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2 3 4	An automatic method for screening clouds and cloud shadows in optical satellite image time series in cloudy regions
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Abstract: Clouds and cloud shadows block land surface information in optical satellite images. 30 Accurate detection of clouds and cloud shadows can help exclude these contaminated pixels in 31 further applications. Existing cloud screening methods are challenged by cloudy regions where 32 most of satellite images are contaminated by clouds. To solve this problem for landscapes where 33 the typical frequency of cloud-free observations of a pixel is too small to use existing methods to 34 35 mask clouds and shadows, this study presents a new Automatic Time-Series Analysis (ATSA) method to screen clouds and cloud shadows in multi-temporal optical images. ATSA has five main 36 steps: (1) calculate cloud and shadow indices to highlight cloud and cloud shadow information; (2) 37 obtain initial cloud mask by unsupervised classifiers; (3) refine initial cloud mask by analyzing 38 time series of a cloud index; (4) predict the potential shadow mask using geometric relationships; 39 and (5) refine the potential shadow mask by analyzing time series of a shadow index. Compared 40 with existing methods, ATSA needs fewer predefined parameters, does not require a thermal 41 42 infrared band, and is more suitable for areas with persistent clouds. The performance of ATSA was 43 tested with Landsat-8 OLI images, Landsat-4 MSS images, and Sentinel-2 images in three sites. The results were compared with a popular method, Function of Mask (Fmask), which has been 44 adopted by USGS to produce Landsat cloud masks. These tests show that ATSA and Fmask can 45 46 get comparable cloud and shadow masks in some of the tested images. However, ATSA can consistently obtain high accuracy in all images, while Fmask has large omission or commission 47 errors in some images. The quantitative accuracy was assessed using manual cloud masks of 15 48 49 images. The average cloud producer's accuracy of these 15 images is as high as 0.946 and the average shadow producer's accuracy reaches 0.884. Given that it can be applied to old satellite 50 51 sensors and it is capable for cloudy regions, ATSA is a valuable supplement to the existing cloud 52 screening methods.

53 Key words: Cloud detection; Cloud shadow; Mask; Optical satellite images; Time series

### 55 **1. Introduction**

Optical satellite images with bands ranging from visible to shortwave infrared are widely 56 used for mapping land cover and land use, monitoring ecosystems, and estimating land surface 57 parameters (Hansen and Loveland, 2012; Zhu and Liu, 2015, 2014). Unfortunately, optical satellite 58 images are easily contaminated by clouds and cloud shadows. This contamination obscures land 59 60 surface features and alters the reflectance of ground objects, reducing the availability of optical images for applications (Fisher, 2013; Zhu and Woodcock, 2014). Masking clouds and cloud 61 shadows is often the first and a necessary step of image preprocessing in most optical remote 62 sensing applications. Although manual digitization can obtain accurate cloud and shadow masks, 63 it requires a lot of time and effort. Therefore, an automatic method for screening clouds and 64 shadows is needed, especially when processing large numbers of images. 65

Automatic detection of clouds and cloud shadows is challenging (Zhu and Woodcock, 2014). 66 First, different types of clouds have different spectral signatures and are easily confused with some 67 68 cloud-free bright objects on the land surface, especially in images with limited spectral bands, such as Landsat Multispectral Scanner (MSS) images. The spectral signals of clouds are usually 69 determined by cloud height, optical thickness, particle size, etc. (Platnick et al., 2003). As a result, 70 71 cloud brightness ranges widely in visible and near infrared bands, and some clouds are easily confused with bright land surfaces, such as concrete surfaces, sand or snow. Second, blurry cloud 72 edges and thin clouds partially obscure land surfaces, making their signal a mixture of cloud and 73 74 land surface elements and making them difficult to separate from clear observations (Cahalan et al., 2001). Another challenge comes from cloud shadows. They are easily confused with dark land 75 76 surfaces, such as moist soil, water bodies and topographic shadow (Fisher, 2013).

77 Despite the above challenges, several methods have been developed to automatically screen

clouds and cloud shadows in optical images. These methods use one or more of the following rules 78 based on cloud and cloud shadow properties: 1) clouds are generally brighter than ground objects, 79 so they have high reflectance in visible, near and shortwave infrared bands; 2) clouds are generally 80 colder than most ground objects, so they have lower brightness in thermal infrared bands; 3) 81 shadows are generally darker than surrounding land surfaces, so they have lower reflectance in 82 83 visible, near and shortwave infrared bands; 4) shadows are paired with clouds, so cloud location and solar angles can help locate cloud shadows; and 5) in a sequence of images, pixels affected by 84 clouds and shadows have larger temporal variations than clear observations in the time series. In 85 general, existing methods for masking clouds and cloud shadows can be divided into two 86 categories: single-image methods (Choi and Bindschadler, 2004; Fisher, 2013; Helmer et al., 2012; 87 Huang et al., 2010; Hughes and Hayes, 2014; Irish et al., 2006; Li et al., 2015, 2017; Luo et al., 88 2008; Martinuzzi et al., 2007; Roy et al., 2010; Scaramuzza et al., 2012; Wilson and Oreopoulos, 89 2013; Zhu and Woodcock, 2012) and multi-temporal or bi-temporal methods (Goodwin et al., 2013; 90 91 Hagolle et al., 2010; Jin et al., 2013; Wang et al., 1999; Zhu and Woodcock, 2014).

Most existing single-image methods use either predefined thresholds or adaptive thresholds 92 to screen clouds in individual images. For example, Luo et al. (2008) identify clouds in MODIS 93 94 images if pixel reflectance satisfies these predefined thresholds: (B1 > 0.18 or B3 > 0.20) and B6 > 0.200.16 and Maximum (B1, B3) > B6 $\times$ 0.67, where B1, B3, and B6 are reflectance of MODIS bands 95 1 (blue), 3 (red), and 6 (shortwave infrared), respectively. This MODIS cloud screening method 96 97 was further adopted for Landsat-8 images (Wilson and Oreopoulos, 2013). Huang et al. (2010) use adaptive thresholds defined in the reflectance-temperature space to mask clouds in Landsat TM 98 99 and ETM+ images. These adaptive thresholds are defined by the mean and standard deviation of 100 pixel values of individual bands in the whole image. The Automatic cloud cover assessment

(ACCA) algorithm consists of twenty-six filters and rules applied to Landsat bands to detect clouds 101 (Irish et al., 2006). ACCA was used to produce web-enable Landsat data (WELD), a consistent, 102 long-term, and large-area data record (Roy et al., 2010). The multi-feature combined (MFC) 103 method uses thresholds in spectral, geometric and texture features to detect clouds in GaoFen-1 104 imagery (Li et al., 2017). Zhu and Woodcock (2012) proposed a method called function of mask 105 106 (Fmask) for detecting clouds in Landsat TM and ETM+ images. Fmask uses all Landsat image bands and several band indices, such as the normalized difference vegetation index (NDVI) and 107 108 the normalized difference snow index (NDSI). It employs more than 20 predefined and adaptive 109 thresholds to mask clouds. Besides the above methods using predefined or adaptive thresholds, machine-learning algorithms have been employed to model the complex relationships between 110 image features and clouds using a training dataset. Then, the trained model is used to screen clouds 111 in other images. These machine learning algorithms include decision trees (Scaramuzza et al., 112 2012), neural networks (Hughes and Hayes, 2014) and support vector machines (Li et al., 2015). 113 114 Of several tested cloud and shadow masking algorithms that use only a single image, Fmask is globally the most accurate one that requires a thermal band (Foga et al., 2017). Of methods not 115 requiring a thermal band, a version of ACCA (Irish et al., 2006) that uses a simulated thermal band 116 117 is better overall, but it is not as accurate as Fmask with the thermal band (Foga et al., 2017). Recently, Fmask was further improved for mountainous areas through integrating Digital 118 119 Elevation Models (DEMs) into the detecting process (Qiu et al., 2017).

In these single-image methods, shadow detection is often subsequent to cloud detection. In general, the possible shadow locations can be calculated from the geometric relationship between sun, sensor, and clouds. The calculation requires cloud heights, which can be estimated with brightness temperature derived from thermal infrared bands, because temperature declines with elevation (Qiu et al., 2017; Zhu and Woodcock, 2012). Some methods also use the fact that cloud
shadows are dark to confirm whether the possible shadow location estimated from geometry is real
cloud shadow, including Fmask (Zhu and Woodcock, 2012) and MFC (Li et al., 2017). In Fmask,
predefined thresholds in the near infrared (NIR) band are used to produce a potential shadow mask,
which is further compared to the possible shadow locations. If there is a high similarity between
potential shadow masks and possible shadow locations, the shadow pixels are finally confirmed
(Zhu and Woodcock, 2012).

For multi-temporal methods, temporal information in the images acquired at different times 131 is used to detect clouds and shadows. Wang et al. (1999) used the brightness difference between a 132 target image and a reference cloud-free image to detect clouds. Lyapustin et al. (2008) developed 133 an algorithm, abbreviated as MAIAC CM, to detect clouds in time series of MODIS images. The 134 general idea of MAIAC CM is to use the low covariance between reference cloud-free image 135 blocks and cloudy image blocks as a criterion to identify clouds in the time series. Hagolle et al. 136 137 (2010) computes differences in the blue band between a target image and a cloud-free reference image. It then flags cloud pixels if variations are larger than a threshold. Goodwin et al. (2013) 138 uses filters to smooth the time-series and then identify clouds and shadows based on reflectance 139 140 differences between each point in the time series and the smoothed time series. Zhu and Woodcock (2014) propose a new algorithm called multiTemporal mask (Tmask) to improve Fmask. Tmask 141 142 fits a time series model of each pixel using remaining clear pixels based on an initial cloud mask 143 from Fmask. Then, it compares model estimates with observations in the time series to detect cloud and shadow pixels which are omitted in the initial screening by Fmask. In general, these multi-144 145 temporal methods are better at detecting clouds and cloud shadows than single-image methods. 146 The temporal information is a valuable complement to the spectral information for differentiating cloud, cloud shadow and clear observations over land surfaces (Goodwin et al., 2013; Zhu and
Woodcock, 2014).

However, these multi-temporal methods still face challenges in areas with persistent cloud 149 cover, such as tropical and subtropical regions (Ju and Roy, 2008). First, in these areas cloud-free 150 observations may be the exception rather than the rule, making it difficult to know whether the fit 151 152 of a time series represents clear or cloudy conditions, which limits the application of existing timeseries methods (Foga et al., 2017). Example limitations include the requirement by the MAIAC 153 CM method of a cloud free image as a reference image (Lyapustin et al., 2008), and the 154 recommendation for Tmask of 15 cloud-free observations for estimating the time series model 155 (Zhu and Woodcock, 2014). Second, most existing methods were designed for images of a specific 156 sensor, so they lack flexibility. For example, Fmask and Tmask were designed for Landsat TM, 157 ETM+, and OLI images, so they cannot be directly applied to the old Landsat MSS data with 158 limited bands. Third, most of the current methods use predefined fixed thresholds to detect clouds 159 160 and shadows in an entire scene. For instance, in Tmask, a pixel with observed green band reflectance of 0.04 higher than the time series model estimation will be identified as cloud. 161 Considering the complex situation of clouds and shadows and the diversity of objects on land 162 163 surfaces and in coastal areas, these fixed thresholds may not always obtain satisfactory results.

To overcome the above limitations of existing methods in cloudy regions, the objective of this study is to develop a new automatic method for accurately screening clouds and cloud shadows in multi-temporal optical images in places with persistent clouds. Our scope of inference is landscapes where are so cloudy that the typical frequency of cloud-free observations of a pixel is too small to use existing methods to mask clouds and shadows with image time series. The new method should have the following strengths: 1) it needs fewer predefined parameters; 2) it is suitable for areas with persistent clouds; and 3) it needs a minimal number of bands. Automatic
Time-Series Analyses (ATSA) method was developed in this study and tested in three pilot sites
using Landsat OLI and MSS images, and Sentinel-2 images. Its performance was compared with
Fmask, a widely recognized method.

#### 174 **2. Test Sites and Data**

175 2.1. A cloudy urban site

Urban landscapes bring more challenges to automatic screening of clouds and shadows than 176 177 other landscapes. The bright built-up area often leads to large commission errors in cloud detection. 178 To test the effectiveness of the proposed method in such challenging cases, we selected Hong Kong, a cloudy subtropical dense city with complex and mixed land-cover types (Fig. 1). This site has an 179 area of 1, 620 km<sup>2</sup> (1200×1500 Landsat pixels), and the central coordinates are 22.367 N and 180 181 114.123° E. It is covered by the Landsat scene of Worldwide Reference System 2 (WRS-2) Path 182 122 Row 44. All 23 available Landsat-8 OLI level-1 images in 2015 were downloaded from USGS 183 Earth Explorer. These images were then converted to Top of Atmosphere (TOA) reflectance with 184 the scaling coefficients in the metadata file. The corresponding Fmask cloud masks of these images 185 were also downloaded from USGS Earth Explorer. Based on Fmask cloud masks, only two images 186 are clear, while the other images have total cloud and shadow coverage ranging from 5.5% to 97%. Sixteen of them have total cloud and shadow coverage larger than 60%, indicating Landsat 187 imagery in this site is seriously contaminated by clouds (Table 1). 188

Fig. 1. True color composition of a Landsat-8 image of 2015, DOY003 in Hong Kong

Table 1. Summary of cloud and shadow coverage of Landsat-8 images from the year 2015 over

204 the Hong Kong site using Fmask product. Only two images have no clouds.

DOY	Cloud coverage %	Shadow coverage %	Total cloud and shadow coverage%
3	0.0	0.0	0.0
19	0.0	0.0	0.0
35	97	0.0	97
51	65	3.8	69
67	90	0.65	90
83	87	3.1	90
99	89	0.0	89
115	82	0.06	82
131	32	5.4	38
147	94	0.25	94
163	64	3.7	68
179	62	10	72
195	48	6.8	55
211	83	2.7	85
227	95	0.00	95
243	95	0.36	95
259	93	0.52	93
275	78	4.2	82
291	6.6	2.2	8.7
307	93	0.71	93
323	42	12	55
339	3.7	1.8	5.5
355	79	14	93

206 2.2. A cloudy forest site

207 Dense time series data are needed for monitoring vegetation dynamics, and monitoring tropical and subtropical forests is very important to quantifying their important role in the global 208 carbon cycle. However, persistent cloud cover poses challenges when monitoring tropical forest 209 vegetation. To investigate the accuracy of the proposed method to screen clouds and shadows in 210 cloudy tropical forest regions, the second site is northeastern Puerto Rico (Fig. 2). This site has an 211 area of 1,836 km<sup>2</sup> (1200×1700 Landsat pixels), and the central coordinates are 18.321° N and 212 65.838° W. The major land cover type is forest, including the EI Yunque National Forest, where 213 extensive tropical montane cloud forests occur that by definition are persistently cloudy. This site 214

also includes bright, wet and dark features that are easily confused with clouds or cloud shadows. 215 It includes much of the capital city of Puerto Rico, San Juan, coastal areas with features like sand, 216 rock and coral reefs, topographic shadow associated with steep topography and many fields with 217 bright, wet or bare soils. The Landsat WRS-2 scene Path 4 and Row 47 covers the area. A total of 218 18 Landsat 8 OLI images from May 26 2013 to May 29 2014 (i.e., one-year length) and their 219 220 corresponding Fmask cloud layers were downloaded from USGS Earth Explorer. The total cloud and shadow coverage of the images as estimated by Fmask ranges from 5% to 81%, and the mean 221 coverage is 45%, indicating this site is also seriously affected by clouds. In this site, another 11 222 Landsat-4 MSS images from the year 1983 were collected to test the performance of the proposed 223 method for screening clouds and shadows in images with limited bands and low radiometric 224 resolution. For these MSS images, corresponding Fmask cloud masks are not available from USGS 225 Earth Explorer because Fmask uses thermal bands, which are not included in MSS images. 226 Through visual inspection, these 11 MSS images have diverse cloud and shadow coverage, from 227 228 almost cloud-free to fully covered by clouds.



Fig. 2. True color composition of a Landsat-8 image from 2013 (DOY210) of northeastern
 Puerto Rico.

232 233

DOY	Cloud coverage %	Shadow coverage %	Total cloud and shadow coverage%
146	41	4.1	45
178	43	8.0	51
210	4.3	1.6	5.9
226	39	5.6	45
242	40	8.2	48
258	28	6.5	35
274	67	14	81
290	27	8.1	35
306	8.5	1.1	9.6
322	30	6.7	37
354	58	13	71
5	24	11	35
21	43	13	56
53	41	12	54
69	31	9.5	41
117	37	6.5	44
133	38	3.0	41
149	59	14	72

Table 2. Summary of cloud and shadow coverage of Landsat 8 OLI images for the Puerto Rico site using Fmask product

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### 236 2.3. A seasonal-change site

Strong seasonality is another challenge for most multi-temporal cloud screening methods. 237 238 The large variation of spectral values due to seasonality may be confused with the variation due to occurrence of clouds and cloud shadows. To investigate the accuracy of the proposed method to 239 screen clouds and shadows in regions with strong seasonality, the third site is Beijing metropolis 240 241 and its surrounding rural areas (Fig. 3). This site is covered by an entire Sentinel-2 tile (about 12,000 km<sup>2</sup>), and the central coordinates are 40.154 N and 116.495 E. This site has a lot of bright 242 land surface and its vegetation is deciduous with strong seasonality. Images from different seasons 243 244 in Fig. 3 show that vegetation grows to a peak greenness in summer and loses leaves in winter. In 245 addition, the high mountains in this site bring difficulties for both cloud and cloud shadow

detection. Twenty Sentinel-2 images in 2016 with varying cloud cover were downloaded from
USGS Earth Explorer (Table 3). The Fmask cloud masks of these Sentinel-2 images were obtained
using the Matlab code (Version 3.3; https://github.com/prs021/fmask) specific for Sentinel-2
images (Zhu et al., 2015). The total cloud and shadow coverage of the images as estimated by
Fmask ranges from 0.1% to 100%, and 7 images have less than 20%, indicating this site has more
clear images than the other two sites.



14 March, 2016



2 June, 2016



30 September, 2016



9 December, 2016

253	Fig. 3. False-color Sentinel-2 images in Beijing from different seasons in 2016 (the yellow box
254	in upper left image is a forest region of interest (ROI) used to demonstrate the seasonality in Fig.
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265Table 3. Summary of cloud and shadow coverage of 20 Sentinel-2 images in 2016 for the Beijing266site using cloud masks by Fmask

Date Cloud coverage %		Shadow coverage %	Total cloud and shadow coverage%
Jan.26	42.4	17.7	60.1
Mar.14	0.7	0.5	1.2
Mar.24	0.1	0.0	0.1
Apr.3	0.2	0.2	0.4
Jun.2	13.2	1.8	15.0
Jun.12	56.4	6.2	62.6
Jul.22	90.3	9.7	100.0
Aug.1	78.6	4.6	83.2
Aug.11	41.2	9.4	50.6
Aug.21	22.6	6.1	28.7
Aug.31	7.6	4.6	12.2
Sep.20	36.4	3.5	39.9
Sep.30	34.4	8.1	42.5
Oct.10	19.3	7.3	26.6
Oct.20	100.0	0.0	100.0
Oct.30	75.7	9.3	85.0
Nov.19	2.6	1.8	4.4
Nov.29	100.0	0.0	100.0
Dec.9	16.0	8.8	24.8
Dec.29	7.1	4.1	11.2

## **3. Methodology**

There are five main steps in ATSA (Fig. 4). Either TOA reflectance or surface reflectance data can be used as inputs. The five main steps are: (1) compute a cloud index and a shadow index from the image bands to highlight cloud and shadow pixels; (2) detect clouds initially with unsupervised clustering of these indices for individual images in the time series; (3) refine the cloud pixels through analyzing the time series of the cloud index; (4) predict the potential shadow locations through the geometric relationships among the sun, clouds, and the Earth surface; (5) confirm the real shadow pixels through analyzing the time series of the shadow index. We detail these steps below.





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shadows

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Given the wide ranges of reflectance values exhibited by diverse cloud- and Earth surface types, individual spectral bands from one image cannot accurately differentiate clouds, cloud

shadows and clear observations. With image time series, significant seasonality of some land cover types (*e.g.*, natural vegetation and agriculture) and land-cover change (*e.g.*, deforestation and urbanization) lead to large temporal variability of reflectance in image time series, which is easily confused with temporal variability caused by clouds and cloud shadows. Therefore, there is a need to combine or transform individual bands to get indices that highlight the clouds and cloud shadows while compressing variability in other land cover types, so that clear observations have values that are as stable as possible in the time series.

As land and water surfaces have very different spectral characteristics (Zhu and Woodcock, 292 2012), the cloud and shadow indices are designed separately for land and water surfaces. A water 293 mask is needed in our method. Fortunately, a water mask can be easily obtained through classifying 294 a cloud-free image in the time series or from an existing water mask. There are now several water 295 masks available at different resolutions, such as a 30-m water mask from a Landsat-based global 296 land cover product (Chen et al., 2015) and a 250-m global water mask from MODIS data (available 297 298 in http://landcover.org/data/watermask/). In our test experiments, we classified a cloud-free image to obtain the water mask. 299

For land surfaces, we used the haze optimal transformation (HOT) as a cloud index. The HOT transformation is derived from an analysis of Red-Blue spectral space. These two bands have a perfect linear relationship for diverse land cover types under clear-sky conditions (Zhang et al., 2002), and Zhang et al. (2002) name this perfect line the clear-sky line (see the red line in Fig. 5 a). For pixels contaminated by haze and clouds, their spectral response in Red-Blue space is very different from the clear-sky line, so the HOT index was designed to quantify the perpendicular distance of a pixel from the clear-sky line:

$$HOT = \frac{|a \times B_{Blue} - B_{Red} + b|}{\sqrt{1 + a^2}} \tag{1}$$

307 where  $B_{Blue}$  and  $B_{Red}$  are pixel values of blue and red bands respectively, and *a* and *b* are the slope 308 and intercept of the clear-sky line.



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Fig. 5. A land-surface subset of a Landsat-8 image and its Red-Blue scatter plot (a) and a water-

311 surface subset Landsat-8 image and its NIR-Blue scatter plot (b)

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In the original HOT transformation (Zhang et al., 2002), the clear-sky line comes from regressing spectral values of pixels selected from areas of a scene that visually are deemed to be the clearest. To make our method automatic, we employed a bin-based approach to search for the clearest pixels in each scene. This approach has three steps: (1) divide the 0-0.15 range of blue

reflectance values into 50 bins with equal intervals, because cloud-free pixel values of most land-317 cover types are within this range; (2) for each bin, select the 20 pixels with the largest reflectance 318 values in the Red band and compute the average value of these selected pixels for red and blue 319 bands, respectively, yielding a pair of red and blue reflectance values for each bin (B<sub>Redi</sub>, B<sub>Bluei</sub>); 320 (3) for all 50 pairs of (B<sub>Redi</sub>, B<sub>Bluei</sub>), regress B<sub>Redi</sub> against B<sub>Bluei</sub> to get the clear-sky line using the 321 322 least absolute deviation (LAD) regression method to avoid the effect of outliers (Bassett and Koenker, 1978). If some images in the time series are completely covered by clouds, no clear 323 pixels can be found for estimating the clear-sky line. For these completely cloud-covered images, 324 325 the average slope and intercept of clear-sky lines derived from other images in the time series are used to compute the HOT index. To demonstrate the effectiveness of a bin-based automatic 326 approach, the retrieved clear-sky line in a sub-image was compared with the result using manually 327 selected clear pixels (Fig. 6). The slope and intercept of the clear-sky line from the bin-based 328 approach is very similar to the results from the manual approach. 329



Fig. 6. Comparison between the clear-sky line of a sub-image estimated by the proposed
automatic bin-based approach (a) and that using manually selected clear pixels marked by red
ROIs (b).

For water surfaces, the cloud-free pixel values of the red and blue bands are not on the clearsky line, leading to large HOT values that are confused with thin clouds. Consequently, a new HOT index, designed specifically for water surfaces, is needed. In the Blue-NIR space, the spectral response of cloud-free water pixels, including turbid or shallow water and coral reefs, is very different from cloudy pixels (Fig. 5b). A new HOT index for water surface,  $HOT_w$  is given as:

$$HOT_w = \frac{|a_w \times B_{NIR} - B_{Blue} + b_w|}{\sqrt{1 + {a_w}^2}}$$
(2)

where  $a_w$  and  $b_w$  are the slope and intercept of the clear-sky line for water bodies and are obtained through the same method as for the land surface. Then, the HOT indices for land and water surfaces are combined to yield a cloud index map (Fig. 7b). In this cloud index map, we can see that the HOT transformation yields an index with a larger difference between cloud and bright non-cloud objects than the individual visible bands. All clouds and haze are highlighted by larger values (i.e. white color) while all cloud-free pixels have a very low value (i.e. dark color).



Fig. 7. False color Landsat-8 image of DOY149 in the Puerto Rico site (a), its corresponding
HOT cloud index (b), initial cloud mask (c), and final cloud mask (d). In (c) and (d): gray is clear
pixels and white is clouds. The time series analysis adds thin clouds to the initial cloud mask,
and the minority analysis removes scattered bright pixels in urban and coastal areas in the upper

left of panel (c), which would otherwise be confused with clouds.

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To further compare the ability of the original bands and HOT for discriminating clouds and clear land surface, the relative difference (*RD*) between cloud and cloud-free pixels in each image was computed:

$$RD = \frac{\bar{B}_{cloud} - \bar{B}_{clear}}{\bar{B}_{cloud}} \tag{3}$$

where  $\bar{B}_{cloud}$  and  $\bar{B}_{clear}$  are average values of cloudy pixels and clear pixels respectively. RD 356 ranges from 0 to 1 and larger values indicate a higher separability between cloudy and clear pixels. 357 Fig. 8 shows the RD values of the Blue band and the HOT index of Landsat-8 images which contain 358 both clear and cloudy pixels in the Hong Kong site. Hong Kong includes both forests and 359 360 considerable bright urban surfaces. It is a challenging site for cloud detection. We can see that in these images HOT index is better than the original Blue band at separating clouds from clear land 361 surfaces. The comparisons of RD values between the Red band and the HOT index, and between 362 363 the NIR band and the HOT index, have a similar pattern (results not shown).



Fig. 8. Relative difference (*RD*) between the average value of cloudy and clear pixels of the Blue

band and the HOT index in Hong Kong Landsat-8 images. A larger *RD* indicates higher
 separability between cloud and clear pixels.

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For cloud shadows, direct solar radiation is blocked by clouds, so the shadow pixels are illuminated by scattered light. Because the atmospheric scattering is weaker at longer wavelengths, the NIR and SWIR bands of shadow pixels are much darker than surrounding clear pixels (Zhu et al., 2015). Therefore, the shadow index (*SI*) is defined as:

$$SI = B_{NIR} + B_{SWIR} \tag{4}$$

However, water also absorbs most radiation at longer wavelengths, so water pixels not obstructed by clouds are as dark as shadow pixels in NIR and SWIR bands (Li et al., 2017). Consequently, for water surfaces, the shadow index is calculated with the blue and green bands:

$$SI_w = B_{Blue} + B_{Green} \tag{5}$$

For old satellite images with fewer bands, such as Landsat MSS images with only green, red, and 2 NIR bands, the green band replaces the blue band in Eqs. (1, 2, and 5), because it is highly correlated with blue band. Also, the second NIR band replaces the SWIR band in Eq. (4), because both the NIR and SWIR bands are good indicators of cloud shadows. Similarly, for other sensors without SWIR bands, such as IKONOS, we anticipate that only one NIR band would be used as the shadow index for land surfaces.

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383 3.2. Detect cloud initially

All cloud index images of the time-series are classified by an unsupervised classifier, k-means, to get the initial cloud mask. First, a certain number (e.g. 10,000) of sample pixels are selected by systematic sampling of all cloud index images. Selecting samples from all images in the time series

ensures that samples of clear surfaces, thin clouds, and thick clouds are included. Using the 387 selected samples rather than all pixels speeds up the k-means optimization in the next step. Second, 388 389 these samples are classified with the k-means method into three classes. The three classes are labeled based on the relative value of the class means, i.e., the lowest class mean is clear pixels, 390 the middle one is thin clouds, and the highest one is the thick clouds. The k-means method uses an 391 392 iterative procedure. At each iteration, each sample is assigned to one class based on the closeness to the class means obtained from the last iteration, and new class means are updated using new 393 394 class labels of samples. The iterative process will be ended when the class labels no longer change (Lloyd, 1982). Third, individual pixels in each cloud index image of the time series are identified 395 as thin clouds, thick clouds, or cloud-free observations based on which class has the smallest the 396 cloud-index distance from the class means of the sample pixels. Finally, an initial cloud mask is 397 produced for each image by combining thin clouds with thick clouds (Fig. 7c). The ranges of cloud 398 index values for the three classes (thin clouds, thick clouds and clear), being derived from all pixels 399 400 in the time series, form a set of thresholds that are adapted to a time series rather than a single image. 401

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3.3. Detect remaining extremely thin clouds and remove bright pixels that are not cloud or haze
Although the initial cloud mask already identifies most cloudy pixels, it may omit some
cloudy pixels, especially extremely thin clouds and cloud edges with lower values of the cloud
index. Therefore, the initial cloud masks need to be further improved with temporal information.
For each pixel, its time series may include both cloudy points and cloud-free observations. In
general, cloudy points have larger variations in spectral values than clear observations. This
temporal property can help to identify cloudy points (Zhu and Woodcock, 2014). However, due to

changes in vegetation phenology or land cover, clear observations of some pixels also undergo 410 temporal variations. However, compared with the original spectral bands, the cloud index derived 411 412 from the HOT transformation depresses the temporal variations from different vegetation growth stages, soil inundation, or land-cover change. For instance, in a forest ROI from the Beijing site 413 (marked by a yellow box in Fig. 3), the time series of the Red band shows a stronger seasonality 414 415 than the HOT index (Fig. 9). The Red band is used to compute the HOT index. It has lower values in summer due to more absorption by vegetation. In contrast, the HOT index is more stable across 416 different seasons and with smaller variability than the original Red bands. 417



Fig. 9. Red reflectance and HOT index of cloud-free pixels in a forest ROI in Beijing (marked by
a yellow box in Fig. 3) across different seasons. The circles are mean values and error bars are ±
1 standard deviation.

422

Therefore, an analysis of the cloud index time series is conducted to refine the initial cloud mask. Fig. 10 gives an example of a cloud index time series of a pixel from the Puerto Rico site (column1173, row 1092). There are two points identified as clouds in the initial mask (the red points). Most of the other points (the black ones) should be clear observations. They are used to find an upper boundary in the HOT index for clear pixels, U(i). Points above this threshold are then also designated as cloudy. For *i*th pixel:

$$U(i) = \operatorname{mean}\{HOT(i,t)|(i,t) \notin \mathbf{C}\} + A \times \operatorname{sd}\{HOT(i,t)|(i,t) \notin \mathbf{C}\}$$
(6)

where sd{·} is the standard deviation of the HOT index through the time series, HOT(i, t) is the 429 HOT index value of the *i*th pixel at time *t*, and **C** is the set of cloudy points from the initial masks 430 for *i*th pixel. A is a standard deviation multiplier that defines the upper boundary. A can be assigned 431 432 a recommended value from 1 to 2. Smaller values would be able to identify thinner clouds, but 433 meanwhile increase the risk of commission errors, *i.e.*, identifying "clear" observations as cloudy 434 points. In existing methods, this parameter is a constant for all pixels in the image (Goodwin et al., 435 2013; Hagolle et al., 2010). However, cloud frequency is different in different parts of the image, 436 so some pixels may include more cloud points in the HOT time series that are omitted in the initial detection than others. Therefore, we need to consider this difference among pixels when we set the 437 438 value of parameter A. In general, clouds cause large variations in the HOT time series. We introduced a new variable, the normalized difference range index (NDRI), to tune the parameter 439 A: 440

$$NDRI(i) = (T_{kmeans} - Range_i) / (T_{kmeans} + Range_i)$$
(7)

$$Range_i = \max\{HOT(i,t)|(i,t) \notin \mathbf{C}\} - \min\{HOT(i,t)|(i,t) \notin \mathbf{C}\}$$
(8)

441 where  $T_{kmeans}$  is the minimum HOT value of all cloud pixels identified by *K*-means in section 3.2. 442 *NDRI* is further used to adjust the parameter *A* in Eq. (6) as a pixel-wise parameter *A*(*i*):

$$A(i) = A + NDRI(i) \tag{9}$$

The value for A(i) is used in Eq. (6) to calculate the pixel-level upper boundary U for each pixel in the time series. A(i) further tunes the pixel-level upper boundary U by adapting the standard deviation multiplier to the temporal variability of each pixel. Because *NDRI* is added to A, we recommend an A value from 0.5 to 1.5 (instead of 1 to 2). Pixels with larger variation in the HOT time series will have a lower upper boundary, i.e. a stricter threshold. Any points above the upper
boundary, e.g. the dashed line in Fig. 10, will be identified as clouds.



449

450 Fig. 10. An example of cloud index time series: the two black points above the dashed line are451 identified as clouds based on the time series analysis.

The cloudy points detected from the time series analysis are the final cloud mask (Fig. 7d). 452 This step adds more thin clouds to the initial mask and also contributes to filtering bright non-453 cloud objects. For instance, very bright land surfaces (e.g., airport runways and beach sand) may 454 455 show consistently high values in the cloud-index time series, leading to a high threshold in Eq. (6). As a result, pixels of these bright land surfaces are not likely to be identified as clouds because 456 457 their cloud index values are unlikely to exceed the high threshold. In addition, assuming that clouds 458 are generally wider than a few pixels at Landsat spatial resolution, isolated pixels identified as 459 being cloudy are removed from the cloud mask using a repeated minority analysis. We removed 460 cloud pixels if 4 or fewer pixels in the 3-by-3 neighborhood of a pixel are cloud pixels. This step removes any remaining isolated, bright pixels in urban and coastal areas that are not clouds. Finally, 461 462 similar to Fmask, all cloud patches are buffered with a width of 1 pixel to further reduce omission errors around cloud edges. 463

## 465 3.4. Estimate potential shadow zones

Shadow pixels are easily confused with dark objects and topographic shadow even in shadow index maps. However, clouds always accompany cloud shadows except at scene edges. This characteristic can help reduce commission (e.g., wet soil, topographic shadow) errors of cloud shadow detection. Actually, the location of cloud shadows can be calculated by the precise geometric relationship among clouds, cloud shadows and the position of the sun (Zhu and Woodcock, 2012). For a cloudy pixel with coordinates (*x*, *y*), the location of its corresponding shadow pixel (*x'*, *y'*) can be calculated using following equations (Luo et al., 2008):

$$x' = x - H \times \tan\theta \sin\phi \tag{10}$$

$$y' = y - H \times \tan\theta \cos\phi \tag{11}$$

where H is the height of clouds above the land surface, and  $\theta$  and  $\phi$  are solar zenith and azimuth 473 angles. Values of  $\theta$  and  $\phi$  can be extracted from the image metadata files, but H is unknown for 474 each cloud patch. In most existing methods, the brightness temperature (BT), derived from thermal 475 infrared bands, is used to estimate cloud height with lapse rates for air temperature, such as -476 9.8K/Km for dry air and -6.5K/Km for moist air (Goodwin et al., 2013; Huang et al., 2010; Zhu 477 and Woodcock, 2012). However, there are two problems with estimating cloud height when 478 locating cloud shadows: (1) the lapse rate varies in different atmospheric conditions, and the BT 479 of thin clouds is also influenced by the land surface; (2) some sensors, especially old ones, do not 480 481 have thermal infrared bands, such as the Landsat MSS sensor, the CBERS IRMSS sensor and the Sentinel MSI sensors. For Landsat 8 also, a method for masking clouds and shadows that does not 482 require a thermal band is needed. The Thermal Infrared Sensor (TIRS) has some error and 483 intermittent availability and has a shorter design life than the multispectral Operational Land 484

Imager (OLI). That thermal data may not always be available is one obstacle to improving Landsat
8 cloud and shadow masks with image time series (Foga et al., 2017; Scaramuzza et al., 2012).

To make the proposed method able to process historical images without thermal infrared bands, a range of possible cloud heights are used to estimate all possible shadow locations of a cloud. We can use a default value of 200 m for minimum cloud heights because it is suitable for most areas (Zhu and Woodcock, 2012). The maximum cloud heights can be determined empirically by visually checking the maximum horizontal distance ( $D_{max}$ ) between clouds and their shadows, or using 12 km based on previous studies (Fisher, 2013; Luo et al., 2008) :

$$H_{max} = \frac{D_{max}}{\sqrt{(\tan\theta\sin\phi)^2 + (\tan\theta\cos\phi)^2}}$$
(12)

Fig. 11b shows an example of potential shadow zones of a subset of image DOY146 in the
Puerto Rico site. We can see that the real shadows are located within the potential shadow zones.





Fig. 11. A subset of the Landsat-8 image DOY146 in the Puerto Rico site (a), its potential shadow
zones (b), shadow darkness as estimated by Inverse Distance Weighting (IDW) (c), and the initial
shadow detected by K-means (d).

501 3.5. Detect shadow within potential shadow zones

502 The potential shadow zones mark the possible locations of cloud shadows. They overestimate 503 the real shadow areas. Therefore, all the pixels within the shadow zones need to be further 504 confirmed as to whether they are real shadow pixels. In the shadow index images, shadows are 505 located at places with regional minima (i.e. "holes") due to their being relatively dark in optical 506 bands. Some existing methods use flood-fill transformation to predict the image without shadows 507 and compare it with real images to identify shadow pixels (Li et al., 2017; Zhu and Woodcock,

2012). However, this approach may often mislabel dark objects, such as water, as cloud shadow 508 (Li et al., 2017). Here, a similar idea is employed, but the new strategy reduces errors as compared 509 510 with the flood-fill method. First, in the shadow index images, pixels in potential shadow zones are predicted from surrounding clear pixels with an inverse distance weighted (IDW) interpolator. 511 Second, for those pixels in potential shadow zones, we estimate their "darkness" as their original 512 513 shadow index minus the predicted values (Fig. 11c, a darker color means higher darkness). This darkness only shows how cloud shadows lower the pixel brightness compared with surrounding 514 clear pixels. Third, similar to initial cloud detection, K-means clustering is applied to these 515 darkened pixels (i.e. pixels with negative darkness values) to classify these pixels into two classes, 516 clear observation and cloud shadow, to yield an initial shadow mask (Fig. 11d). 517

After the initial shadow detection, we apply a time series analysis, similar to the cloudy point 518 refinement, to refine the initial shadow mask. This process aims to reduce both omission and 519 commission errors in the initial shadow mask. Because cloud shadows have darkening effects, 520 521 which lead to lower shadow index values in the time series of a pixel, a lower boundary L is used as a threshold to identify real shadow points. Considering differences in earth-sun-sensor geometry, 522 atmospheric conditions and vegetation phenology, the shadow index of land surfaces needs to be 523 524 normalized to minimize these differences prior to the time series analysis. Here, the histogram matching method is used given its simplicity (Helmer and Ruefenacht, 2005). Although histogram 525 526 matching is a linear correction, and changes in vegetation phenology across an image can be 527 nonlinear (Helmer and Ruefenacht, 2007), we found that histogram matching worked well for mitigating the temporal variability in shadow-index time series. First, the image with the fewest 528 529 clouds in the time series is selected as a base image. Then, the shadow index of other images is 530 normalized to this base image using the gain and bias:

$$gain = \frac{\sigma_B}{\sigma_t} \tag{13}$$

$$bias = \mu_B - \mu_t \times gain$$

where  $\mu_B$  and  $\mu_t$  are the mean value of clear pixels in the base image and the image at time *t* respectively,  $\sigma_B$  and  $\sigma_t$  are the standard deviations of clear pixels in the base image and the image at time *t* respectively. The normalized shadow index value of image at time *t*,  $SI_N(i, t)$ , can be computed as:

$$SI_N(i,t) = SI(i,t) \times gain + bias$$
 (14)

535 This lower boundary L is defined using "good" points, which are those points not identified as 536 shadow in the initial shadow mask (Fig. 12):

$$L(i) = \operatorname{mean}\{SI_N(i,t)|(i,t) \in \operatorname{"good"}\} - B \times \operatorname{sd}\{SI_N(i,t)|(i,t) \in \operatorname{"good"}\}$$
(15)

where B is a standard deviation multiplier that serves as a parameter to tune the threshold, L(i) is 537 the lower threshold for pixel *i*, mean is the mean shadow index (SI) of pixel *i* for the time series, 538 and sd is the standard deviation of the SI for the time series of pixel i. Pixels with SI brighter than 539 L(i) are deemed too bright to be cloud shadow. The recommended value of B is from 1 to 3, and a 540 larger value will select *darker* shadows, *i.e.*, it will darken the threshold for designating whether 541 pixels are shadow. Therefore, the parameter B should be set to balance the omission and 542 commission errors for shadow detection. For the initial shadow points, they are confirmed as real 543 shadow if their shadow index values are lower than the mean value of "good" points. This step 544 reduces the commission errors in initial shadow detection. For other points in the time series which 545 are marked as potential shadow using sun-cloud geometry, they will be identified as final shadow 546 points if their shadow index values are lower than L (Fig. 12). This step reduces the omission errors 547

548 in the initial shadow detection. It should be noted that although the potential shadow zones and 549 time series analysis can greatly prevent classifying topographic shadow as cloud shadow, 550 topographic shadow within the potential shadow zones may be identified as cloud shadows if it is 551 as dark as cloud shadow.



552

Fig. 12. An example of shadow index time series analysis: the points below the dashed line areidentified as cloud shadow. The dashed line represents *L*, the lower threshold.

555

556 Similar to the cloud mask, isolated shadow pixels are also filtered out by a repeated minority 557 analysis in a 3-by-3 neighborhood, and then all shadows are buffered, with a width of 1 pixel, to 558 obtain the final shadow mask. The final shadow mask is combined with the final cloud mask to 559 get the final product of cloud and shadow mask.

560

### 561 3.6. Evaluation and comparison

To demonstrate the accuracy and effectiveness of the proposed method, it was compared with Fmask (Zhu and Woodcock, 2012), one of the most advanced single-image methods and used by USGS to produce the standard cloud mask for Landsat images. The results of Fmask can be considered as a benchmark to assess the performance of ATSA. Both ATSA and Fmask were applied to Landsat-8 OLI and Sentinel-2 images, while only ATSA was applied to Landsat-4 MSS images in the second site, because MSS images lack not only thermal bands but also other bands that are needed by Fmask. We found that Fmask detected many clouds as snow in some images in the Hong Kong and Puerto Rico sites (Fig. 13). Because these two test sites are subtropical and never have snow, we merged snow into clouds before the comparison, but this adjustment was not made for cloud masks in Beijing site because it can snow in winter.

572



Fig. 13. Landsat-8 image DOY178 in the Puerto Rico site (a) and its original Fmask cloud mask(b) showing where clouds are classified as snow (light blue color).

576

573

In the comparison, the agreement between these two methods was evaluated. First, the 577 percentage of clouds and cloud shadows of all methods were plotted together to check their 578 difference. Second, matrices were built comparing the proposed ATSA and Fmask methods, and 579 the overall agreement derived from these matrices was used to assess the pixel-wise agreement 580 between ATSA and Fmask. Third, representative images selected from the time series were 581 digitized to produce reference cloud and shadow masks. The digitizing work was done by 582 583 experienced experts who were not involved in the development of ATSA. Then, these digitized maps were used to quantitatively evaluate the accuracy of both methods. It should be noted that 584

585 the manual mask of cloud and cloud shadow is not 100% accurate. It may include some 586 commission or omission errors.

587 **4. Results** 

588 4.1. Hong Kong site

Among 23 images to which we applied the two cloud and shadow masking methods, the two 589 methods detect similar cloud cover for 19 of them (Fig. 14), while for the other 4 images (DOY 590 131, 179, 195, and 339) there are large differences. For the 19 images with similar cloud coverage, 591 592 visual inspection confirms that both methods successfully detect clouds (see image of DOY 51 as 593 an example in Fig. 16). For images of DOY 131 and 339, ATSA detected many more clouds than Fmask. On the other hand, for images of DOY 179 and 195, ATSA detected far fewer clouds than 594 595 Fmask. Unlike cloud coverage, shadow coverage detected by the two methods slightly differs in most of the 23 images except the image DOY 355 (Fig. 14). Visual inspection of this image 596 597 confirms that Fmask detected all water surface as cloud shadow. There are 5 images in the time 598 series with large disagreement between ATSA and Fmask (Fig. 15). Agreement between ATSA and 599 Fmask for the images of DOY 131, 179 and 339 is even lower than 50%.

600









Fig. 15. Overall agreement of cloud and shadow masks of 23 Landsat-8 OLI images in the HongKong site between ATSA and Fmask.

604

In the cloud masks of the three images with the least agreement between ATSA and Fmask, 608 it is clear from Fig. 16 that ATSA more accurately identified clouds. Fmask underestimated clouds 609 in two images of DOY 131 and 339, and it overestimated clouds in the image of DOY 179. 610 Specifically, Fmask failed to screen many of the thin clouds in the center of the image of DOY 611 131, and it failed to identify many of the thick clouds in the image of DOY339, even though these 612 613 clouds appear very bright in all visible and NIR bands. In the image of DOY179, Fmask 614 misidentified most of the clear water and some clear land surface (see the island in the lower right) as clouds, which led to serious overestimation of cloud cover. For the cloud shadows, it appears 615 616 that ATSA successfully identified most shadows adjacent to clouds. Fmask identified some clear 617 pixels as shadow that were near the misidentified cloud patches (see image of DOY 179 in Fig. 618 16).



Fig. 16. False color composite of selected Landsat images (upper row) and their cloud masks by
Fmask (middle row) and ATSA (lower row) for the Hong Kong site (gray: clear pixels; black:
shadows; white: clouds)

Quantitative accuracy assessment for the four images in Fig. 16 using manual masks shows 625 that ATSA and Fmask obtain comparable overall accuracy for the image of DOY51, but ATSA's 626 627 overall accuracy is much higher than Fmask for the other three images (Table 4). For cloud detection, ATSA obtained user's accuracies ranging from 0.85 to 0.99 and producer's accuracies 628 ranging from 0.89 to 0.99. The accuracy of ATSA cloud mask for the image of DOY179 is lower 629 than that of the other images due to the errors in haze detection on the water surface (see haze in 630 the lower right part of this image). In contrast, the cloud producer's accuracy of Fmask is low for 631 images of DOY131 (0.41) and DOY339 (0.04) because of large omission errors. The cloud user's 632

633	accuracy of Fmask is low for images of DOY179 (0.08) because it misidentified many clear pixels
634	as clouds. For shadow detection, ATSA can obtain producer's accuracy higher than 0.82. The
635	shadow user's accuracy is also high except the image of DOY179, in which ATSA overestimated
636	the shadow area. In the context of applications with cloudy images, the producer's accuracy is
637	more important than user's accuracy, because end users hope to exclude all contaminated pixels in
638	their analysis, and meanwhile they can allow commission errors to some extent (Zhu and
639	Woodcock, 2012). Both user's and producer's accuracies of shadow detection by Fmask are much
640	lower than ATSA. Specifically, Fmask detected fewer shadows in the image of DOY51 and
641	identified clear pixels near the wrong cloud patches as shadow in other three images in Fig. 16.

Table 4 Accuracy assessment of cloud masks of the 4 images in Fig. 16 in the Hong Kong site:

648	overal	l accuracy (	[oa]	), user	's accuracy	(ua)	) and	prod	ucer	's accuracy	(pa)	).
-----	--------	--------------	------	---------	-------------	------	-------	------	------	-------------	------	----

			Cloud		Cloud Shad	
DOY		oa	Ua	ра	иа	ра
51	Fmask	0.93	0.97	0.98	0.60	0.49
51	ATSA	0.99	0.99	0.99	0.95	0.87
121	Fmask	0.45	0.99	0.41	0.04	0.10
131	ATSA	0.98	0.99	0.99	0.93	0.82
170	Fmask	0.29	0.08	0.98	0.08	0.26
1/9	ATSA	0.97	0.85	0.89	0.67	0.90
220	Fmask	0.06	0.95	0.04	0.00	0.00
539	ATSA	0.99	0.99	1.00	0.89	1.00

# 650 4.2. Puerto Rico site

#### 651 4.2.1. Landsat-8 OLI images

Among the 18 images, ATSA and Fmask obtained similar cloud coverage in 15 images, while 652 ATSA detected many more clouds in three images (DOY178(2013), 306(2013), and 053(2014)) 653 than Fmask (Fig. 17). For shadow coverage, ATSA detected slightly more shadows than Fmask in 654 most images. Through visual inspection of these images, we found that Fmask underestimated 655 656 shadows surrounding small cloud patches, which leads to smaller shadow percentage than ATSA. On the other hand, in the images DOY 274(2013), 354(2013), 53(2014), and 149(2014), ATSA 657 detected fewer shadows than Fmask. Visual inspection shows that these four images only have 658 large cloud patches. Fmask overestimated shadow cover of these large cloud patches. The 659 quantitative assessment of pixel-wise agreement between the two methods is good (higher than 660 80%) for masks of most images, but the masks of two images, DOY 178(2013) and 306(2013) 661 have agreement between ATSA and Fmask that is lower than 60% (Fig. 18). For the images with 662 good agreement between ATSA and Fmask, both methods successfully detect clouds (e.g. image 663 DOY146(2013) in Fig. 19). In the two images with the least agreement between ATSA and Fmask, 664 Fmask omitted a lot of thin clouds in west region in the image of DOY178(2013), and it missed a 665 lot of cloudy pixels, even of thick clouds, in the image of DOY306(2013) (Fig. 19). 666



668 Fig. 17. Cloud and shadow coverage of 18 Landsat-8 OLI images in the Puerto Rico site detected

# 669 by ATSA and Fmask

670



671

Fig. 18. Overall agreement of cloud and shadow mask between ATSA and Fmask for the Puerto

673

Rico site.



Fig. 19. False color composite of the three Landsat images in the Puerto Rico site (upper row)

676	and their cloud and shadow masks by Fmask (middle row) and ATSA (lower row) (gray: clear
677	pixels; black: shadows; white: clouds)

The quantitative accuracy assessment of the cloud masks of these three images in Fig. 19, 678 using manual masks, shows that the overall accuracy of ATSA cloud and shadow masks ranges 679 from 0.97 to 0.98, which is much higher than Fmask (Table 5). Cloud producer's and user's 680 681 accuracy of ATSA reaches 0.97 in all three images, while cloud producer's accuracies of Fmask are only 0.12 to 0.52 for the image of DOY178 and 306. Shadow producer's and user's accuracies 682 of ATSA are lower than the cloud mask accuracy, but it is still much higher than Fmask. ATSA 683 omitted some thin shadows on land surfaces in the lower part of image of DOY306 (Fig. 19) 684 leading to a relatively lower producer's accuracy of 0.86. Similar to the Hong Kong site, Fmask 685 detected fewer shadows than the real situation, leading to low producer's accuracy in shadow 686 detection. 687

688

Table 5. Accuracy assessment of cloud masks of three images in Fig. 19 in the Puerto Rico site:

			Cle	Cloud		dow
DOY(Year)		oa	иа	pa	иа	pa
146(2013)	Fmask	0.90	0.85	0.98	0.58	0.34
140(2013)	ATSA	0.98	0.97	1.00	0.92	0.94
178(2013)	Fmask	0.53	0.98	0.52	0.07	0.08
178(2013)	ATSA	0.98	0.98	0.99	0.97	0.96
306(2013)	Fmask	0.28	0.98	0.12	0.03	0.00
500(2015)	ATSA	0.97	1.00	0.98	0.97	0.86

690 overall accuracy (*oa*), user's accuracy (*ua*) and producer's accuracy (*pa*).

691

692 4.2.2. Landsat-4 MSS images

Only ATSA was applied to the 11 Landsat-4 MSS images in the Puerto Rico site, because
 Fmask needs more bands than MSS images have. Four images with representative cloud coverages

(6.8% to 99% clouds; 0% to 12.6% cloud shadows) were selected for further assessment (Fig. 20). 695 The cloud coverages of these four selected images are 6.8% (DOY40), 31.1% (DOY280), 41.4% 696 (DOY24), and 66.1% (DOY200). Visual inspection shows that ATSA successfully identified most 697 clouds and shadows in these MSS images (Fig. 21), including the thin clouds in the image of 698 DOY200. We can see that the cloud user's and producer's accuracy of all four MSS images are 699 700 higher than 0.95 (Table 6), indicating that ATSA successfully screened clouds in these images with very small omission and commission errors. For shadow accuracy, in terms of producer's accuracy 701 (more important for applications in our opinion), it is high enough in image DOY 40, and 280, 702 703 reaching 0.97. The image DOY200 has shadow producer's accuracy of 0.83 which is caused by the identification of shadows as clouds in the lower right part of this image (Fig. 21). 704



705

Fig. 20. Cloud and shadow coverage of 11 Landsat MSS images in the Puerto Rico site detectedby ATSA.



		Cloud		Sha	dow
DOY	oa	Ua	ра	иа	ра
24	0.96	0.98	0.99	0.88	0.87
40	0.99	1.00	0.99	0.89	1.00
200	0.97	0.99	0.97	0.90	0.83
280	0.98	1.00	0.98	0.92	0.97

719 4.3. Beijing site

In general, the cloud coverage detected by ATSA is smaller than that of Fmask in the 20 Sentinel-2 images. The cloud coverage difference between ATSA and Fmask is larger than or equal to 20% in three images, Aug. 21, Sep.30, and Oct.10 (Fig. 22). Through visual inspection of these 723 cloud masks, we found that both ATSA and Fmask misclassify some pixels in very bright urban surfaces as clouds, but this commission error of Fmask is more serious than ATSA. Fmask detected 724 nearly all urban pixels as cloud or snow in the three images of Aug.21, Sep.30, and Oct.10. Similar 725 to cloud coverage, shadow coverage detected by ATSA is generally lower than Fmask, except for 726 images dated Oct. 30 and Nov. 29. The larger shadow coverage detected by Fmask results from 727 the commission errors of cloud detection. In other words, Fmask detected many clear pixels as 728 shadow surrounding areas wrongly detected as clouds. Regarding the pixel-wise agreement 729 between ATSA and Fmask (Fig. 23), there are 7 images with overall agreement lower than 80% 730 731 and 3 images lower than 70%. These three images are Jan.26, Aug.11, and Sep.30.





Fig. 22. Cloud and cloud shadow coverage of 20 Sentinel-2 images in the Beijing site detected

by ATSA and Fmask.



Fig. 23. Overall agreement of cloud and shadow mask between ATSA and Fmask for the 20
Sentinel-2 images from the Bejing site.

735

For the images with high agreement between ATSA and Fmask, both methods successfully 739 detect clouds and shadows (see Sep.20 image as an example in Fig. 24). In the three images with 740 the least agreement between ATSA and Fmask, ATSA is generally more successful than Fmask for 741 identifying clouds and shadows. Specifically, in the Jan. 26 image, both ATSA and Fmask detected 742 most of clouds. ATSA does not have snow detection step, so the snow in the northwest was 743 misclassified as cloud in ATSA, while Fmask successfully detected these snows but it also 744 745 identified many clear pixels as snow. Fmask also detected many clear and thin-cloudy pixels as snow in summer image where no snow events should happen. In addition, Fmask detected many 746 clear urban pixels as cloud in the Sep. 30 image (Fig. 24). For cloud shadow detection, ATSA is 747 more successful although it detected topographic shadows as cloud shadows in the Jan. 26 image. 748 Images of mountainous areas have more topographic shadows in spring and winter due to the lower 749 sun elevation. It may lead to larger commission errors in cloud shadow detection if these 750 topographic shadows are within the potential shadow zones. In contrast, Fmask failed to detect 751

- shadows which are distant from the cloud patches. The possible reason is that Fmask for Sentinel-
- 2 assumes the cloud height between 200 m and 1,200 m for all images (Zhu et al., 2015).



754

Fig. 24. False color composite of the four Sentinel-2 images in the Beijing site (upper row) and
their cloud mask by Fmask (middle row) and ATSA (lower row) (gray: clear pixels; black:
shadows; white: clouds; snow: light blue)

The quantitative accuracy assessment also demonstrates that ATSA can obtain more accurate cloud and shadow masks than Fmask (Table 7). For the Sep. 20 image, both ATSA and Fmask can obtain acceptable accuracy in cloud detection. For other three images, cloud producer's accuracy of ATSA ranges from 0.81 to 0.96 and cloud user's accuracy ranges from 0.92 to 0.99. In contrast, cloud user's accuracy of Fmask for the Sep.30 image is very low because of large commission

errors, and the low cloud producer's accuracy of Fmask for the Aug.11 image is caused by large 764 omission errors. For the Jan. 26 image, the producer's accuracy of cloud detection for both Fmask 765 and ATSA is only 0.81 because both methods omitted extremely thin clouds. For shadow detection, 766 ATSA obtains good producer's accuracy ranging from 0.81 to 0.96, which is much higher than 767 Fmask ranging from 0.20 to 0.50, indicating that Fmask omitted considerable cloud shadow in 768 769 these images. For the Jan. 26 image, the user's accuracy of shadow detection by ATSA is only 0.5, because it detects many black rocks and topographic shadows as cloud shadows. In this 770 mountainous area, some snow and ice pixels were misclassified as clouds, which makes the black 771 772 rocks and topographic shadows within the potential shadow zone. This issue can be solved if the commission error in cloud detection is reduced, especially for distinguishing snow and ice from 773 clouds. 774

775

Table 7. Accuracy assessment of cloud masks of images in Fig. 24: overall accuracy (*oa*), user's
accuracy (*ua*) and producer's accuracy (*pa*).

			Cloud		Shadow	
Date		oa	иа	ра	иа	ра
Sep.20	Fmask	0.89	0.80	1.00	0.20	0.41
	ATSA	0.98	0.99	1.00	0.62	0.81
Jan.26	Fmask	0.67	0.74	0.81	0.33	0.50
	ATSA	0.79	0.92	0.81	0.50	0.87
Aug.11	Fmask	0.68	0.77	0.65	0.21	0.46
	ATSA	0.97	0.99	0.96	0.94	0.86
Sep.30	Fmask	0.58	0.00	0.92	0.00	0.20
	ATSA	1.00	0.99	0.85	0.96	0.96

778

## 779 **5. Discussion and conclusions**

Masking clouds and cloud shadows is necessary for many applications of optical satellite
 images, because it is difficult to acquire totally cloud-free images in most places, particularly when

time series are needed to monitor change. Many methods have been developed to screen clouds 782 and cloud shadows automatically in optical images. However, they may not perform well in very 783 784 cloudy regions. Aiming to produce more accurate cloud and shadow masks of optical imagery in cloudy regions, an automatic time series analysis based method, ATSA, was developed in this study. 785 ATSA was tested in three sites with different dominant land covers. Landsat-8 OLI images, 786 787 Landsat-4 MSS images, and Sentinel-2 images were used to evaluate the performance of ATSA for screening clouds and cloud shadows in images with different band configurations and quality. 788 789 Results show that ATSA can obtain accurate cloud and shadow masks in all sites and all data sets 790 except the images with snow and ice cover. The comparison with an advanced algorithm, Fmask, also confirms that ATSA can yield robust and accurate cloud and shadow masks in cloudy regions. 791 The good performance of ATSA can be attributed to the following strengths. 792

First, ATSA only needs a minimum number of input bands. Only 5 bands, blue, green, red, 793 NIR, and SWIR bands, are required, and this requirement can be reduced to 3 bands if the images 794 795 do not have blue and SWIR bands. The low requirement of input bands brings two advantages. The first advantage is that in the regions tested, the results can be more robust than existing 796 methods when processing images with various conditions. Although the spectral similarities 797 798 among different land surfaces and clouds and cloud shadows are complex, being different among locations and times, a common characteristic is that they affect the pixel values from visual to near 799 800 infrared bands, i.e., clouds brighten these bands and shadows darken them. In general, adding more 801 bands into the screening process, such as thermal bands or a cirrus band, can improve the accuracy of cloud and shadow masks, especially for the single-image cloud detection methods (Foga et al., 802 803 2017; Zhu et al., 2015). However, it may also lead to more uncertainties and errors in some extreme 804 cases. For example, Fmask uses visible, near infrared, and thermal bands to identify clouds; it also

uses the cirrus band in Landsat-8 images to detect clouds (Zhu et al., 2015), while ATSA uses 805 neither the thermal nor cirrus bands. In the Hong Kong site, for the image DOY339, Fmask misses 806 most clouds. A further investigation of all bands of this image reveals that the thermal band is 807 cooler in a small sub-area of this mostly cloudy image (Fig. 25). As a result, Fmask only detects 808 clouds in this cold area and omits other warmer clouds. According to the USGS product guide, 809 810 Fmask has a known issue that either too large or too small temperature differentials will lead to errors in cloud detection. The second advantage of using fewer bands is that the algorithm is more 811 812 flexible and applicable than existing methods when processing images from different optical 813 sensors. For cloudy places, we expect ATSA to: (1) extend the history for automated Landsat time series analyses with cloud and cloud shadow masks that are highly accurate, but automatically 814 derived, back to the MSS era of the 1970s (instead of only the TM era of the 1980s); and (2) in the 815 era of Sentinal-2, allow for denser time series in intra-annual analyses such as those examining 816 vegetation phenology. The past and ongoing optical sensors have different configurations of 817 818 spectral bands. However, most of these optical sensors have visible and near infrared bands. ATSA can be applied to all images with these basic bands, which is very important and necessary when 819 we process historical satellite images with limited bands. 820



False color image of DOY339 in HK site

821







Band 10-Thermal

Fig. 25. False color composite of the Landsat-8 images DOY339 in the Hong Kong site (left), its cloud mask by Fmask (center), and the thermal band of this image (right)

825	Second, ATSA has fewer predefined parameters than most existing methods. In ATSA, there
826	are only two important predefined parameters, i.e., $A$ in Eq. (8) and $B$ in Eq. (12) tune the threshold
827	for identifying clouds and shadows in the time series respectively. As standard deviation
828	multipliers of variation through the time series, these two parameters regulate the degree of
829	strictness for masking clouds and shadows. In other words, they balance the omission errors and
830	commission errors of cloud and shadow detection. In our tests, $A$ and $B$ are 0.5 and 1.5 for the
831	Hong Kong site, 1.0 and 1.5 for the Puerto Rico site, and 1.2 and 2.0 for the Beijing site. Fig. 26
832	shows the cloud user's and producer's accuracy for the Landsat-4 MSS image DOY200 in Fig. 21
833	when using different values of parameter $A$ within the recommended range 0.5-1.5. Larger values
834	of parameter $A$ improve the user's accuracy but meanwhile decrease the producer's accuracy. The
835	parameter B shows a similar effect on the accuracy of cloud shadow detection. Fig. 26 also suggests
836	that the detection accuracy is not very sensitive to the parameter. There is a wide range of parameter
837	A able to obtain both producer's accuracy and user's accuracy higher than 0.95. Users can tune
838	these two parameters according to their specific applications. For example, studies using time
839	series to model land surface parameters, such as forest biomass and crop yield, are very sensitive
840	to clouds, even the extremely thin clouds. These studies may hope to mask out all possible clouds
841	and accept some commission errors, so smaller values of parameter $A$ and $B$ should be used. In
842	addition, ATSA also use the statistics of each image in the time series to determine some parameters
843	to increase the adaptability of ATSA. For example, the HOT transformation has been used in many
844	cloud screening methods, such as MFC (Li et al., 2017) and Fmask (Zhu and Woodcock, 2012).
845	However, these methods apply one HOT formula to all images. For instance, both MFC and Fmask
846	use $HOT=B_{blue}$ -0.5B <sub>red</sub> for all images. However, the coefficients in the HOT transformation vary

from scene-to-scene, so it is necessary to estimate the HOT parameter for individual images (Chen et al., 2016; Zhang et al., 2002). ATSA regresses the coefficients in HOT transformation model in each image by an automated strategy which can get optimal cloud index images.



850

Fig. 26. User's and producer's accuracy of cloud detection for the Landsat-4 MSS image
 DOY200 in the Puerto Rico site using different values of parameter A

853

854 Third, ATSA uses the minimal clear observations in image time series over cloudy regions to ensure accurate cloud and shadow masking without fitting a time series model of these 855 observations. For both cloud and shadow detection, there are two hierarchies in ATSA. In the first 856 hierarchy, ATSA selects samples from all images in the time series for optimizing the class centers 857 in the K-means classifier. As we know, it is quite common that image scenes are totally covered 858 by clouds. If the K-means classifier (K=2 or 3) is applied to each individual image, it cannot detect 859 all clouds in a totally cloud-covered image. In the second hierarchy, ATSA only uses "clear" 860 861 observations in the time series to estimate the adaptive threshold, and further detect clouds and 862 shadows omitted in the first hierarchy. Another multi-temporal method, Tmask, also uses clear observations in the time series to refine the initial cloud mask from Fmask. It can detect more thin 863 864 clouds than Fmask (Zhu and Woodcock, 2014). However, Tmask is not appropriate in our test sites

in cloudy regions. Fig. 27 shows the number of clear observations of individual pixels in the time-865 series data of Hong Kong and Puerto Rico site. We can see that both sites have considerable pixels 866 with fewer than 6 clear observations. The clear observations are not enough for Tmask (15 clear 867 observations are recommended) to accurately estimate the parameters in the time series model, 868 leading to misclassifying cloudy pixels as clear pixels (Foga et al., 2017). Unlike Tmask, ATSA 869 870 does not fit a time series model using many clear observations. It can be an alternative to Tmask for screening clouds in time-series data of cloudy regions or short time series (e.g., one year) which 871 872 is unlikely to have enough clear observations.



Fig. 27. Number of clear observations of individual pixel in the Landsat-8 time series at both

- 875 Hong Kong and Puerto Rico sites.
- 876

There are also some limitations in ATSA. First, ATSA currently does not have a snow detection module. In tropical and subtropical regions, which are among the cloudiest regions (Ju and Roy, 2008), images in these regions do not have snow in all seasons except at the highest elevations. If the images include snow, ATSA is likely to detect snow as clouds (see Jan.26 image in Fig. 24). This outcome may be acceptable in many applications, such as vegetation studies, in which, like clouds, snow would often be excluded. Actually, most current algorithms often confuse

snow and clouds even if they have a snow detection module, like the Fmask results shown in Fig. 883 13 and Fig. 24. If more powerful snow detection methods are developed in the future, they can be 884 integrated with ATSA. Second, although the HOT transformation can suppress the pixel values of 885 various land covers (also see an experiment in a desert landscape shown in the Supplementary 886 Data), the very bright pixels may be identified as clouds. A recent study proposed an iterative HOT 887 888 (IHOT) algorithm which can better suppress surface reflectance (Chen et al., 2016), but it needs more computing time. IHOT can be used as an alternative to HOT if the computing time is not a 889 890 restriction factor. Third, land cover changes may happen in the time series. It may bring temporal 891 variability in the HOT time series which could further affect the cloud detection by ATSA. An experiment reported in the Supplementary Data shows that ATSA may be not affected by many 892 types of land cover changes, but other methods (e.g. Tmask) which can model land cover change 893 may obtain better results than ATSA when substantial land cover changes exist. Fourth, ATSA may 894 omit some cloud shadows on water surfaces or cloud shadows on the land surface that are 895 896 extremely thin. Omission of cloud shadows on water surfaces may not affect mapping the water bodies, but it may affect water quality modeling. Thin cloud shadow on land surfaces may also 897 affect quantitative information retrieval. Omission errors from missed cloudy pixels are the most 898 899 common errors in cloud shadow masking methods (Foga et al., 2017); however, more accurate cloud detection with ATSA in the types of landscapes tested should reduce this error. These errors 900 can be corrected by a further manual checking. Fifth, ATSA requires a time series, albeit with fewer 901 902 dates than existing methods. Last, due to the limitation of resources and support, ATSA was tested in several typical sites and on data sets from three satellite sensors. More comparison and 903 904 validation are needed, and they are our future studies. Due to its simple principles, ATSA has an 905 acceptable efficiency for processing time-series data. ATSA only used 11 minutes and 13 minutes

for the Landsat-8 time series in Hong Kong and Puerto Rico sites respectively (program coded in
interactive data language and run on a windows laptop with a 2.50GHz CPU and 8 GB RAM). We
welcome other researchers to test ATSA in more areas and different data sets. The code of ATSA
is available upon request.

In conclusion, a new cloud and cloud shadow screening method, ATSA, was developed in this study. Its target is time series optical images in cloudy regions. ATSA is a valuable supplement to the family of cloud and cloud shadow masking algorithms. It will support studies of land surface dynamics using dense optical time series, such as studies of forest phenology in tropical regions using Landsat or Sentinel-2 images.

915

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926

#### 927 Appendix A. Supplementary data

928 Supplementary data to this article can be found online at ##.

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