et al. (1999), who showed that for a 10-m slope with slope angles between 15°-45°, 284 the difference in FS values obtained by a fine mesh  $(60 \times 60)$  and a coarse mesh  $(20 \times 20)$ 285 is smaller than 4%. As part of the verification process in this study, a separate analysis 286 was conducted with 4 times the number of elements (each 0.25 m  $\times$  0.5 m in size), and the 287 subsequent changes to the Sobol' indices, conditional mean and SD curves are mostly under 288 5%. 289

Based on these spatial variability features, unconditional random field simulations are first performed without specifying any sample locations in the domain. 1,000 realizations of  $c_u$  profiles are generated by Latin Hypercube Sampling with Dependence (LHSD), which is an extension of LHS and aims to introduce stratification while maintaining the spatial correlation of random variables. Details of LHSD and its incorporation into Cholesky decomposition are described in Packham and Schmidt (2010) and Lo and Leung (2017a), who also showed that when LHSD is coupled with PCE, 1,000 realizations are sufficient to obtain
 robust estimates of the PCE coefficients and hence probability density of the FS for slopes
 with similar features of spatially variable soils.

Using this approach, the unconditional mean and standard deviation of FS are found to 299 be E(FS) = 1.697 and SD(FS) = 0.308 in this case. Through decomposition of the R matrix 300 and Eqs. (9) and (10), the Sobol' index map is then derived, and is shown in Fig. 1(a). The 301 optimal sample location is found at the depth of 6.25 m from the slope top, with horizontal 302 separation of 5.8 m from the slope toe, where the Sobol' index is at the maximum of 0.397. 303 This means that on average (considering all possible sample values), an approximate SD 304 reduction of 22.3% can be achieved if a sample is obtained at this location. The concentric 305 shape of the Sobol' index variation is related to the ratio between  $\theta_{\ln,x}$  and  $\theta_{\ln,y}$ . Since  $\theta_{\ln,x}$  is 306 more than 5.5 times larger than  $\theta_{\ln,y}$  in this example, the corresponding contour is elongated 307 in the x-direction. In case of a layered soil profile with  $\theta_{\ln,x} = \infty$ , the Sobol' index contour 308 will also display a 'layered' feature as all locations at the same depth are equally important 309 because of the highly correlated properties. 310

The conditional mean equation, considering standardized residuals in original space (E(FS|e')), is then calculated from the PCE through Eq. (12) and plotted in Fig. 1(b), where the coefficients are found to be  $(s_0, s_1, s_2) = (1.76, 0.47, -0.175)$  for this case. As mentioned earlier, the conditional variance equation can be obtained by maximizing the log-likelihood function (Eq. (3)) involving  $V(\theta, e)$ . In this study, the generalized logistic function (Richards 1959) is adopted for  $V(\theta, e)$  since it has a flexible functional form with an asymmetric S-shape:

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$$Var(FS|e) = V(\theta, e) = l + \frac{u - l}{\{1 + t \exp[-b(e - m)]\}^{1/t}}$$
(20)

The shape of the function is controlled by the parameters  $\boldsymbol{\theta} = \{l, u, t, b, m\}$ , where l is the lower bound, u is the upper bound, t is the asymmetry parameter, b is the growth rate, and *m* is the point of inflection. The maximization is performed by Differential Evolution, with the following optimized parameters:  $\{l, u, t, b, m\} = \{0.013, 0.12, 1.6, 1, 0.56\}$  in this case. Through this 'sensitivity function', the relationship between uncertainty in FS and the potential sample values can be developed. For example, the conditional variance Var(FS|*e*) is converted to the conditional SD curve in terms of original residual, i.e. SD(FS|e'), and plotted in Fig. 1(c). The function allows rapid determination of the system reliability once the sample value becomes available.

These results by the Sobol' index approach can be validated through conducting sepa-328 rate conditional random field simulations, where sample values are assigned at the optimal 329 sampling point mentioned earlier. Eight individual sets of analyses are performed, with dif-330 ferent sample values assigned at that location: e' = (-1.5, -1, -0.5, 0, 0.5, 1, 1.5, 2), which 331 corresponds to about 95% of the possible cases considering the lognormal distribution of 332  $c_u$ . For each set of analysis, 1,000 realizations of the conditional random field are simulated 333 and assessed using the LHSD-PCE approach. The corresponding conditional mean and SD 334 results are also plotted in Figs. 1(b) and (c), showing close agreements with the relationships 335 developed by the proposed Sobol' index approach, and therefore validating its accuracy. It 336 is important to note that an additional 8,000 FLAC analyses are required by conditional 337 random field modeling, whereas the proposed Sobol' index approach only requires 1,000 un-338 conditional random field simulations. The associated reductions in computational demands 339 are substantial. 340

In fact, the influence of sampling at any (non-optimal) location can be evaluated by the proposed approach, by adopting another c vector corresponding to the intended sample location i. To illustrate such effects, a sensitivity analysis is performed with another sample location indicated in Fig. 1(a). The prediction intervals, taken as  $E(FS|e')\pm SD(FS|e')$ , are shown in Fig. 1(d) for both sample locations. The prediction interval arising from the optimal location is steeper and narrower than that of the non-optimal location, which again demonstrates that FS is more sensitive to the  $c_u$  value at the optimal location.

#### Design charts for slopes in $c_u$ soils

Based on the proposed Sobol' index approach, a series of design charts can be developed for efficient assessments of slope reliability according to the available sample value of shear strength parameters. For example, a long embankment may be designed based on mean shear strength parameters across the site. Making use of the charts in this study, the failure probabilities of individual sections can be easily evaluated when a soil sample is retrieved at any of those locations.

Table 1 presents the various slope geometries and subsurface conditions considered in 355 this study. For slopes in  $c_u$  soils, the FLAC model geometries are similar to those described 356 earlier, with slope height H = 5 m, and model boundary 15 m below the top of the slope. 357 Four different slope angles are considered:  $\beta = 20^{\circ}$ ,  $30^{\circ}$ ,  $35^{\circ}$  and  $40^{\circ}$ , and the width of 358 the slope is therefore  $W = H/\tan\beta$ .  $c_u$  is modeled as a lognormal random field, with the 359 mean values  $(\mu_{cu})$  chosen such that deterministic analyses (assuming uniform soil profiles) 360 would lead to FS of 1.0, 1.5 or 2.0. Such 'deterministic FS' provides a useful indicator of the 361 degree of strength mobilization, which can be easily applied by most practitioners in a typical 362 design process. Various patterns of spatial variability for  $c_u$  are considered in developing the 363 design charts. The coefficient of variation (i.e.  $CV_{cu}$ ) is assigned to be 0.15 or 0.4, while the 364 autocorrelation distance  $\theta_{\ln,x} = 0.5W$ , 2W or  $\infty$ , and  $\theta_{\ln,y} = 0.25H$  or 0.5H. Considering 365 these variations, 36 series of probabilistic analyses are performed for each slope angle. 366

For each series of analyses, a PCE is constructed using 1,000  $c_u$  profile realizations. It 367 should be noted that for all analyses in this study, the value of  $Q^2$ , which measures accuracy 368 of PCE coefficients, exceeds 0.92. With the coefficients determined, the optimal sampling 369 locations are then obtained using the Sobol' index approach. It was found that the sampling 370 locations are insensitive to the mean values of the property, and therefore cases with different 371  $\mu_{cu}$  values are averaged. According to Fig. 2(a), the optimal sampling depths D are all below 372 the slope, i.e., D/H > 1, which is expected since deep-seated failures are more common for 373 slopes in  $c_u$  soils. In general, D becomes shallower as  $\beta$  increases or with higher variability 374

in  $c_u$ , since the possibilities of shallow slope failures increase with potential weak spots at shallow depths. Fig. 2(a) also shows the horizontal distance between the optimal sampling location and the slope toe (L), normalized by the width of slope (W). In most cases, it would be reasonable to adopt L = 0.5W.

Based on these optimal sample locations, the sensitivity functions (conditional mean 379 E(FS|e') and conditional standard deviation SD(FS|e') relationships) can be evaluated by 380 the procedures described earlier. Fig. 3 shows that under the same deterministic FS (of 381 1.5) and spatial variability conditions (i.e.  $CV_{cu}$ ,  $\theta_{\ln,x}$ ,  $\theta_{\ln,y}$ ), the sensitivity functions are 382 relatively insensitive to the slope angle (from  $20^{\circ}$  to  $40^{\circ}$ ), with a narrow range around the 383 average value. Therefore, the obtained functions are averaged across various slope angles in 384 the subsequent figures. Also, various sets of analyses are performed with the model scale 385 doubled, as illustrated in Table 1. The results are also shown in Fig. 3 and indicated that 386 the sensitivity functions are scale-invariant. In other words, regardless of the size of the 387 slope, the functions will be identical for the same  $\theta_{\ln,x}/W$  and  $\theta_{\ln,y}/H$  ratios. Meanwhile, 388 the intention of Fig. 3 is not to imply that autocorrelation distances will be scaled up or 389 down with the slope geometries at a particular site. In fact, for site-specific applications, it 390 is necessary to evaluate the  $\theta_{\ln,x}/W$  and  $\theta_{\ln,y}/H$  ratios based on the site conditions, and use 391 the corresponding sensitivity functions pertaining to those conditions. 392

For example, Fig. 4 represents the design charts with sensitivity functions for different 393 values of deterministic FS, and demonstrates the influence of spatial correlation features 394 on slope reliability. In particular, the E(FS|e') function becomes steeper with larger values 395 of  $\theta_{\ln,x}$ ,  $\theta_{\ln,y}$  and/or  $CV_{cu}$ , which means the sample has a larger conditioning effect to FS 396 in these cases. In Figs. 4(a) and (b), the E(FS|e') functions for  $\theta_{\ln,x} = 2W$  are omitted 397 for brevity, but the curves lie approximately midway between those of  $\theta_{\ln,x} = 0.5W$  and 398  $\theta_{\ln,x} = \infty$  for positive e' values, and are close to the latter for negative e' values. The 399 corresponding SD(FS|e') curves are also shown in Figs. 4(c) to (h), and they are found to 400 increase with  $CV_{cu}$ ,  $\theta_{\ln,x}$  and  $\theta_{\ln,y}$ . Using these relationships, the reliability of a slope can 401

<sup>402</sup> be quickly assessed for any sampled value of  $c_u$  represented as standardized residual e'. For <sup>403</sup> other values of deterministic FS, it is reasonable to interpolate the results as a monotonic <sup>404</sup> trend is observed among the various curves. As discussed earlier, these relationships are <sup>405</sup> independent of the problem scale and may also be applied to other slope dimensions.

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#### Design charts for slopes in $c - \phi$ soils

Similar procedures are applied to slopes in  $c-\phi$  soils to establish the corresponding design 407 charts. In this section, a slope height of H = 10 m is adopted, and the model boundary is 408 15 m below the top of the slope. The H to model boundary ratio is smaller than that for 409  $c_u$  soils, since slopes in 'sandy' soils often fail with relatively shallow slip surfaces, and this 410 phenomenon will be discussed here in terms of sensitivity analyses. This section will focus 411 on the variability of friction angle  $\phi$ , while c is assumed to be a constant of 5 kPa. As shown 412 in Table 1,  $\phi$  is modeled as a lognormal random field, with mean values ( $\mu_{\phi}$ ) designated to 413 achieve deterministic FS of 1.0, 1.5 or 2.0. The coefficient of variation of  $\phi$  (i.e.  $CV_{\phi}$ ) is 414 assigned to be 0.05 or 0.1, while other variations are similar to those adopted for  $c_u$  slopes. 415 'Dry' slopes are first investigated and the influence of water table will be discussed later. 416

Again, the optimal sampling locations are found to be insensitive to the  $\mu_{\phi}$  values, and 417 therefore cases with different  $\mu_{\phi}$  are averaged. The normalized sampling positions are shown 418 in Figs. 2(c) and (d), where the optimal sampling depths are all within the slope, with 419 D/H < 1. In other words, stability of  $c - \phi$  slopes are mostly influenced by shear strength 420 close to the slope face, which is another manifestation of the shallow slip surfaces usually 421 observed in these slopes. Similar to the  $c_u$  slopes, a larger slope angle or  $CV_{\phi}$  value would 422 lead to a shallower optimal sampling depth, and D/H ranges from around 0.75 to 0.95 for 423 all the studied cases. Meanwhile, L is approximately 0.4W for most cases. 424

Fig. 5 shows the conditional mean and SD curves for  $c - \phi$  slopes with sample obtained from the optimal location. Similar to the case of  $c_u$  slopes, additional analyses are performed with a double model size, and the results are found to be invariant to the model scale (Fig. 3). In addition, the curves are found to be insensitive to the horizontal autocorrelation distance of  $\phi$  once  $\theta_{\ln,x}$  exceeds 2*W*. This may be attributed to the relatively shallow slip surfaces, which tend to intercept several horizontal soil layers, making  $\theta_{\ln,y}$  generally more influential when  $\theta_{\ln,x}$  extends beyond the width of the slope. On the other hand, Figs. 5(a) and (b) shows that the E(FS|*e'*) curves becomes steeper with increasing values of  $\theta_{\ln,y}$ , which means the sample significance increases as the soil becomes more uniform.

Interestingly, the SD(FS|e') curves in Figs. 5(c) to (h) are shifted downwards with in-434 creasing  $\theta_{\ln,y}$ , which is an opposite trend compared to  $c_u$  slopes. In fact, the conditional SD 435 curves depend both on the value of  $\theta_{\ln,y}$ , and the conditioning effects of the sample. For a 436 probabilistic assessment without any sample value, the 'unconditional' SD of slope perfor-437 mance will always increase with  $\theta_{\ln,y}$ , since there are less random effects that 'average out' 438 the soil variability, leading to more uncertain performance. This trend is, however, counter-439 acted by the knowledge of sample value at a designated location, in which case a larger  $\theta_{\ln,y}$ 440 means a greater conditioning effect of the sample, as a thicker layer becomes associated with 441 that value. The conditioning effects appear to be less dominant for  $c_u$  slopes that involve 442 deep-seated failure mass, but are more influential for  $c - \phi$  slopes with shallow slip surfaces. 443 Similar influence of conditioning to the trend reversal of SD is also observed in probabilistic 444 assessments of footing performance presented by Lo and Leung (2017a). 445

The influence of water table to the reliability of  $c - \phi$  slopes is also investigated in 446 this study, where the water level is modeled such that 3/4 of the slope is submerged, as 447 shown in the inset of Fig. 6(a). The scenario with deterministic FS=1.5 and slope angle of 448 30° is selected for comparisons of E(FS|e') and SD(FS|e'), as shown in Fig. 6. It should be 449 noted that the mean friction angles are different for dry and submerged slopes in order to 450 achieve the same deterministic FS, with  $\mu_{\phi} = 32.5^{\circ}$  for the former case and  $\mu_{\phi} = 42.5^{\circ}$  for 451 the latter. Fig. 6 shows that the E(FS|e') relationships are steeper for submerged slope, 452 resulting in higher E(FS|e') for positive e' and lower E(FS|e') for negative e'. Meanwhile, 453 the SD(FS|e') relationships are shifted upwards when water table is present. However, the 454 differences in both mean and standard deviation of FS are not substantial comparing the 455

dry and submerged cases. In general, the optimal sampling depths for the submerged slope is slightly deeper (by about 0.05H) compared to the dry slope case.

#### 458

#### 58 Example applications of the design charts

To illustrate the application of the presented charts, this section considers a scenario 459 where a 5 m high cut slope of  $30^{\circ}$  is to be constructed. Based on previous geotechnical 460 investigation and/or knowledge of the site, suppose the mean  $c_u$  of the clayey deposits is 461 estimated to be 27 kPa, with  $CV_{cu} = 0.4$ . The standard deviation of  $c_u$  (i.e.,  $\sigma_z$ ) is therefore 462 10.8 kPa, while  $\theta_{\ln,x} = 17.3$  m (= 2W) and  $\theta_{\ln,y} = 2.5$  m. With this mean  $c_u$  value, the 463 deterministic FS is 1.5. To refine the estimates of failure probability for a particular section 464 of the slope, a soil sample is to be obtained. Making use of Fig. 2, the optimal location of 465 the sample would be about 7.5 m deep from the top of slope, and 4.3 m away from the toe 466 of the future slope profile. Suppose a clay sample is then retrieved from this location with 467  $c_u$  determined to be 22.7 kPa. This would correspond to e' = -0.4 (the sampled value is 468  $-0.4\sigma_z$  away from the mean). From the design charts in Fig. 4, taking deterministic FS= 1.5, 469  $\theta_{\ln,y} = 0.5H$  and with e' = -0.4, the conditional mean and SD of FS become 1.27 and 0.18, 470 respectively. Assuming the FS has a normal distribution, the slope has a failure probability 471 of 0.067. 472

Consider another possible scenario at the same site, where prior samples or knowledge 473 of the soil properties are limited, and the mean  $c_u$  at the site cannot be determined with 474 confidence. Nonetheless, the optimal sample location is still the same since it does not 475 depend on the mean property value. Suppose a sample is retrieved and the  $c_u$  of that sample 476 is determined to be 22.7 kPa. Without other information, this is taken as the mean value 477 and the deterministic FS is found to be 1.26. Making use of the design charts in Fig. 4, the 478 conditional mean and SD of FS may be further approximated, assuming e' = 0 and with 479 estimates of  $CV_{cu}$  and  $\theta_{\ln,y}$  according to published literature or the engineers' judgement. It 480 is also possible to test how the failure probability changes with different CV or  $\theta$  assumptions, 481 and these estimates may provide useful information to support the engineering decisions. 482

#### 483 Applications for multiple sample points

This section illustrates the application of Sobol' index approach considering multiple sample points, making use of the sequential algorithm described earlier. The slope is in  $c_u$ soils with a slope height of 5 m and slope angle of 35°. The *FLAC* model boundary is 15 m below the top of the slope.  $c_u$  is simulated as a lognormal random field with a mean value 27.3 kPa, corresponding to a deterministic FS of 1.5.  $CV_{cu} = 0.4$ , while  $\theta_{\ln,x} = 2W = 14.3$  m and  $\theta_{\ln,y} = 0.25$  H = 1.25 m. Based on these parameters, the unconditional mean and standard deviation of FS are E(FS)=1.28 and SD(FS)=0.18.

Considering the scenario where six soil samples are to be obtained from two boreholes 491 in the slope, the optimal sample locations and resulting conditional FS distribution can be 492 evaluated by the proposed approach. The first sample is determined by searching throughout 493 the entire slope profile, and the corresponding lateral coordinates become those of the first 494 borehole. If the optimal location of subsequent samples falls within a close horizontal distance 495 (taken as 3 m in this case) from the first borehole, then a constrained search is performed 496 such that the sample would share the same lateral coordinates as this borehole. Otherwise 497 this sample indicates the location of the second borehole. 498

Fig. 7(a) shows the location of the six samples thus determined, while Table 2 summarizes 499 their coordinates and the corresponding Sobol' index values. The cumulative reductions in 500 SD(FS) are also shown in percentages as the number of samples increases, approximated by 501  $100[1 - \sqrt{1 - S(e)}](\%)$  as indicated in Eq. (13). The vertical spacing between the samples 502 is 2 to 2.5 m, which is about 2 times  $\theta_{\ln,y}$ . The horizontal spacing between the two boreholes 503 is 7.3 m, which is about 0.5 times  $\theta_{\ln,x}$ , which is consistent with the recommendation by Li 504 et al. (2016b). Once the sample locations are determined, sensitivity analysis is conducted 505 to obtain the conditional mean equation E(FS|e), according to Eq. (17). The coefficients of 506 the conditional mean equation are given in Table 3. 507

The improvements in response prediction through the proposed sampling strategy are further illustrated through 1,000 realizations of  $c_u$  profiles, simulated by the LHSD technique

described by Lo and Leung (2017a). The profiles are then analyzed by FLAC to obtain the 510 FS values, denoted by  $FS_1, FS_2, \ldots, FS_{1000}$ . If an unconditional random field assessment was 511 performed, the prediction errors can be represented by  $FS_i - E(FS)$  (*i* from 1 to 1,000). On 512 the other hand, the conditional predictions can be performed using the sampling strategy 513 presented in Table 2, and 1,000 sample combinations can be realized as  $e_1, e_2, \ldots, e_{1000}$ . The 514 associated prediction errors are therefore  $FS_i - E(FS|\boldsymbol{e}_i)$ , where the coefficients of  $E(FS|\boldsymbol{e}_i)$ 515 are given in Table 3 and no separate conditional random field simulations are required. 516 Fig. 7(b) shows the density plots of the unconditional errors and the conditional errors, 517 with their standard deviations being 0.184 and 0.102, respectively, and hence the percentage 518 reduction in SD is about 44.3%. This SD reduction is identical to the value shown in Table 2, 519 which is calculated using a different approach, i.e., directly through the Sobol' index. In other 520 words, this numerical testing also verifies the accuracy of the proposed approach for multiple 521 sample points. 522

#### 523 EXAMPLE APPLICATION OF DESIGN CHARTS TO CASE STUDY

The case study of the James Bay hydroelectric project in Quebec was described in detail 524 by Christian et al. (1994), and is revisited in this study to illustrate the application of the 525 proposed design charts and their differences with the FOSM method. The project involved 526 construction of embankments on soft sensitive clays, with the first stage consisting of two 527 berms with a total height (H) of 12 m. The two berms are separated from each other in 528 the horizontal direction and the total width (W) is approximately 90 m (Fig. 8a). Field 529 vane shear tests had been conducted in 35 boreholes across the site to characterize  $c_u$  of the 530 marine clay and lacustrine clay below the embankment, and the details are also shown in 531 Fig. 8b. The spatial variability features were investigated by DeGroot and Baecher (1993) 532 and Christian et al. (1994) (Fig. 8c), with the autocorrelation distances in horizontal and 533 vertical directions estimated to be 37.3 m and 1.1 m from their results, assuming similar 534 spatial correlation features for marine and lacustrine clays. 535

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Some approximations are inevitable when applying the design charts for the case study.

Based on the slope geometries and spatial correlation features, the charts for  $\theta_{\ln,y} = 0.25H$ 537 and  $\theta_{\ln,x} = 0.5W$  are adopted. The deterministic FS was estimated to be 1.453 ( $\approx 1.5$ ) 538 by Christian et al. (1994), with the critical failure plane intersecting the lower lacustrine 539 clay that has  $CV_{cu} = 0.272$ . Therefore, when applying the design charts, the results from 540  $CV_{cu} = 0.15$  and  $CV_{cu} = 0.4$  are averaged. Considering the total H and W of the two 541 berms, the average slope angle would be 8°, and the corresponding optimal (single) sample 542 location is at L = 0.5W and D = 2.4H by extrapolating Fig. 2. The sampled value at that 543 depth is very close to the mean  $c_u$ , and therefore e' = 0. 544

Based on these approximations and by averaging the results from Figs. 4(a) and 4(b), 545 the conditional mean FS is estimated to be 1.382, and the conditional SD of FS is 0.083 from 546 Figs. 4(g) and 4(h). This SD value is lower than the value estimated with FOSM by Christian 547 et al. (1994), which was 0.205 considering only the contributions from spatial variations of 548  $c_u$  (total SD was 0.257, which included factors such as variations in the fill and crust layers). 549 The discrepancy can be attributed to two main reasons. Firstly, the FOSM method assumes 550 the soil to be homogeneous  $(\theta_{\ln,y} = \theta_{\ln,x} = \infty)$  when the partial derivative  $\partial FS/\partial c_u$  was 551 evaluated, which leads to an increase in SD estimates. For example, if the design charts are 552 applied to the scenario of  $\theta_{\ln,x} = \infty$  and  $\theta_{\ln,y} = 0.5H$ , with other conditions being the same 553 (i.e., Figs. 4(c) and 4(d)), the estimated SD will become 0.154, and will increase further as 554  $\theta_{\ln,y}$  approaches the assumptions of FOSM. In addition, the proposed approach and design 555 charts explicitly considered the influence of the soil sample, which also reduces the SD in FS 556 estimates. It should be noted, however, that the James Bay project involves multiple samples 557 which cannot be accounted for only by using the design charts. Yet, the estimates through 558 the charts can be useful indicators for practitioners. If all samples need to be considered in 559 a more rigorous manner, the approach presented in Eqs. (15) to (19) can be adopted. 560

#### 561 CONCLUSION

This paper extends the Sobol' index approach for the simulation of spatially variable soil properties, and explores its applications to the reliability assessment of slopes. The approach can be used to identify the optimal locations of samples which will bring maximum reduction
to the uncertainty in system performance. It also allows the derivation of sensitivity functions
for site-specific risk assessment of system performance, which is computationally efficient
since the influence of different sample values can be evaluated without performing additional
conditional random field simulations.

Design charts are presented for the case of a single sample in  $c_u$  slopes and  $c - \phi$  slopes. 569 These charts enable efficient assessment of the failure probability of slopes, represented as 570 the conditional mean and standard deviations under various conditions of slope geometries 571 and spatial variability. It should be noted, however, that probabilistic methods should not 572 be taken as replacement of the understanding of local geology, which can be helpful in 573 identifying potential anomalously weak layers that could be one of the main causes of slope 574 failures. Also, the charts are developed based on the assumption of statistical stationarity in 575 the shear strength parameters of the soil, with constant mean values of  $c_u$  or  $\phi$ . They should 576 be applied with proper engineering judgement, especially when the geotechnical conditions 577 deviate significantly from these assumptions. In those cases, it is also possible to conduct 578 detailed risk assessments through the proposed Sobol' index approach. 579

The proposed approach is extended to consider multiple sampling points through evalua-580 tion of multi-ordered Sobol' index. This is illustrated through an example where the optimal 581 locations of six samples along two boreholes are determined. A significant reduction in the 582 performance uncertainty can be achieved, and is verified by the much smaller prediction 583 errors in the FS of slope, compared to the case when no samples are available. The proposed 584 Sobol' index approach is shown to be a useful post-processing tool on the existing random 585 field simulation results, and is capable of revealing high risk zones related to the specific 586 geotechnical applications. This can be integrated into a risk assessment framework to assist 587 the decision-making process associated with the uncertainty of system performance arising 588 from spatial variability of geotechnical properties. 589

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#### SUPPLEMENTAL DATA 590

Descriptions on the details of risk-based design process utilizing the proposed Sobol' index 591 approach, together with Fig. S1, are available online in the ASCE Library (www.ascelibrary.org). 592

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#### **APPENDIX** 598

According to Anderson (1984), if  $\boldsymbol{X} = \{X_1, \dots, X_p\}^T$  is a multivariate normal distribu-599 tion, i.e.  $\boldsymbol{X} \sim N(\boldsymbol{\mu}, \boldsymbol{\Sigma})$ , then the characteristic function  $\phi$  is given by: 600

$$\phi(\boldsymbol{t}) = \mathbf{E}\left[\exp(i\boldsymbol{t}^{T}\boldsymbol{X})\right] = \exp\left(i\boldsymbol{t}^{T}\boldsymbol{\mu} - \frac{1}{2}\boldsymbol{t}^{T}\boldsymbol{\Sigma}\boldsymbol{t}\right)$$
(21)

The *n*-th non-central moment of **X** can be obtained by differentiating  $\phi$ : 602

$$\mathbf{E}\left[X_1\dots X_n\right] = \frac{1}{i^n} \frac{\partial \phi}{\partial t_1\dots \partial t_n} \bigg|_{t=0}$$
(22)

The first four non-central moments are hence derived: 604

1

 $\operatorname{E}[X_1] = \mu_1$ 605

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603

$$\mathbb{E}[\Lambda_1]$$

607

$$E[X_{1}X_{2}] = \frac{1}{i^{2}} \frac{\partial \phi}{\partial t_{1} \partial t_{2}} \Big|_{t=0} = \sigma_{12} + \mu_{1}\mu_{2}$$

$$E[X_{1}X_{2}X_{3}] = \frac{1}{i^{3}} \frac{\partial \phi}{\partial t_{1} \partial t_{2} \partial t_{3}} \Big|_{t=0} = \sigma_{12}\mu_{3} + \sigma_{13}\mu_{2} + \sigma_{23}\mu_{1} + \mu_{1}\mu_{2}\mu_{3}$$

$$E[X_{1}X_{2}X_{3}X_{4}] = \frac{1}{i^{4}} \frac{\partial \phi}{\partial t_{1} \partial t_{2} \partial t_{3} \partial t_{4}} \Big|_{t=0} = \sigma_{12}\sigma_{34} + \sigma_{13}\sigma_{24} + \sigma_{14}\sigma_{23}$$
(23)

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 $+\sigma_{12}\mu_{3}\mu_{4}+\sigma_{13}\mu_{2}\mu_{4}+\sigma_{14}\mu_{2}\mu_{3}+\sigma_{23}\mu_{1}\mu_{4}+\sigma_{24}\mu_{1}\mu_{3}+\sigma_{34}\mu_{1}\mu_{2}$ 

<sup>609</sup> which lead to these covariances:

610  

$$Cov[X_{1}, X_{2}X_{3}] = E[X_{1}X_{2}X_{3}] - E[X_{1}]E[X_{2}X_{3}] = \sigma_{12}\mu_{3} + \sigma_{13}\mu_{2}$$

$$Cov[X_{1}X_{2}, X_{3}X_{4}] = E[X_{1}X_{2}X_{3}X_{4}] - E[X_{1}X_{2}]E[X_{3}X_{4}]$$

$$= \sigma_{13}\sigma_{24} + \sigma_{14}\sigma_{23} + \sigma_{13}\mu_{2}\mu_{4} + \sigma_{14}\mu_{2}\mu_{3} + \sigma_{23}\mu_{1}\mu_{4} + \sigma_{24}\mu_{1}\mu_{3}$$
(24)

Since  $\boldsymbol{e} \sim N(\boldsymbol{0}, \mathbf{R})$ ,

614

$$Cov(e_i, e_{k_1}e_{k_2}) = 0$$
$$Cov(e_{i_1}e_{i_2}, e_{k_1}e_{k_2}) = R_{i_1k_1}R_{i_2k_2} + R_{i_1k_2}R_{i_2k_1}$$
(25)

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**TABLE 1.** Slope geometries and shear strength variability parameters adopted in Sobol'index analyses

Parameters		$c_u$ soils	$c_u$ soils	$c - \phi$ soils	$c - \phi$ soils
			(double scale $)$		(double scale $)$
Slope angle, $\beta$ (for all		20°, 30°	°, 35°, 40°		
Soil unit weight (for a		20 k	${ m N/m^3}$		
Slope Height, $H$	$5 \mathrm{m}$	10 m	$10 \mathrm{~m}$	20 m	
Deterministic FS	1,1.5,2	1.5	1,1.5,2	1.5	
Variability of shear	$\mathrm{CV}_{cu}$	0.15,0.4	0.15,0.4	-	-
strength parameters	$\mathrm{CV}_{\phi}$	-	-	0.05,  0.1	0.05,  0.1
	$\theta_{\ln,x}$ (for all cases)		$0.5W,2W,\infty$		
	$\theta_{\ln,y}$ (for all cases)		0.25H	H, 0.5H	

n	Borehole	Depth from	$S(e_1, \ldots, e_n)$	Cumulative $\%$
(No. of samples)	number	slope top (m)		reduction in SD
1	1	6.25	0.24	12.8
2	1	8.75	0.37	20.6
3	2	4.2	0.47	27.2
4	1	10.75	0.56	33.7
5	1	13.25	0.63	39.2
6	2	2.2	0.69	44.3

TABLE 2. Six sample locations from two boreholes for a slope in  $c_u$  soils

Order of coefficient	r	Value
0th order	$r_0$	1.318
1st order	$r_1$	0.080
	$r_2$	0.057
	$r_3$	0.050
	$r_4$	0.047
	$r_5$	0.038
	$r_6$	0.039
2nd order	$r_{11}$	-0.0078
	$r_{12}$	0.0012
	$r_{13}$	-0.0097
	$r_{14}$	0.0105
	$r_{15}$	0.0098
	$r_{16}$	-0.0095
	$r_{22}$	-0.0103
	$r_{23}$	0.0046
	$r_{24}$	-0.0019
	$r_{25}$	0.0118
	$r_{26}$	-0.0028
	$r_{33}$	-0.0026
	$r_{34}$	0.0078
	$r_{35}$	0.0137
	$r_{36}$	-0.0073
	$r_{44}$	-0.0096
	$r_{45}$	0.0016
	$r_{46}$	0.0056
	$r_{55}$	-0.0090
	$r_{56}$	0.0085
	$r_{66}$	0.0023

## TABLE 3. Conditional mean coefficients for multiple samples in $\ensuremath{c_u}$ slope

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FIG. 1. Comparison between Sobol' index approach and individual conditional random field simulations



FIG. 2. Optimal sampling locations normalized by slope height (D/H) and width (L/W) for slopes in (a)  $c_u$  soils and (b)  $c - \phi$  soils. Data points represent the results for four slope angles and the lines are trendlines



FIG. 3. Conditional mean and standard deviation functions with deterministic FS=1.5, slope angle varying from 20° to 40° and different model scales



FIG 4. Conditional mean and standard deviation functions for slopes in  $c_u$  soils, with gray areas bounded by  $\theta_{n,y} = 0.25H$  and 0.5H.  $\theta_{n,y} = 0.5H$  results in higher SD and mean FS for positive e' (*x*-axes are identical for sub-figures)



FIG 5. Conditional mean and standard deviation functions for slopes in  $c - \phi$  soils, with gray areas bounded by  $\theta_{n,y} = 0.25H$  and 0.5H.  $\theta_{n,y} = 0.5H$  results in lower SD and higher mean FS for positive e'(*x*-axes are identical for sub-figures)



FIG. 6. Influence of water table on sensitivity functions



FIG. 7. (a) Locations of six samples along two boreholes; (b) Prediction errors in FS comparing 1,000 realizations with unconditional and conditional mean estimates



FIG. 8. (a) Cross section of embankment in James Bay project; (b) Variations of  $c_u$  in foundation clay; (c) Autocovariance functions for  $c_u$  in vertical and horizontal directions (adapted from Christian et al. 1994 and DeGroot and Baecher 1993)