

# **An elastoplastic model for gap-graded soils based on homogenization theory**

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## 24 **Abstract**

25 Naturally formed soils (e.g., residual soils and deposit clays) usually show an absent range of  
26 particle size. Frequently used by geotechnical communities worldwide, such gap-graded soils  
27 can be simplified as binary mixtures composed of fine soil matrix and coarse rock aggregates.  
28 In this study, an elastoplastic model is proposed for gap-graded soils based on a volume  
29 average scheme and homogenization theory. The proposed model incorporates a structural  
30 variable to account for the evolution of the inter-granular skeleton of rock aggregates. The  
31 model is then implemented in a numerical code by the linearized integration technique  
32 proposed by Bardet and Choucair (1991). It is shown that the model can predict a wide range  
33 of variations of the overall shear responses with the increase in volume fraction of rock  
34 aggregates. An isotropic loading induces a nonuniform stress distribution in gap-graded soils,  
35 where the stress in the soil matrix is lower than that of the rock aggregates. The stress path of  
36 the matrix is approximately parallel with that of the rock aggregates during triaxial shear  
37 loading. The proposed model contains only one additional structure parameter compared with  
38 the generalized modified Cam clay model, which can be easily calibrated from the data of a  
39 conventional triaxial compression tests. Comparison between our model predictions and the  
40 experimental data from literature indicates that the propose model can well reproduce the  
41 mechanical responses of gap-graded soils within a wide range fraction of rock aggregates.

42

43 **Keywords:** Elastoplastic model; Gap-graded soils; Volume average scheme; Homogenization  
44 theory

## 45 1. Introduction

46 Natural soils are usually composed of fine-grained soil and rock aggregates with a great range  
47 gap of size distribution (Yang and Juo, 2001; Zhao *et al.*, 2007; Ueda *et al.*, 2011; Change *et*  
48 *al.*, 2014; Ng *et al.*, 2016; Deng *et al.*, 2017; Cui *et al.*, 2017; Xu and Coop, 2017; Yang *et al.*,  
49 2018). Such gap-graded soils have frequently been treated by mixture theory as binary  
50 mixtures consisting of soft soil matrix and stiff rock inclusions (Vallejo, 2000; Peters and  
51 Berney, 2010; Zhou *et al.*, 2016; Shi and Yin, 2017). In arid and semi-arid areas, the fine-  
52 grained soils originate from the disintegration of parent rocks (from surface inwards) due to  
53 weathering, e.g., wetting-drying cycles or temperature oscillations. Since the disintegration is  
54 mainly a physical process, the soil matrix has the same mineral composition as the rock  
55 blocks. The soil matrix may also be a resultant of erosion and transportation of sedimentary  
56 soils from other places, followed by subsequent deposition (Chandler, 2000). However, this  
57 may probably happen in wet areas. The binary gap-graded soils are used as geotechnical  
58 structures worldwide, such as riprap, dam and high-fill subgrade (Zhao *et al.*, 2007; Vallejo,  
59 2000; Chen and Cui, 2017). The rock fraction has a considerable influence on the workability  
60 of these geo-structures (Vallejo, 2000; Peters and Berney, 2010; Zhou *et al.*, 2016).

61 The mechanical behavior of gap-graded soils was documented by many previous  
62 researchers, including laboratory work (Graham *et al.*, 1989; Kumar, 1996; Yin, 1999;  
63 Vallejo, 2000; Monkul *et al.*, 2005; Monkul and Ozden, 2007; Ueda *et al.*, 2010; Shi and Yin,  
64 2018; Shi *et al.*, 2018) and numerical simulations (González *et al.*, 2004; Dai, 2015; Ng *et al.*,  
65 2016; Zhou *et al.*, 2016; Shi and Herle, 2017; Wu *et al.*, 2017). Most of these works focus on  
66 the qualitative analysis of the test data, with relatively scarce theoretical work on the gap-  
67 graded soils reported. To this end, a model incorporating the coarse fraction effect is proposed  
68 here for gap-graded soils, which is further validated by reported experimental data.

## 69 **2. Structure of gap-graded soils**

### 70 *Structure evolution with increasing volume fraction of rock aggregates*

71 For a small volume fraction of rock inclusions, the coarse aggregates may suspend in the  
72 matrix, and the overall mechanical behavior depends on the soil matrix and the interaction  
73 between the matrix and inclusion phases. With increasing volume fraction of the aggregates,  
74 contacts between the aggregates gradually form. However, these may only be partial contacts.  
75 The inter-granular skeleton of aggregates can support a higher stress than the matrix due to  
76 the partial contacts and soil bridges between the aggregates (Jafari and Shafiee, 2004; Fei,  
77 2016; Shi and Yin, 2017). When the volume fraction of soil matrix approximates the  
78 maximum porosity of the inclusions (loosest packing state of coarse aggregates), a continuous  
79 inter-granular skeleton forms. The corresponding fraction of the matrix is noted as ‘transition  
80 fines content’ (Monkul and Ozden, 2007). In this study, we only consider the fine content  
81 beyond the ‘transition fines content’, which corresponds to soils in intense weathering areas.  
82 An extremely small fine content (e.g., close or equal to zero) may result in macro-pores  
83 between the rock aggregates, which is beyond the scope that the mixture theory can treat.

### 84 *Volume fraction of rock aggregates*

85 Due to its dual-level configuration, it is challenging to provide an exact description of the  
86 structure of a gap-graded soil. To this end, the volume fraction concept is introduced in the  
87 subsequent analysis (see, e.g., de Boer and Ehlers, 1986; Didwania and de Boer, 1999). This  
88 concept leads to a substitute (smeared) continua with reduced physical quantities of the  
89 constituents, which can be easily incorporated into the mixture theory.

90 A gap-graded soil is simplified as a mixture of matrix and inclusions, with the matrix  
91 being the soft soil and the inclusions being the rock aggregates. As mentioned above, the  
92 structure transition of gap-graded soils is controlled by the volume fraction of the rock

93 aggregates  $\phi_a$ , provided that the rock aggregates are randomly arranged in the soil matrix.  
 94 Therefore, the coarse volume fraction  $\phi_a$  can be introduced as a bridge between the overall  
 95 behavior of the mixture and that of the soft soil matrix. The compressibility of the matrix is  
 96 much higher than that of the rock inclusions, thus the volume fraction of rock aggregates  
 97 increases with increasing compression loading. The volume fraction of rock aggregates is  
 98 hence a state dependent variable. For a given coarse mass fraction, it can be formulated as a  
 99 function of the overall void ratio  $e$  and the void ratio of the soil matrix  $e_m$ :

$$100 \quad \phi_a = \frac{e_m - e}{(1 + e)e_m} \quad (1)$$

101 The void ratio of the soil matrix  $e_m$  is given as

$$102 \quad e_m = \frac{\psi_a \rho_m + (1 - \psi_a) \rho_a}{(1 - \psi_a) \rho_a} e \quad (2)$$

103 where  $\psi_a$  is the dry mass fraction of rock aggregates;  $\rho_m$  and  $\rho_a$  are the particle densities of  
 104 the soil matrix and the rock aggregates, respectively. Two fractions of aggregates are used in  
 105 this work: the volume fraction  $\phi_a$  and the dry mass fraction  $\psi_a$ .  $\phi_a$  is used for homogenizing  
 106 state variables of binary gap-graded soils in the sequel analysis. The dry mass fraction  $\psi_a$  is  
 107 commonly adopted in laboratory tests, since it is constant during compression and shearing  
 108 process. In numerical simulations, the volume fraction of aggregates depends on the stress  
 109 state which is computed from the overall void ratio and the dry mass fraction of aggregates  
 110 (Eqs. (1) and (2)). Note that this is not applicable for a mixture with very high coarse fractions,  
 111 in which the macro-pores may exist between the rock aggregates.

## 112 *Volume average stresses and strains*

113 Due to the difference of stiffness between the two phases of a mixture, the interaction at the  
 114 interface may result in a nonuniform stress (strain) field. As the essential load-carrying

115 members of the mixture, the hard rock aggregates sustain a higher loading than that of the  
 116 ductile matrix, and the loading increases with the volume fraction of the aggregates (Tandon  
 117 and Weng, 1988). Correspondingly, with increasing volume fraction of the rock aggregate,  
 118 the strain experienced by the matrix phase decreases, and the magnitude of stress in the matrix  
 119 drops.

120 In the following, the focus will be placed on modeling the mechanical behavior in the  
 121 frame of continuum mechanics rather than describing the microstructure of the mixture media.  
 122 As suggested by Tandon and Weng (1988), the mean-field theory provides a reasonable  
 123 approximation for describing the behavior of geomaterials. Using the volume fraction concept,  
 124 all physical and geometric quantities can be defined in a predefined space (e.g., deformation,  
 125 motion, and stress invariants). In the sequel, the stress and strain variables are approximated  
 126 by the statistical average values of the real ones (de Boer, 2006). It is of convenience to use  
 127 two subscripts, ‘a’ and ‘m’, to denote quantities pertaining to the rock aggregates and soil  
 128 matrix, respectively. Following the volume average scheme, the overall stress tensor  $\sigma'_{ij}$  and  
 129 overall strain tensor  $\varepsilon_{ij}$  can be expressed as

$$130 \quad \sigma'_{ij} = \frac{1}{V_t} \int_{V_t} \tilde{\sigma}'_{ij}(\mathbf{x}) dV = \frac{1}{V_t} \int_{V_a} \tilde{\sigma}'_{ij}(\mathbf{x}) dV + \frac{1}{V_t} \int_{V_m} \tilde{\sigma}'_{ij}(\mathbf{x}) dV = \phi_a \sigma'_{ij,a} + (1 - \phi_a) \sigma'_{ij,m} \quad (3a)$$

$$131 \quad \varepsilon_{ij} = \frac{1}{V_t} \int_{V_t} \tilde{\varepsilon}_{ij}(\mathbf{x}) dV = \frac{1}{V_t} \int_{V_a} \tilde{\varepsilon}_{ij}(\mathbf{x}) dV + \frac{1}{V_t} \int_{V_m} \tilde{\varepsilon}_{ij}(\mathbf{x}) dV = \phi_a \varepsilon_{ij,a} + (1 - \phi_a) \varepsilon_{ij,m} \quad (3b)$$

132 where  $V_t$  is the representative elementary volume (REV) of gap-graded soils,  $\tilde{\sigma}'_{ij}(\mathbf{x})$  and  $\tilde{\varepsilon}_{ij}(\mathbf{x})$   
 133 are local stress and strain over the defined REV.  $V_a$  and  $V_m$  are the volumes of the rock  
 134 aggregates and soil matrix, respectively.  $\sigma'_{ij,a}$ ,  $\sigma'_{ij,m}$ ,  $\varepsilon_{ij,a}$  and  $\varepsilon_{ij,m}$  are the stress and strain  
 135 variables of the two constituents. Note that the stiffness of rock aggregates is extremely high,  
 136 thus, a negligible deformation can be expected within the conventional stress range, i.e.,

137  $\varepsilon_{ij,a} \approx 0$ .

138 The constitutive relationship of the gap-graded soils depends on the following factors: (1)  
139 the stress-strain relationships for the two phases. The rock aggregates are extremely hard  
140 with negligible deformation, and the soil matrix shows a plastic deformation when subjected  
141 to an external loading. (2) The homogenization approaches which builds a bridge between  
142 the overall compliance (stiffness) and the respective ones of the two phases. These two  
143 factors will be addressed in the following two sections.

144

### 145 **3. Modeling the soil matrix**

146 Natural soil-rock mixtures usually contain a fraction of soil matrix higher than the ‘*transition*  
147 *fines content*’. In this case, the overall behavior of the mixtures depends on that of the soil  
148 matrix, partial contacts between the coarse aggregates and the interaction at the interface  
149 between the matrix and aggregates. In the absence of a continuous inter-granular skeleton, the  
150 mechanical behavior of the soil matrix provides a frame of reference for assessing the overall  
151 behavior of the gap-graded soil. It is assumed that the soil matrix follows an incremental  
152 stress-strain relationship. A numerical scheme based on the tangent homogenization is  
153 adopted to compute the overall compliance of the soil matrix (Ju and Sun, 2001).

#### 154 *Elastic deformation*

155 Following the convention of classical soil mechanics, compressive stress and strain are  
156 taken as positive. An incremental elasto-plastic description is adopted for the ductile soil  
157 matrix. The incremental strains of the soil matrix  $\varepsilon_{ij,m}$  is decomposed into an elastic part  $\varepsilon_{ij,m}^e$   
158 and a plastic part  $\varepsilon_{ij,m}^p$ :

$$159 \quad d\varepsilon_{ij,m} = d\varepsilon_{ij,m}^e + d\varepsilon_{ij,m}^p \quad (4)$$

Logarithmic volumetric strain is adopted in this study. It is assumed that the logarithmic value of the specific volume  $v_m$  changes linearly with the effective mean stress  $p'_m$  of the matrix (Butterfield, 1979) for both virgin compression and swelling curves. Following this assumption, the elastic incremental stress-strain relationship can be expressed as

$$d\varepsilon_{ij,m}^e = \frac{1}{K_{e,m}} \left[ \frac{(1+\mu_m)}{3(1-2\mu_m)} \delta_{ik} \delta_{jl} - \frac{\mu_m}{3(1-2\mu_m)} \delta_{ij} \delta_{kl} \right] d\sigma'_{kl,m} \quad (5)$$

where  $K_{e,m} = \frac{p'_m}{\kappa_m}$  is the elastic modulus,  $\kappa_m$  is the slope of the swelling line of the matrix in double logarithmic  $\ln v_m : \ln p'_m$  relationship,  $\mu_m$  is the Poisson's ratio of the soil matrix,  $\delta_{ij}$ ,  $\delta_{ik}$ ,  $\delta_{jl}$  and  $\delta_{kl}$  are Kronecker's symbols.

#### Plastic flow

The fabric of a natural gap-graded soil depends on its history of formation, and the soil may be anisotropic due to a preferred orientation of the rock aggregates during erosion, depositional and post-depositional processes (Zhou *et al.*, 2017). As a preliminary investigation, only the isotropic case is considered in this study. It is widely accepted that the critical state type models (Roscoe and Burland, 1968; McDowell and Hau, 2004; Yao *et al.*, 2004, 2012; Gao and Zhao, 2012, 2015, 2017; Zhao and Gao, 2016) can well reproduce the stress-strain relationship of reconstituted soils. A generalized form of the Modified Cam clay model proposed by McDowell and Hau (2004) is adopted. The yield surface for the soil matrix  $f_m$  is given as

$$f_m : q_m^2 + \frac{M_m^2}{1-k_m} \left( \frac{p'_m}{p'_c} \right)^{\frac{2}{k_m}} p_c'^2 - \frac{M_m^2 p_c'^2}{1-k_m} = 0; \quad (k_m \neq 1) \quad (6)$$

where  $q_m$  is the deviatoric stress of the matrix,  $p'_c$  represents the size of the yield surface,  $M_m$  is a strength parameter corresponding to a unique critical state line in  $p'_m : q$  stress plane,



181 and  $k_m$  controls the shape of the yield surface. Note that the Critical State Line (CSL) in the  
 182 compression plane changes with the shape parameter  $k_m$ .

183 In the sequel, the stress-strain relationship of the soil matrix will be presented following  
 184 the incremental plasticity theory presented by Scott (1985) which has been adopted in plastic  
 185 fractional order plasticity (Sun and Shen, 2017; Sun *et al.*, 2018). The size of the yield surface  
 186  $p'_c$  acts as a hardening variable. Consistency condition of the yield surface gives

$$187 \quad \frac{\partial f_m}{\partial \sigma'_{kl}} d\sigma'_{kl} + \frac{\partial f_m}{\partial p'_c} dp'_c = 0 \quad (7)$$

188 In many critical state models, the evolution of hardening variable  $p'_c$  is assumed as a  
 189 function of the plastic volumetric strain increment  $d\varepsilon_{v,m}^p$  (e.g., Yao *et al.*, 2009; Yao and  
 190 Zhou, 2013; Hong *et al.*, 2014):

$$191 \quad dp'_c = \frac{dp'_c}{d\varepsilon_{v,m}^p} d\varepsilon_{v,m}^p = \frac{p'_c}{\lambda_m - \kappa_m} d\varepsilon_{v,m}^p \quad (8)$$

192 where  $\lambda_m$  is the slope of the Normal Compression Line (NCL) in double logarithmic  
 193  $\ln v_m : \ln p'_m$  plot. Note that a linear relationship between  $\ln v_m$  and  $\ln p'_m$  is assumed for the  
 194 virgin compression of the soil matrix here.

195 The plastic (volumetric) strain increment ( $d\varepsilon_{v,m}^p$  or  $d\varepsilon_{ij,m}^p$ ) of the soil matrix is related to  
 196 the maximum gradient of the plastic potential surface  $g_m$ :

$$197 \quad d\varepsilon_{v,m}^p = d\zeta_m \frac{\partial g_m}{\partial p'_m} \quad (9a)$$

$$198 \quad d\varepsilon_{ij,m}^p = d\zeta_m \frac{\partial g_m}{\partial \sigma'_{ij,m}} \quad (9b)$$

199 where  $d\zeta_m$  is a positive plastic multiplier. Substitution of Eqs. (8) and (9a) into Eq. (7) gives

$$d\zeta_m = -\frac{1}{K_{p,m}} \left( \frac{\partial g_m}{\partial \sigma'_{ij,m}} \right)^{-1} m_{ij,m} n_{kl,m} d\sigma'_{kl,m} \quad (10)$$

where  $K_{p,m}$  is the plastic modulus of the soil matrix, the unit vectors  $m_{ij,m}$  and  $n_{kl,m}$  represent the normal to the potential surface and yield surface of the matrix, respectively:

$$K_{p,m} = \frac{p'_c}{\lambda_m - \kappa_m} \frac{\frac{\partial f_m}{\partial p'_c} \frac{\partial g_m}{\partial p'_m}}{\left\| \frac{\partial f_m}{\partial \sigma'_{kl,m}} \right\| \left\| \frac{\partial g_m}{\partial \sigma'_{ij,m}} \right\|}; \quad n_{kl,m} = \frac{\frac{\partial f_m}{\partial \sigma'_{kl,m}}}{\left\| \frac{\partial f_m}{\partial \sigma'_{kl,m}} \right\|}; \quad m_{ij,m} = \frac{\frac{\partial g_m}{\partial \sigma'_{ij,m}}}{\left\| \frac{\partial g_m}{\partial \sigma'_{ij,m}} \right\|} \quad (11)$$

where  $\|x_{ij}\| = \sqrt{x_{ij}x_{ij}}$ . Substitution of Eqs. (5), (9b) and (10) into Eq. (4) gives

$$d\varepsilon_{ij,m} = C_{ijkl,m} d\sigma'_{kl,m} \quad (12a)$$

$$C_{ijkl,m} = \frac{1}{K_{e,m}} \left[ \frac{(1 + \mu_m)}{3(1 - 2\mu_m)} \delta_{ik} \delta_{jl} - \frac{\mu_m}{3(1 - 2\mu_m)} \delta_{ij} \delta_{kl} \right] - \frac{1}{K_{p,m}} m_{ij,m} n_{kl,m} \quad (12b)$$

#### 4. A new homogenization approach for gap-graded soils

The majority of early homogenization studies have been developed based on linear elasticity consideration of the constituents (Eshelby, 1961; Hill, 1965; Mori and Tanaka, 1973; Lielens *et al.*, 1998). However, a gap-graded soil is not typical soils that can be described in the classical mixture theory for three-fold reasons: (1) the soil matrix in a gap-graded soil is dominantly plastic, (2) the modulus of the rock aggregates is normally much larger compared with that of the matrix, and (3) the interface between the constituents is not perfect. To this end, a new homogenization model is proposed based on mixture theory for gap-graded soils.

##### *Effective compliance tensor*

The microstructure of natural gap-graded soils can be defined by selecting a suitable REV with randomly distributed rock aggregates. The work by Tu *et al.* (2005) on mixtures with

219 this kind of microstructure reveals that increasing the modulus of aggregates significantly  
 220 improve the overall modulus for small modulus ratios (the ratio of the modulus of the  
 221 aggregates to that of the matrix). Further increase of the modulus ratio brings the overall  
 222 modulus to a limit state (named as ‘saturation state’ by Tu et al., 2005). Therefore, it is  
 223 reasonable to relate the overall modulus of a gap-graded soil to the one of the soil matrix and  
 224 the inter-granular skeleton regardless of the modulus of the rock-aggregates.

225 The overall modulus of a gap-graded soil increases with the volume fraction of rock  
 226 aggregates, and it should meet the following two requirements: (1) for a gap-graded soil with  
 227 negligible fraction of aggregates, e.g.,  $\phi_s \approx 0$ , the overall elastic modulus  $K_e$  and plastic  
 228 modulus  $K_p$  are assumed to be approximately close to the corresponding ones of the soil  
 229 matrix, i.e.,  $K_e \approx K_{e,m}$  and  $K_p \approx K_{p,m}$ . (2) when the inter-granular void ratio (the ratio of the  
 230 volume of inter-granular space to that of the aggregates) approaches the minimum void ratio  
 231 of the rock aggregates, additional external load is mainly sustained by the inter-granular  
 232 structure. Hence, the overall modulus should be much larger than that of the soil matrix.

233 The homogenization model proposed by Shi and Yin (2017) for the compression behavior  
 234 of sand-marine clay mixtures is modified for the gap-graded soils:

$$235 \quad \ln K_e = \chi \ln K_{e,m} - \ln(1 - \phi_a) \quad (13a)$$

$$236 \quad \ln K_p = \chi \ln K_{p,m} - \ln(1 - \phi_a) \quad (13b)$$

237 where  $\chi$  is a structure variable representing the inter-granular structure evolution given by

$$238 \quad \chi = \left( \frac{\phi_a^0}{\phi_a^0 - \phi_a} \right)^\xi \quad (14)$$

239 where  $\xi$  is a structure parameter controlling the sensitivity of the structure variable on the  
 240 volume fraction of rock aggregate,  $\phi_a^0$  is the maximum volume fraction of aggregates for  
 241 pure coarse inclusions, and it corresponds to the minimum void ratio of the rock aggregates

242  $e_{\min}$  :

$$243 \quad \phi_a = \frac{1}{1+e_{\min}} \quad (15)$$

244 Analogous to the stress-strain relationship of the soil matrix, the overall one can be  
245 expressed as

$$246 \quad d\varepsilon_{ij} = d\varepsilon_{ij}^e + d\varepsilon_{ij}^p = C_{ijkl} d\sigma'_{kl} \quad (16a)$$

$$247 \quad C_{ijkl} = \frac{1}{K_e} \left[ \frac{(1+\mu_m)}{3(1-2\mu_m)} \delta_{ik} \delta_{jl} - \frac{\mu_m}{3(1-2\mu_m)} \delta_{ij} \delta_{kl} \right] - \frac{1}{K_p} m_{ij} n_{kl} \quad (16b)$$

248 where  $m_{ij}$  and  $n_{kl}$  are unit vectors representing the normal to the potential surface and yield  
249 surface of the gap-graded soils, respectively:

$$250 \quad n_{kl} = \frac{\frac{\partial f}{\partial \sigma'_{kl}}}{\left\| \frac{\partial f}{\partial \sigma'_{kl}} \right\|}; \quad m_{ij} = \frac{\frac{\partial g}{\partial \sigma'_{ij}}}{\left\| \frac{\partial g}{\partial \sigma'_{ij}} \right\|} \quad (17)$$

#### 251 *Stress concentration tensor*

252 The overall compliance tensor of the gap-graded soil can be expressed as a function of the  
253 soil matrix using an incremental stress (or strain) concentration tensor. By applying the  
254 volume average scheme, the stress concentration tensor is defined as

$$255 \quad d\sigma'_{ij,m} = \Theta_{ijkl} d\sigma'_{kl} \quad (18)$$

256 Considering that the deformation of the rock aggregates is negligible, the overall  
257 incremental strain of the gap-graded soils is given as

$$258 \quad d\varepsilon_{ij} = (1-\phi_a) d\varepsilon_{ij,m} \quad (19)$$

259 Combination of Eqs. (12a), (16a), (18) and (19) leads to the following stress concentration  
260 tensor

$$\Theta_{ijkl} = \frac{1}{1-\phi_a} C_{pqij,m}^{-1} C_{pqkl} \quad (20)$$

## 262 *Simplification of the full constitutive model*

263 To reproduce the stress-strain curves of the gap-graded soils, one must identify the yield  
 264 surface  $f$  and the plastic potential surface  $g$  in order to determine the yield direction vector  
 265  $n_{kl}$  and the plastic flow direction vector  $m_{ij}$ . In the sequel, the unit flow and loading vectors  
 266 of the gap-graded soil are derived by assuming an associated flow rule.

267 Based on a similar form of the incremental total strain (Eq. (19)), the overall incremental  
 268 plastic strain of the gap-graded soils is related to that of the soil matrix:

$$d\varepsilon_{ij}^p = (1-\phi_a) d\varepsilon_{ij,m}^p \quad (21)$$

270 The overall plastic strain increment of the gap-graded soils is proportional to the  
 271 maximum gradient of the corresponding plastic potential surface  $g$  :

$$d\varepsilon_{ij}^p = d\zeta \frac{\partial g}{\partial \sigma'_{ij}} \quad (22)$$

273 Substitution of Eqs. (9b) and (22) into (21) yields the following equation:

$$\frac{\partial g}{\partial \sigma'_{ij}} = \frac{d\zeta_m}{(1-\phi_a) d\zeta} \frac{\partial g_m}{\partial \sigma'_{ij,m}} \quad (23)$$

275 where  $d\zeta_m$ ,  $d\zeta$ , and  $\phi_a$  are scalars. In consideration of the definitions of yield direction and  
 276 flow direction vectors in Eqs. (11) and (17), it follows that

$$m_{ij} = m_{ij,m} \quad (24)$$

278 An associated flow rule is assumed for the gap-graded soils, i.e., the yield surface  $f$  is the  
 279 same as the plastic potential surface  $g$ , so that

$$n_{ij} = m_{ij} \quad (25)$$

281 The compliance tensor of the gap-graded soils is represented by the following equation:

$$C_{ijkl} = \frac{1}{3K_e(1-2\mu_m)} \left[ (1+\mu_m)\delta_{ik}\delta_{jl} - \mu_m\delta_{ij}\delta_{kl} \right] - \frac{1}{K_p} m_{ij,m} m_{kl,m} \quad (26)$$

## 5. Model parameter calibration and numerical simulations

### Calibration of model parameters

The proposed elastoplastic model in the Section 4 contains seven parameters:  $M_m$ ,  $N_m$ ,  $\lambda_m$ ,  $\kappa_m$ ,  $\mu_m$ ,  $k_m$ ,  $\xi$ . Six of them are for the constitutive model of the soil matrix, denoted by a subscript ‘ $m$ ’.  $M_m$  is a strength parameter of the soil matrix, which can be calibrated from the critical state data in  $p'_m : q_m$  stress plane;  $N_m$  and  $\lambda_m$  describe the Normal Compression Line of the soil matrix in double logarithmic  $\ln v_m : \ln p'_m$  plot;  $\kappa_m$  corresponds to the slope of the swelling line of the clay matrix in  $\ln v_m : \ln p'_m$  compression plane;  $\mu_m$  is Poisson’s ratio of the soil matrix, which can be determined from the initial stiffness in triaxial compression test;  $k_m$  is a shape parameter controlling the shape of the yield surface, which can be calibrated from the critical state line in  $\ln v_m : \ln p'_m$  compression plane;  $\xi$  is a structure parameter describing the evolution of inter-granular skeleton with increasing volume fraction of rock aggregates.

A minimum of three conventional tests are required for the calibration of the seven model parameters: an oedometer test or isotropic compression test on the pure soil matrix, and triaxial shear tests on both the pure soil matrix and a gap-graded soil with a predefined mass fraction of the rock aggregates.  $N_m$ ,  $\lambda_m$  and  $\kappa_m$  can be determined from the loading-reloading curves of an oedometer (isotropic compression test) of the soil matrix.  $M_m$ ,  $\mu_m$  and  $k_m$  can be calibrated from a triaxial shear test on the pure soil matrix. The structure parameter  $\xi$  is calibrated by trial and error using the data of a triaxial shear test on a gap-

303 graded soil.

#### 304 *Stress integration of the constitutive model*

305 The model presented above is a rate-type stress-strain relationship, which can be solved by  
306 the linearized integration technique proposed by Bardet and Choucair (1991). The explicit  
307 integration scheme is utilized in this work to describe the material point response of various  
308 rock-soil mixtures from literature. It can be readily implemented into finite element codes for  
309 boundary value problems, which reduces the difficulties arising from the high non-linearity of  
310 the mechanical behavior of mixture soils. The overall loading constraints for the gap-graded  
311 soils in laboratory testing conditions can be linearized into the following equation (Bardet and  
312 Choucair, 1991):

$$313 \quad P_{ijk} d\sigma_{jk} + Q_{ijk} d\varepsilon_{jk} = dY_i \quad (27)$$

314 where  $P_{ijk}$  and  $Q_{ijk}$  are constant coefficients,  $dY_i$  is a loading increment during a loading  
315 process. In consideration of the overall stress-strain relationship (Eq. 16), Eq. (27) becomes

$$316 \quad (P_{ijk} + Q_{ipq} C_{pqjk}) d\sigma'_{jk} = dY_i \quad (28)$$

317 It is more convenient to use the stress increment of the soil matrix  $d\sigma'_{jk,m}$  as the principal  
318 invariants for a boundary value problem. Substitution of Eq. (18) into Eq. (28) gives

$$319 \quad (P_{ist} + Q_{ipq} C_{pqst}) \Theta_{stjk}^{-1} d\sigma'_{jk,m} = dY_i \quad (29)$$

320 The numerical stress integration procedure of the proposed model is outlined as follows:

321 (1) Suppose an initial isotropic overall stress state  $\sigma'_{jk(j \neq k)} = 0$  kPa,  $\sigma'_{jk(j=k)} = 10$  kPa, with a  
322 uniform stress distribution  $\sigma'_{kl,m} = \sigma'_{kl,a} = \sigma'_{kl}$ . Under this assumption, there may be a small  
323 deviation at the initial stage of loading, which is negligible when the stress level is  
324 significantly higher than 10 kPa.

325 (2) Determine the stress and strain constraint tensors  $P_{ijk}$  and  $Q_{ijk}$  based on the loading

326 conditions in laboratory testing. The test can be either stress controlled or strain controlled,  
 327 or (stress-strain) mixed controlled.

328 (3) For a given overall effective stress (or strain) increment at the current computation step  
 329  $dY_i$ , the stress increment of the soil matrix  $d\sigma'_{jk,m}$  is calculated.

330 (4) If the assumed stress state of the soil matrix  $\sigma'_{jk,m} + d\sigma'_{jk,m}$  is still within the yield surface  
 331 (Eq. (6):  $f_m(\sigma'_{jk,m} + d\sigma'_{jk,m}) \leq 0$ ), the stresses and strains are updated as

$$332 \quad \sigma'_{jk,m} \leftarrow \sigma'_{jk,m} + d\sigma'_{jk,m}; \quad \varepsilon_{jk,m} \leftarrow \varepsilon_{jk,m} + C_{jkrt,m} d\sigma'_{rt,m} \quad (30a)$$

$$333 \quad \sigma'_{pq} \leftarrow \sigma'_{pq} + \Theta_{pqjk}^{-1} d\sigma'_{jk,m}; \quad \varepsilon_{pq} \leftarrow \varepsilon_{pq} + (1 - \phi_s) C_{pqrt,m} d\sigma'_{rt,m} \quad (30b)$$

334 (5) If the assumed stress state is beyond the yield surface, the stress increment is reduced so  
 335 that the new stress increment  $\beta d\sigma'_{jk,m}$  pushes the stress state onto the yield surface, i.e.,

336  $f_m(\sigma'_{jk,m} + \beta d\sigma'_{jk,m}) = 0$ . So that the current stresses and strains are

$$337 \quad \sigma'_{jk,m} \leftarrow \sigma'_{jk,m} + \beta d\sigma'_{jk,m}; \quad \varepsilon_{jk,m} \leftarrow \varepsilon_{jk,m} + \beta C_{jkrt,m} d\sigma'_{rt,m} \quad (31a)$$

$$338 \quad \sigma'_{pq} \leftarrow \sigma'_{pq} + \beta \Theta_{pqjk}^{-1} d\sigma'_{jk,m}; \quad \varepsilon_{pq} \leftarrow \varepsilon_{pq} + (1 - \phi_s) \beta C_{pqrt,m} d\sigma'_{rt,m} \quad (31b)$$

339 (6) Update the following state variables: the compliance tensor of the soil matrix  $C_{jkrt,m}$ , the  
 340 structure variable  $\chi$ , the overall compliance tensor  $C_{jkrt}$ , the stress concentration tensor  
 341  $\Theta_{pqjk}^{-1}$ .

342 (7) Reset the loading increment  $dY_i \leftarrow (1 - \beta) dY_i$ , and compute the stress increment  $\sigma'_{jk,m}$  due  
 343 to the plastic deformation (based on Eq. (29)).

344 (8) The stress and strain tensor are computed by using Eq. (30), and the size of the yield  
 345 surface of the soil matrix  $p'_c$  is updated. Repeat steps (3)-(8) to proceed with the next  
 346 round of computation until the loading is completed.



### 347 *Simulations of the proposed model*

348 Following the numerical integration procedure presented above, simulations of drained  
349 triaxial tests of gap-graded soils using the proposed model is performed in this section. The  
350 calibrated model parameters for the proposed model are given in Table 1. The shape  
351 parameter  $k_m=2$  is assigned for the yield surface of the soil matrix, which is reduced to an  
352 ellipse adopted in the Modified Cam clay model. The maximum volume fraction of the pure  
353 rock aggregates is assumed to be 0.65, and the two phases (soil matrix and rock aggregates)  
354 have the same value of particle density:  $2650 \text{ kg/m}^3$ . An initial (isotropic) effective stress of  
355 10 kPa is assumed, and the initial state of the soil matrix is assumed on the Normal  
356 Compression Line. The sample is then isotropically compressed to 200 kPa, followed by a  
357 shear process.

358 The simulation results of the drained triaxial test of gap-graded soils are shown in Fig.1. A  
359 small value of the structure parameter is assumed  $\xi=0.05$  first, and four different rock mass  
360 fractions are considered (0.00, 0.10, 0.20, 0.40). It is not surprising that the sample with a  
361 higher rock fraction shows a smaller deformation, and an increase of the rock fraction  
362 improves the overall stiffness of the gap-graded soils remarkably. However, the effect of rock  
363 fraction on the overall shear strength is negligible. Note that the kinks in Fig. 1d is induced by  
364 the change of stress path (from isotropic compression to triaxial shear). To provide an insight  
365 into this phenomenon, the stress-strain curves of the soil matrix is shown in Fig. 2. It is seen  
366 that an isotropic loading leads to a non-uniform stress distribution in the gap-graded soils. The  
367 final stress (at the end of the isotropic loading process) in the soil matrix is smaller than the  
368 corresponding overall value. The stress paths of the soil matrix and the rock aggregates are  
369 almost parallel during the subsequent triaxial shear loading stage. The overall shear strength  
370 of normally consolidated mixtures is related to the critical stress state of the soil matrix and  
371 rock aggregates. A lower stress in the matrix in conjunction with a higher stress in the

aggregates results in a comparable shear strength to that of the pure soil matrix.

The overall behavior of gap-graded soils with a high value of structure parameter ( $\xi=0.35$ ) is presented in Fig. 3. It reveals a different coarse fraction effect from the one with a lower value of structure parameter ( $\xi=0.05$ , see Fig. 1). Both the initial stiffness and ultimate shear strength increase continuously with increasing rock fraction. This is consistent with the results of some gap-graded soils from literature (Jafari and Shafiee, 2004; Fei, 2016; Ruggeri et al., 2016). To further evaluate the performance of the proposed model, more simulations with different values of structure parameter are performed (where the dry mass fraction is assumed to be 0.40). The results are presented in Fig. 4. It indicates that overall shear strength increases as the structure parameter increases. The structure parameter is controlled by the particle shape and the particle size distribution of the coarse aggregates.

## 6. Validation of the proposed model

The shear strength of a gap-graded soil is affected by the volume fraction, the particle shape and particle size distribution of the rock aggregates (Jafari and Shafiee, 2004; Fei, 2016; Ruggeri et al., 2016). For some gap-graded soils the shear strength is insensitive to the volume fraction of rock aggregates until the rock particles form a continuous skeleton (Wood and Kumar, 2000). However, the shear strength of other gap-graded soils may increase continuously with increasing rock fraction (Jafari and Shafiee, 2004; Fei, 2016; Ruggeri et al., 2016). Benchmark analysis in the previous section reveals that the proposed model can simulate the shear strength behavior of the above two cases by assigning different values for the structure parameter. Three gap-graded soils from literature are used to validate the proposed model: (1) the natural gap-graded soils presented by Ruggeri et al. (2016); (2) the Kaolin clay-gravel mixtures from Jafari and Shafiee (2004); (3) the Kaolin clay-sand mixtures (data from Wood and Kumar, 2000).

397 *Natural gap-graded soils (Ruggeri et al., 2016)*

398 The poorly graded soil investigated by Ruggeri *et al.* (2016) consists of coarse grain particles  
399 and fine grey soil matrix, with a composition of 8% clay, 27% silts, 37% sand and 28% gravel.  
400 The shape of its PSD (particle size distribution) curve shows an absence of fine sand fraction.  
401 The pure soil matrix consists of 22% clay and 78% silts, and it has a liquid limit of 30% and a  
402 plastic limit of 18%. Three different mixtures were tested based on the PSD of the coarse  
403 aggregates: (1) HTP: the first series contains aggregates smaller than 16 mm, (2) HTP10: the  
404 grain size of the second series is smaller than 2.0 mm, and (3) HTP40: the third one has a  
405 grain size finer than 0.425 mm. The mixtures were prepared by mixing the soil matrix with  
406 the coarse aggregates in dry conditions. The reconstituted sample was first consolidated in a  
407 consolidometer at a vertical stress of 200 kPa to hold the sample together, then it was further  
408 consolidated at 400 kPa followed by triaxial shear under drained conditions. Four different  
409 cases of mass fractions of the coarse aggregates (10%, 20%, 30%, and 40%) are compared  
410 with the proposed model predictions.

411 *Kaolin clay-gravel mixtures (Jafari and Shafiee, 2004)*

412 Jafari and Shafiee (2004) have performed a series of triaxial tests on clay-gravel mixtures.  
413 The soil matrix is a commercial Kaolin clay. The particle density of the Kaolin material is  
414  $2740 \text{ kg/m}^3$ . The liquid limit and plastic limit are 69% and 31%, respectively. The gravel was  
415 retrieved from a riverbed. It consists of sub-rounded aggregates with a particle density of  
416  $2660 \text{ kg/m}^3$ . The size of gravel particles varies within a narrow range of 4.75 mm to 6.30 mm,  
417 with an average size of 5.55mm. The minimum void ratio of the gravel material was not given  
418 by the authors, and it was assumed as 0.41 following the summary of granulometric properties  
419 of granular materials by Herle and Gudehus (1999). Three initial volume fractions of the  
420 gravel aggregates were considered: 20%, 40%, and 60%. The gravel was first mixed with dry

421 Kaolin clay according to designated gravel fractions. The specimens were then compacted  
422 layer by layer (ASTM1999: standard compaction test). Finally, the specimens were saturated,  
423 consolidated and compressed under undrained strain-controlled conditions. Since the Normal  
424 Compression Line of the pure Kaolin matrix was not provided by the authors, an alternative  
425 one done by Atkinson *et al.* (1987) was used for calibrating parameters, since it has  
426 approximately the same Atterberg limits as the commercial Kaolin clay used by Jafari and  
427 Shafiee (2004).

#### 428 *Kaolin clay-sand mixtures (Wood and Kumar, 2000)*

429 The mixture tested by Wood and Kumar (2000) consists of Kaolin matrix and coarse uniform  
430 sand inclusions. The Kaolin clay has a liquid limit and plastic limit of 80% and 39%,  
431 respectively. Most of the soil particles (95%) of soil matrix are finer than 0.002 mm. The size  
432 of sand particles is more or less uniform around 2.0 mm, and the particle shape is sub-angular  
433 to sub-rounded. The maximum and minimum porosity of the coarse sand is 0.50 and 0.37,  
434 respectively. The particle densities are 2620 and 2650 kg/m<sup>3</sup> for kaolin and sand, respectively.  
435 First, water was added to the dry Kaolin powder to reach a desired water content of 120%.  
436 The slurry was mixed homogeneously, and then coarse sand particles were added. Finally, the  
437 sample was pre-consolidated in a consolidometer, followed by a further pre-consolidation  
438 (400 kPa), (reloading) and shearing in a triaxial cell. Three different consolidation ratios of  
439 the mixture were considered: OCR = 1.0, 1.3, and 4.0.

#### 440 *Model predictions on the three gap-graded mixtures*

441 It is assumed that the mixtures have a uniform initial stress of 10 kPa, followed by an  
442 isotropic compression and a further triaxial shearing process. The model parameters for the  
443 gap-graded soils are given in Table 2, which were determined from the procedure summarized  
444 in Section 5. Note that the values of the structure parameter  $\zeta$  for natural gap-graded soils are

445 0.57/0.60/0.63 for HTP/HTP10/HTP40, respectively. The predictions based on our model are  
446 compared against tests data for all three gap-graded mixtures are shown in Figs. 5-16.

447 Figs. 5-10 present a comparison of our model predictions with the experimental  
448 observations made by Ruggeri *et al.* (2016) on natural gap-graded soils. It is seen that the  
449 proposed model can well reproduce the effect of coarse fraction on the mechanical responses  
450 and volumetric deformation behavior of the tested natural gap-graded soils. The experimental  
451 results of the kaolin-clay and kaolin-gravel mixtures obtained by Jafari and Shafiee (2004)  
452 and the numerical simulations using the proposed model are presented in Figs. 11-12. The  
453 shear strength is moderately underestimated by the proposed model at the confining stress of  
454 100 kPa, which may be attributed by the overconsolidation due to the compaction during  
455 sample preparation. The simulations are consistent with the experimental data for volume  
456 fractions of 20% and 40%, However, a difference arises between the experimental data and  
457 the simulation curves for a high fraction of aggregates (60%). This may be due to the  
458 following two reasons: (1) large pores may exist in the soil matrix or the interface between the  
459 two phases; (2) an associated flow rule is assumed for the gap-graded soils, which may be not  
460 applicable in case where the inter-granular skeleton of aggregates controls the deformation  
461 process.

462 Different values of shape parameter are calibrated from the data of drained triaxial tests  
463 ( $k_m=1.6$ ) and from the undrained triaxial tests ( $k_m=2.0$ ) (Wood and Kumar, 2000). If  $k_m=2.0$  is  
464 adopted for all numerical simulations, comparison of predictions with the experimental data  
465 of kaolin clay-sand mixtures (Wood and Kumar, 2000) are shown in Figs. 13-16. Noticeably,  
466 the proposed model cannot well capture the overall shear stress and overall volumetric  
467 deformation in drained triaxial tests. This may be due to the fact that the shape parameter is  
468 calibrated from the undrained tests. For a better fitting of the experimental data of the pure  
469 Kaolin clay,  $k_m=1.6$  is adopted for a further comparison for the drained case (solid lines Figs.

13 and 14). Evidently, it is seen that the difference of overall shear stress between the simulations and experimental data significantly decreases in the drain case, whereas the overall volumetric strain remains underestimated. This can be interpreted by the deficiency of the generalized modified Cam clay model which underestimates the volumetric strain of the pure kaolin matrix in drained triaxial tests.

## **7. Conclusions**

An elastoplastic constitutive model has been proposed for gap-graded soils based on mixture theory and a volume-average homogenization scheme. Validation of the model against experimental data has been presented. A summary of the features of present model and conclusions are presented below:

(1) The effect of inter-granular skeleton is considered by incorporating a structure parameter which evolves with the volume fraction of the rock aggregates. A small value of the structure parameter yields a negligible increase of the overall shear strength. However, a higher value of structure parameter can simulate a continuous increase of overall shear strength with increasing rock fraction.

(2) Simulation of the proposed model provides insights into the mechanisms governing the evolution of inter-granular skeleton. An isotropic loading may induce a nonuniform stress distribution in gap-graded soils, where the stress in the soil matrix is lower than that of the rock aggregates. The stress paths of the phases are almost parallel during subsequent triaxial compression loading.

(3) Compared with the generalized Modified Cam clay model, the proposed model has only one additional structure parameter, which can be estimated by trial and error using the data of a triaxial compression test on gap-graded soils with a prescribed fraction of rock aggregates. The other model parameters can be calibrated from an oedometer (or isotropic compression)

495 test and a triaxial test on the pure soil matrix.

496 Test data of three different gap-graded soils from the literature are compared with the  
497 predictions of the proposed model, revealing that the proposed model can well reproduce the  
498 stress strain relationship of gap-graded soils. However, it is noteworthy that the proposed  
499 model has been targeted for binary gap-graded soils, based on the following hypotheses: The  
500 inter-granular space is fully filled with fine soil matrix. In intense weathering areas, fine  
501 content is usually beyond the minimum porosity of the pure aggregates, and no macro-pores  
502 prevail between the large aggregates (Kavvasdas *et al.* 1996; Vallejo and Mawby, 2000; Zhou  
503 *et al.*, 2017). Therefore, it can be simplified as binary mixtures and be treated using mixture  
504 theory. However, The macro-pores would arise in case that the volume fraction of matrix is  
505 less than the minimum porosity of pure aggregates, and the decreasing of fine fraction leave  
506 increasing macro-pores between the coarse aggregates. This kind of soils cannot be properly  
507 modelled within the mixture theory, and further efforts need to be devoted to address this  
508 issue.

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Table 1. Model parameters for benchmark analysis of the proposed model

Parameters	$M_m$	$N_m$	$\lambda_m$	$\kappa_m$	$\mu_m$	$k_m$	$\xi$
Value	1.4	0.6	0.05	0.01	0.22	2.0	0.00/0.05/0.20/0.35

Table 2. Model parameters for validation of the proposed model

Parameters	Natural gap-graded soils	Kaolin-gravel mixtures	Kaolin-sand mixtures
$M_m$	0.97	0.98	0.80
$N_m$	0.817	1.269	1.35
$\lambda_m$	0.056	0.089	0.085
$\kappa_m$	0.019	0.030	0.020
$\mu_m$	0.35	0.23	0.30
$k_m$	2.0	1.1	2.0 (1.6)
$\zeta$	0.57/0.60/0.63	0.07	0.05

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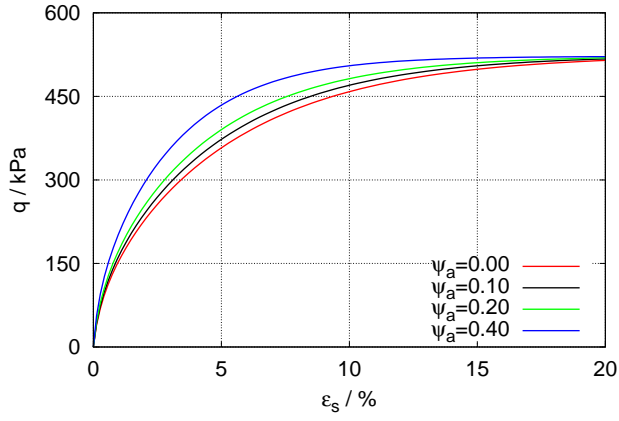
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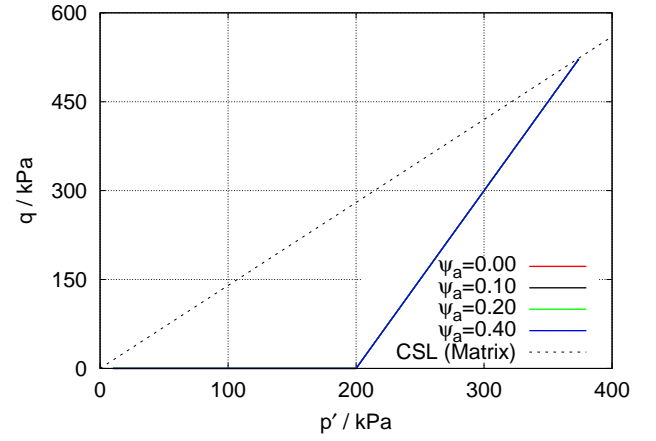
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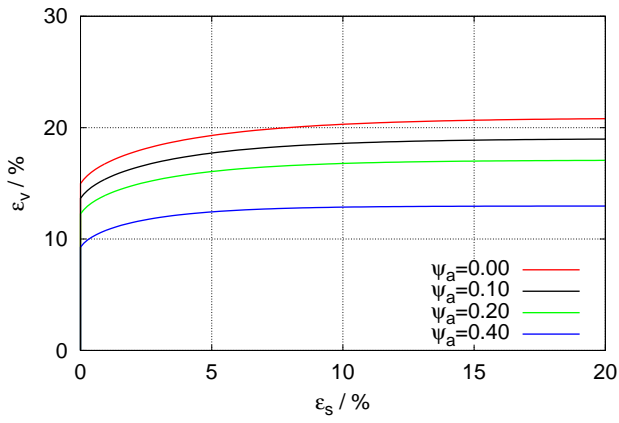
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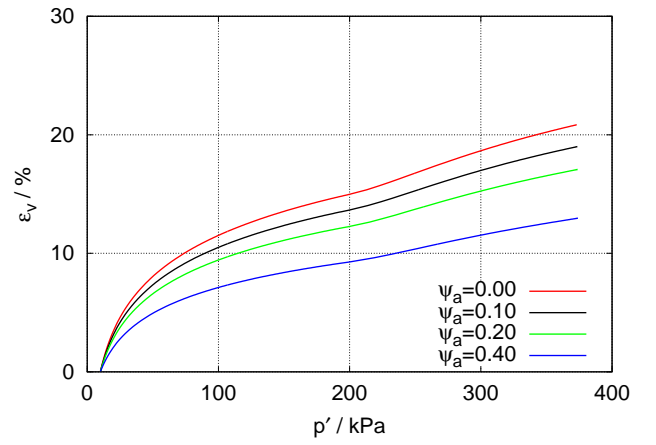
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(b)  $p'$ - $q$

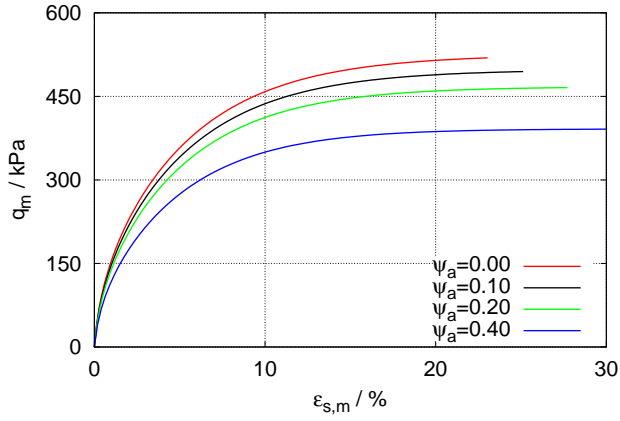


(c)  $\varepsilon_s$ - $\varepsilon_v$

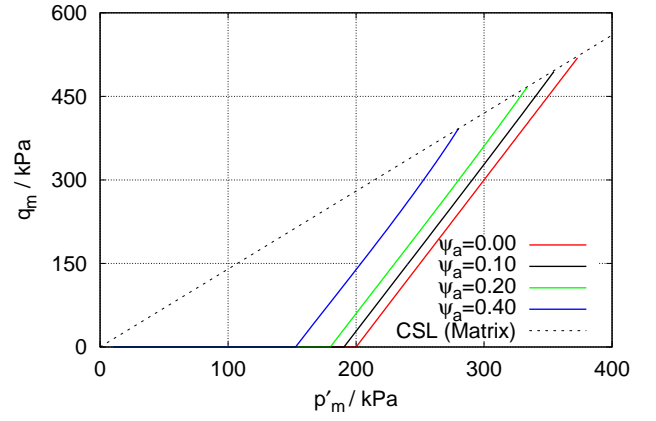


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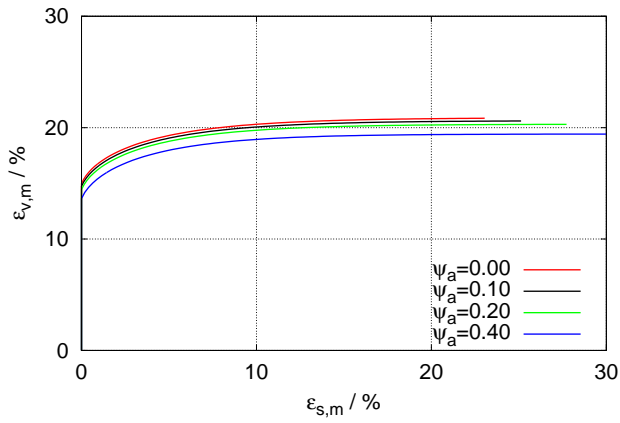
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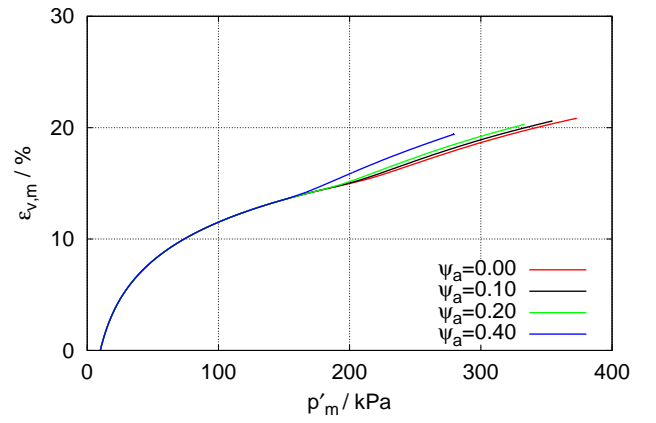
(a)  $\varepsilon_{s,m}-q_m$



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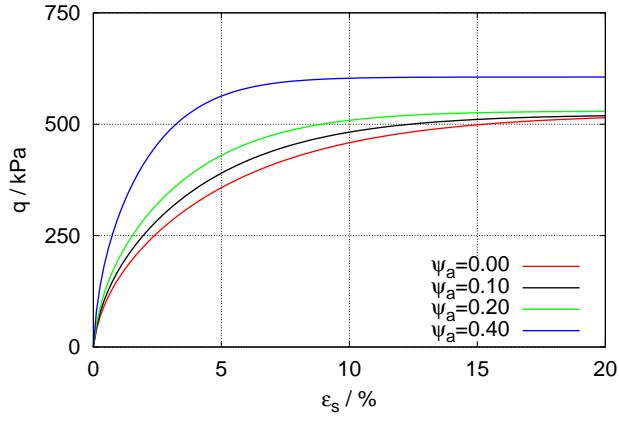


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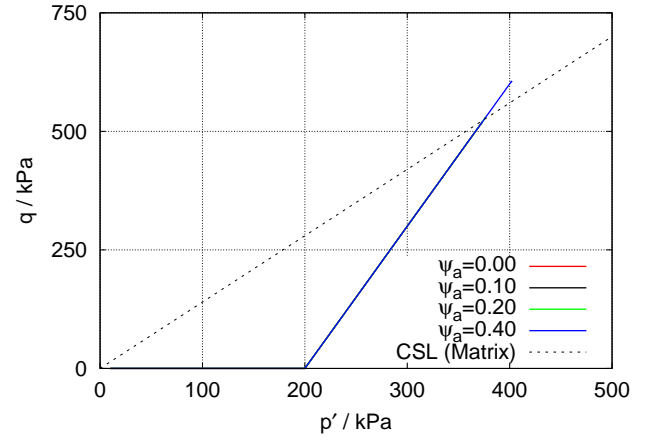


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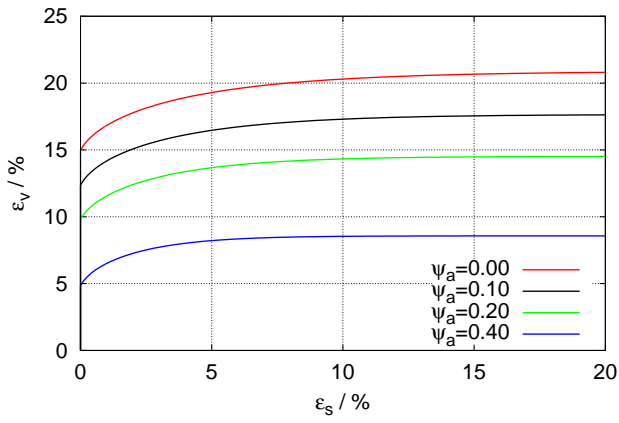
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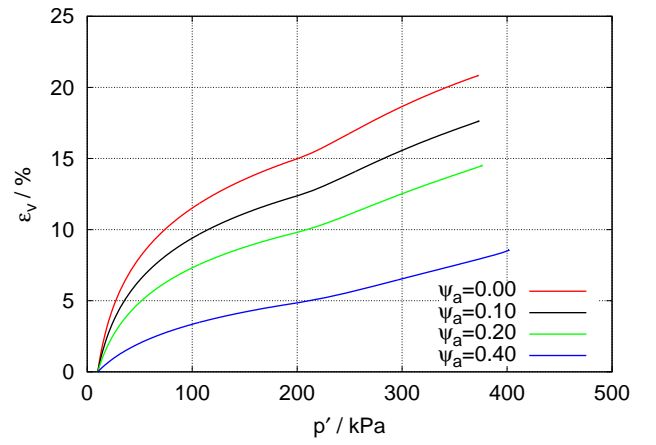
(a)  $\varepsilon_s$ - $q$



(b)  $p'$ - $q$

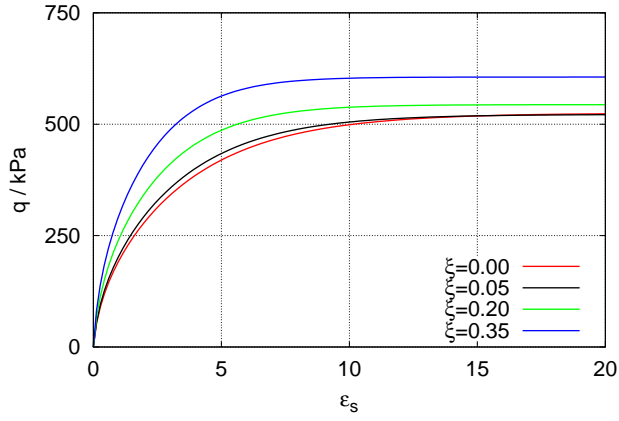


(c)  $\varepsilon_s$ - $\varepsilon_v$

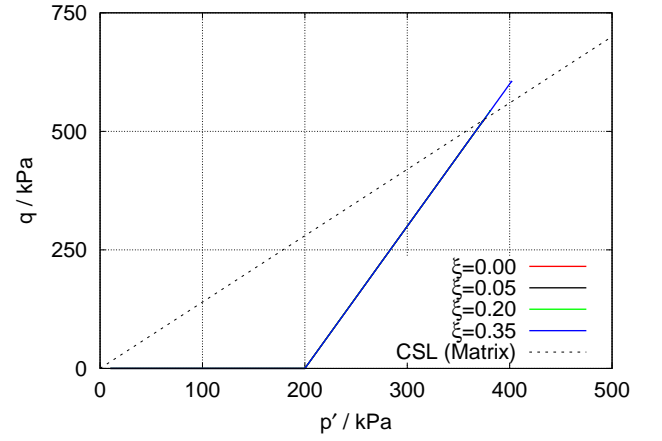


(d)  $p'$ - $\varepsilon_v$

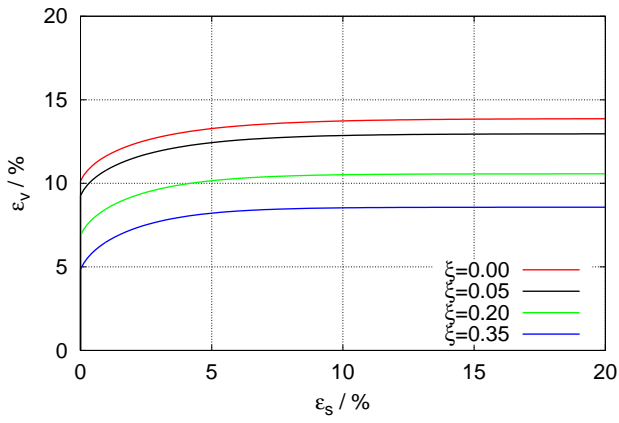
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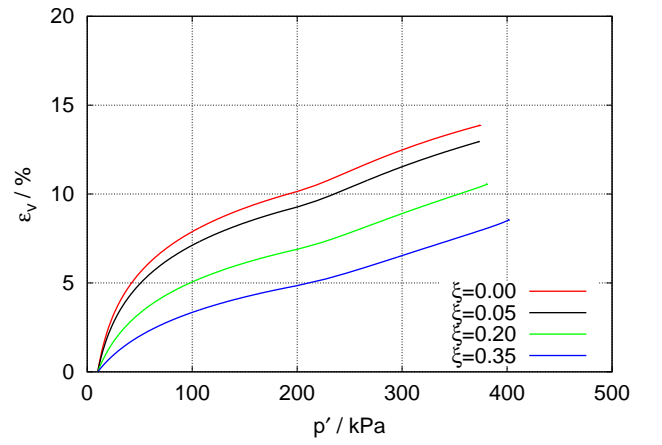
(a)  $\varepsilon_s$ - $q$



(b)  $p'$ - $q$

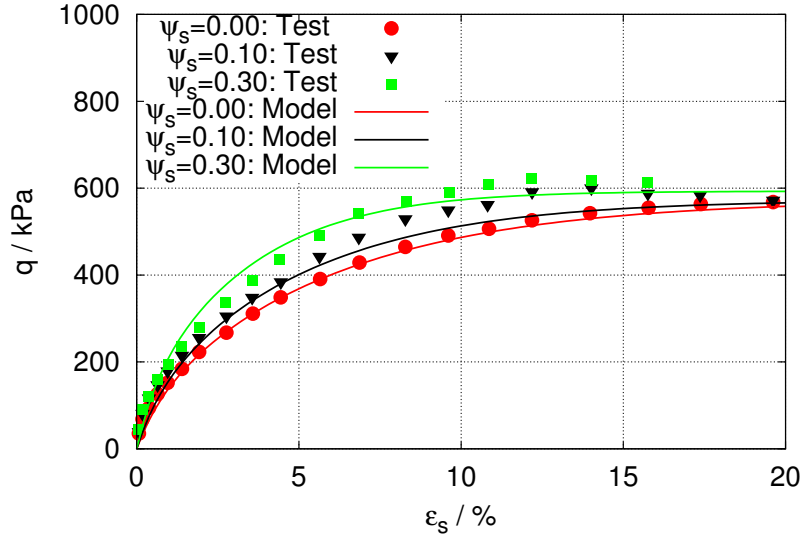


(c)  $\varepsilon_s$ - $\varepsilon_v$

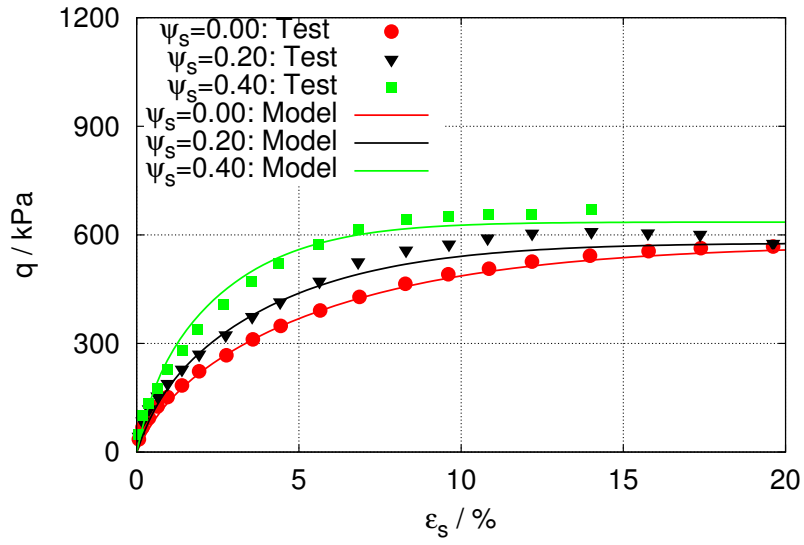


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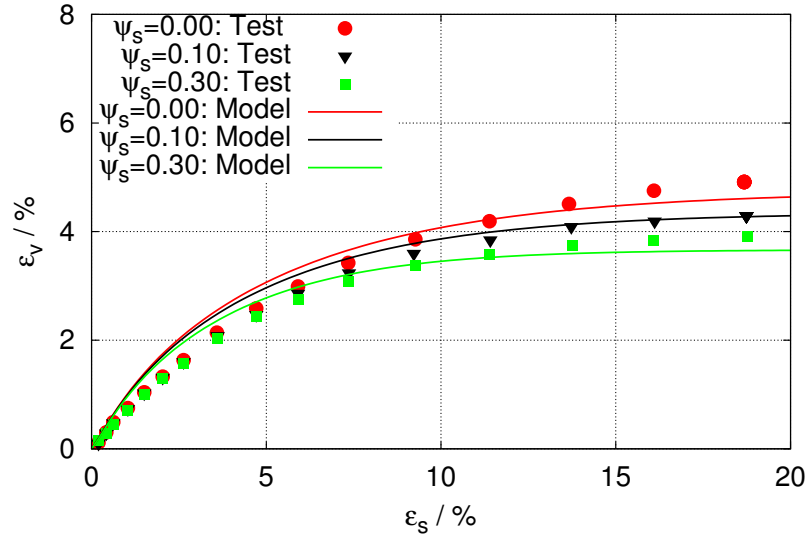
(a)  $\psi_a = 0.10, 0.30$



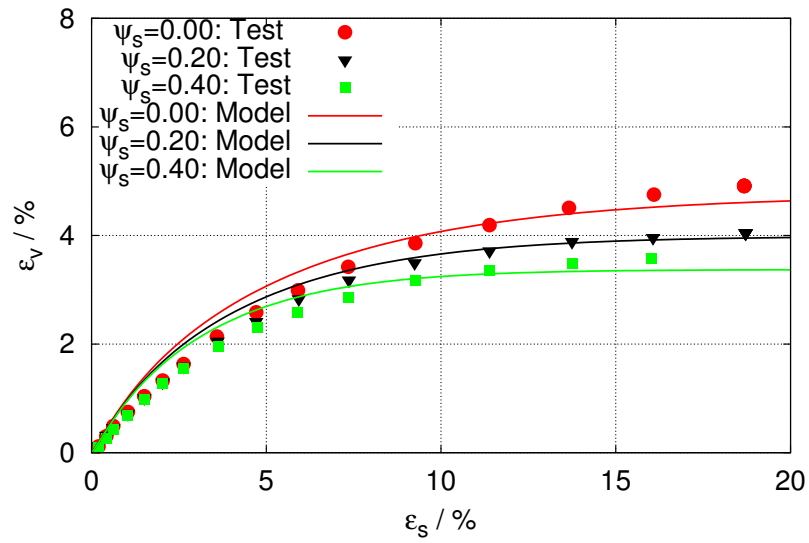
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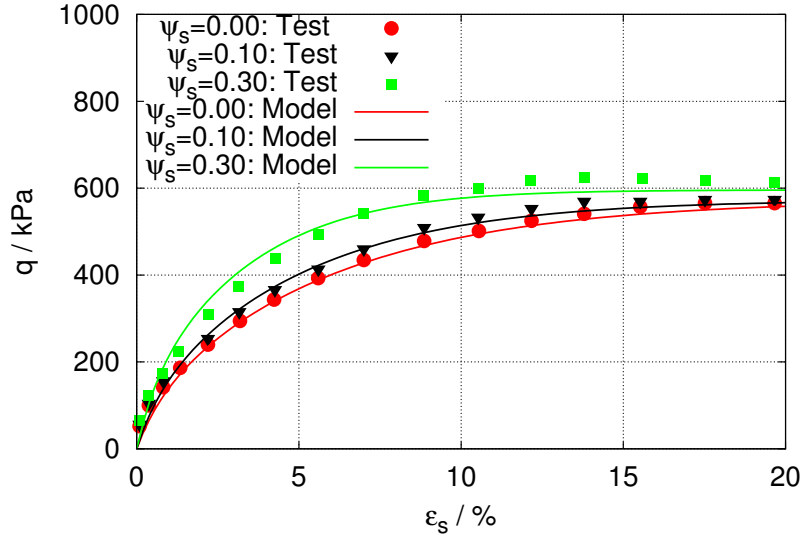


(a)  $\psi_a=0.10, 0.30$

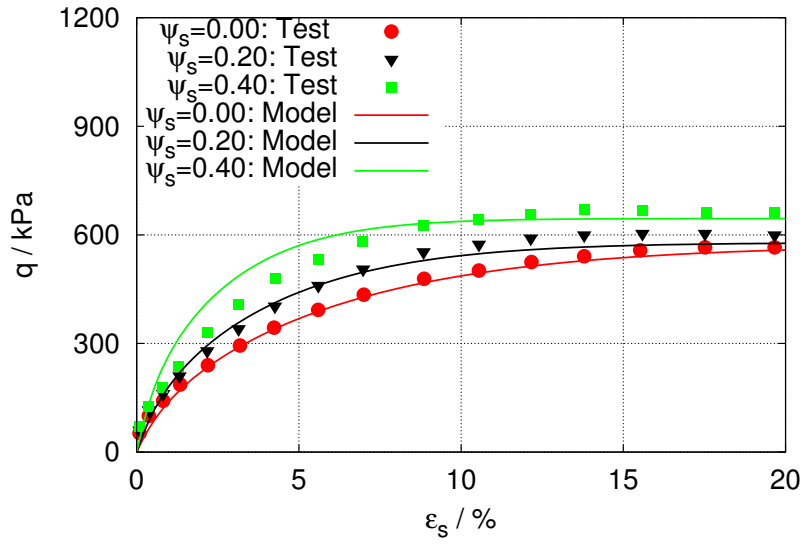


(b)  $\psi_a=0.20, 0.40$

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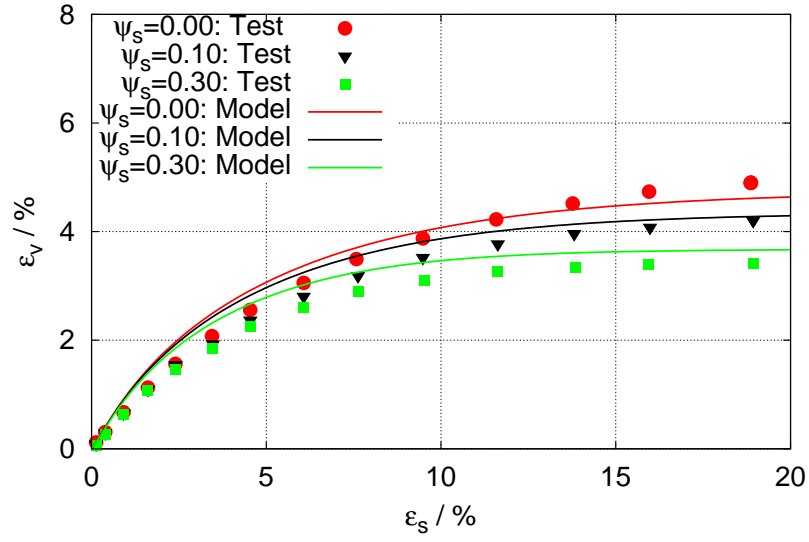


(a)  $\psi_a=0.10, 0.30$

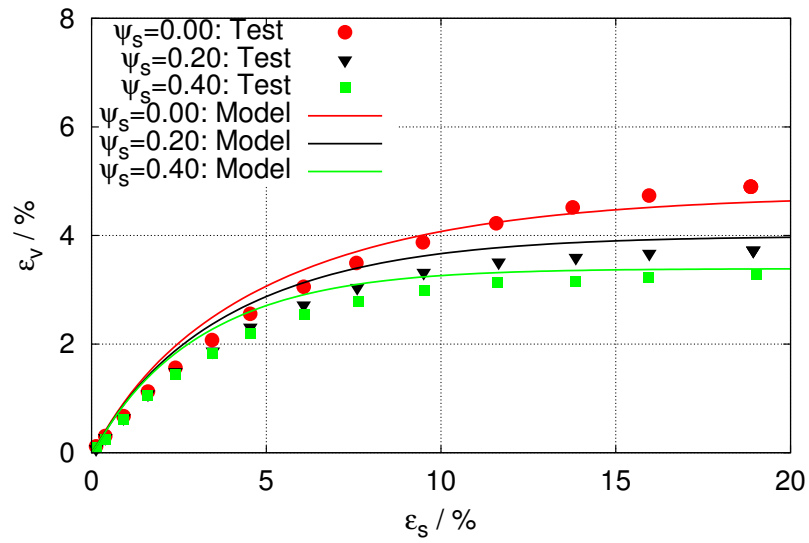


(b)  $\psi_a=0.20, 0.40$

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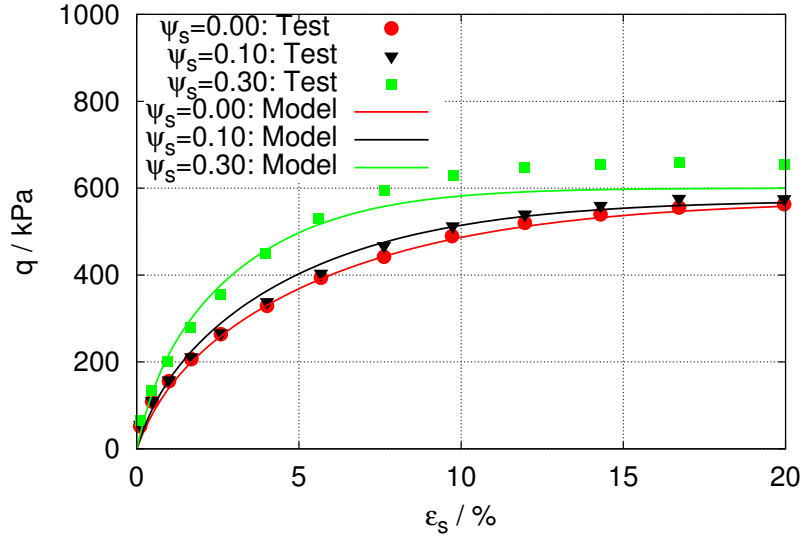


(a)  $\psi_a = 0.10, 0.30$

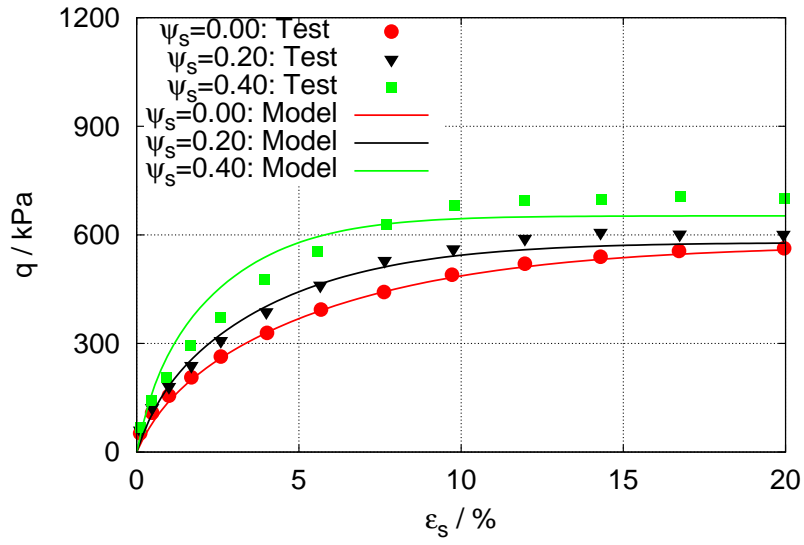


(b)  $\psi_a = 0.20, 0.40$

Figure 8: Experimental volumetric strain and numerical simulations of HTP10 series (Natural gap-graded soils)

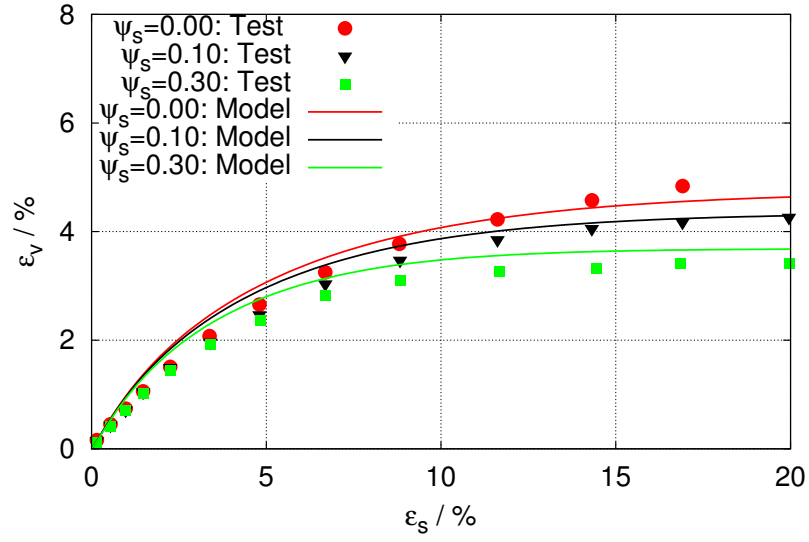


(a)  $\psi_a = 0.10, 0.30$

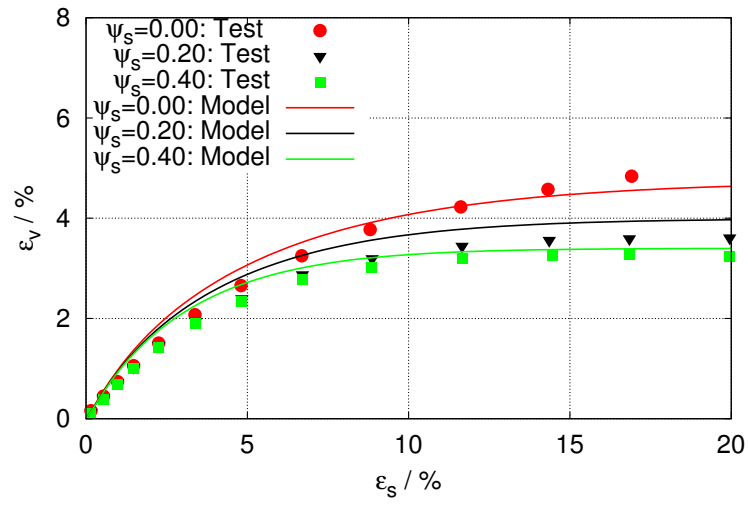


(b)  $\psi_a = 0.20, 0.40$

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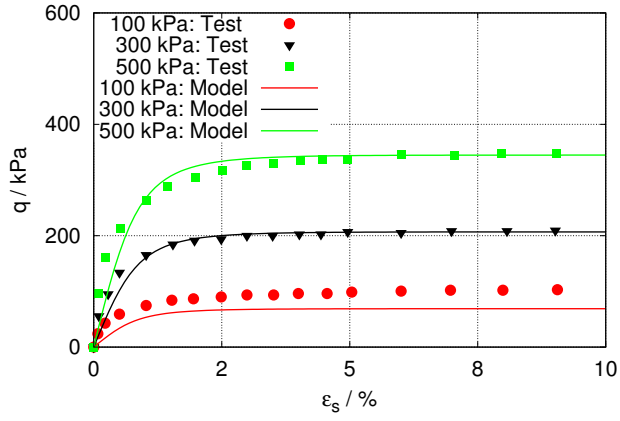


(a)  $\psi_a = 0.10, 0.30$

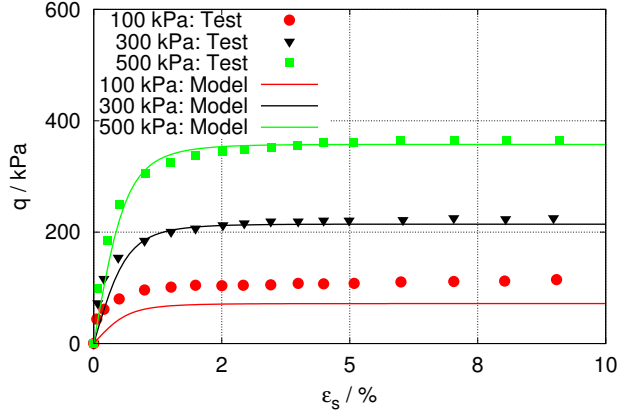


(b)  $\psi_a = 0.20, 0.40$

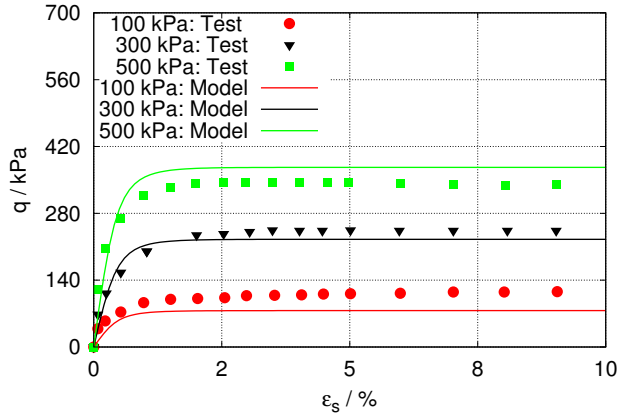
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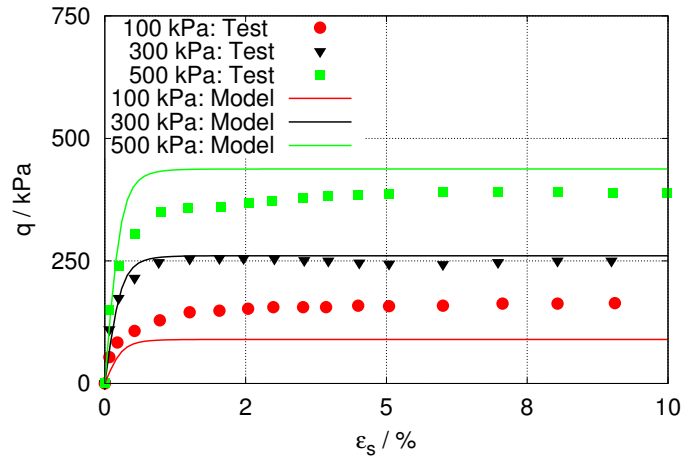
(a)  $\phi_a=0.00$



(b)  $\phi_a=0.20$

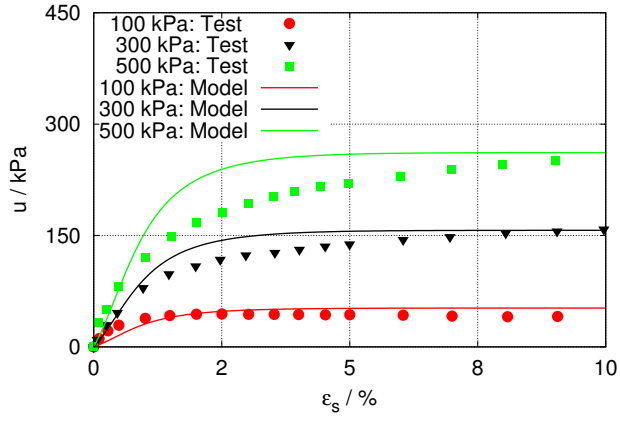


(c)  $\phi_a=0.40$

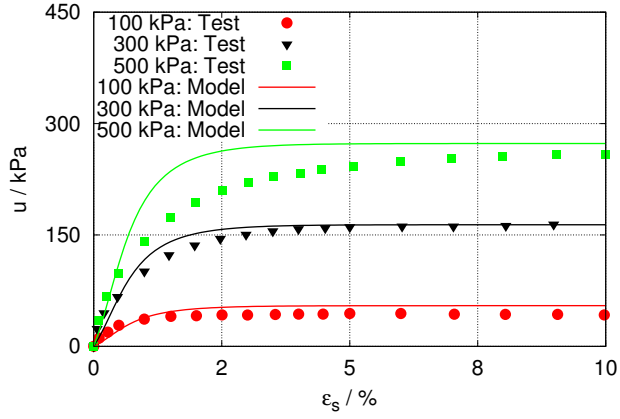


(d)  $\phi_a=0.60$

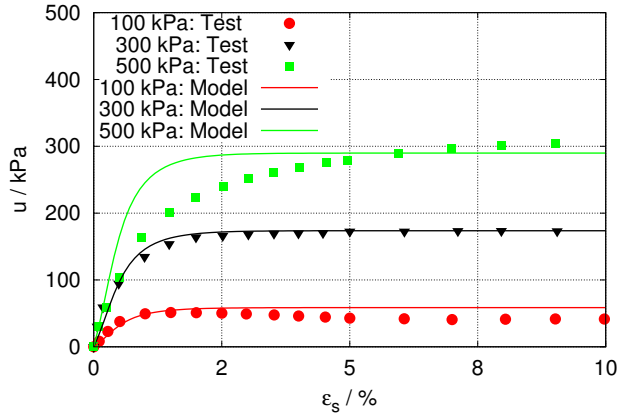
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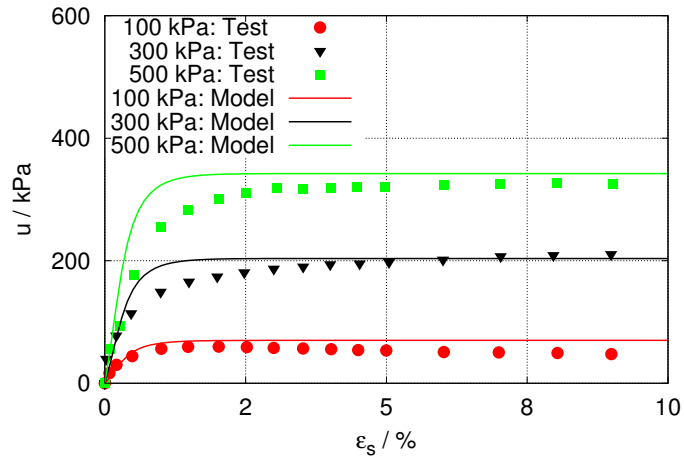
(a)  $\phi_a=0.00$



(b)  $\phi_a=0.20$

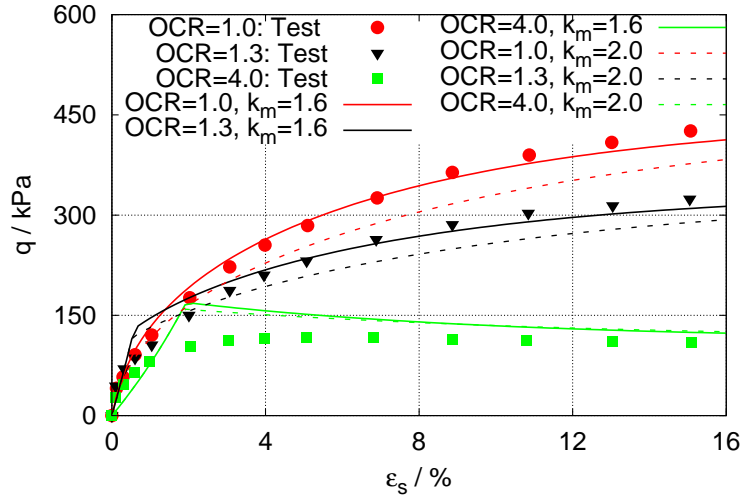


(c)  $\phi_a=0.40$

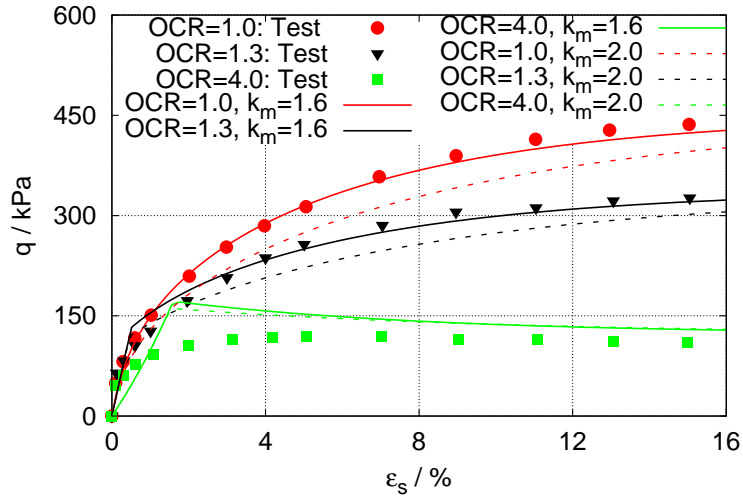


(d)  $\phi_a=0.60$

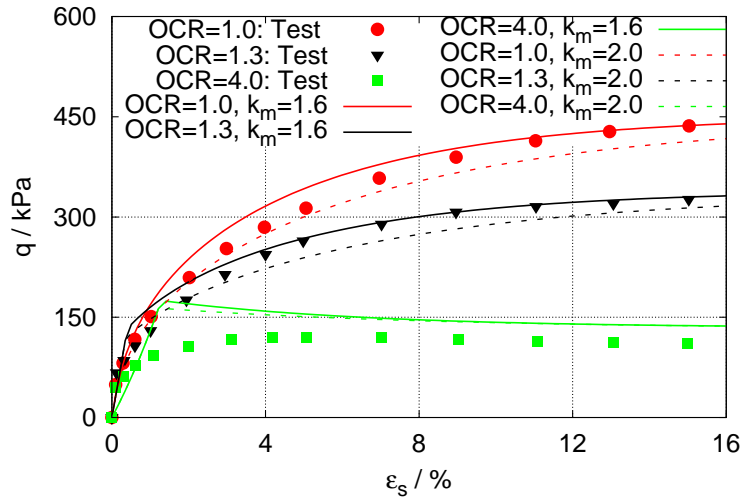
Figure 12: Experimental data of excess pore water pressure dissipation and numerical simulations (Kaolin clay-gravel mixtures)



(a)  $\psi_a=0.00$



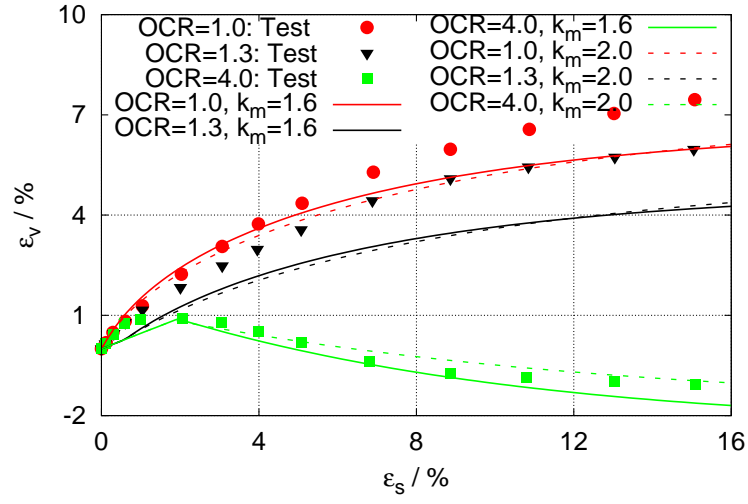
(b)  $\psi_a=0.30$



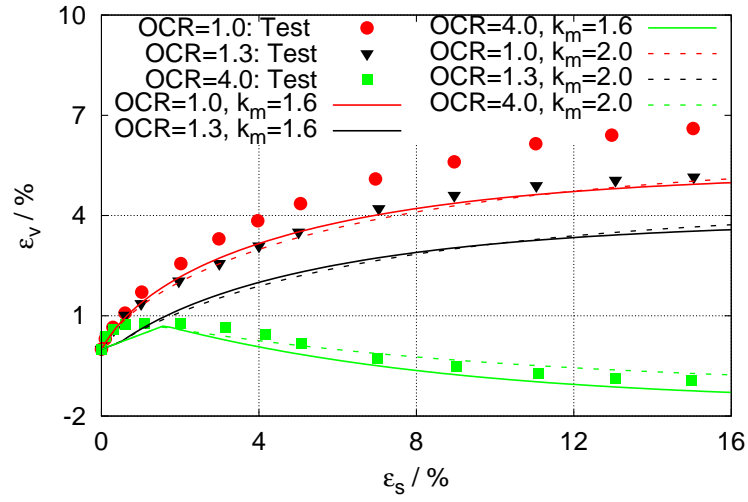
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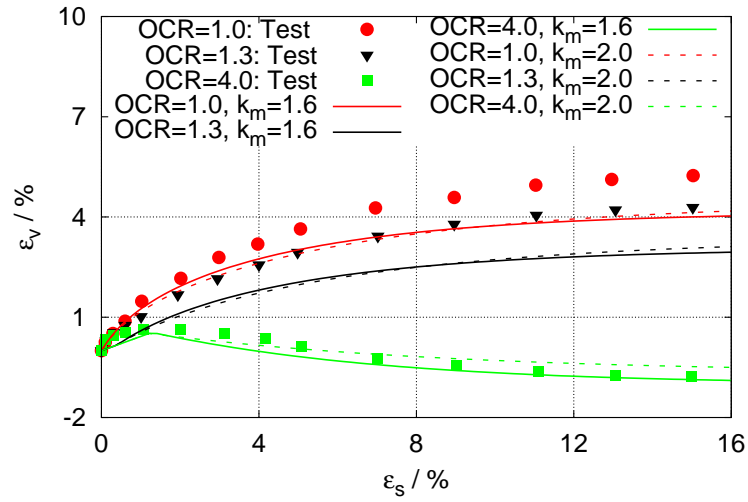




(a)  $\psi_a=0.00$

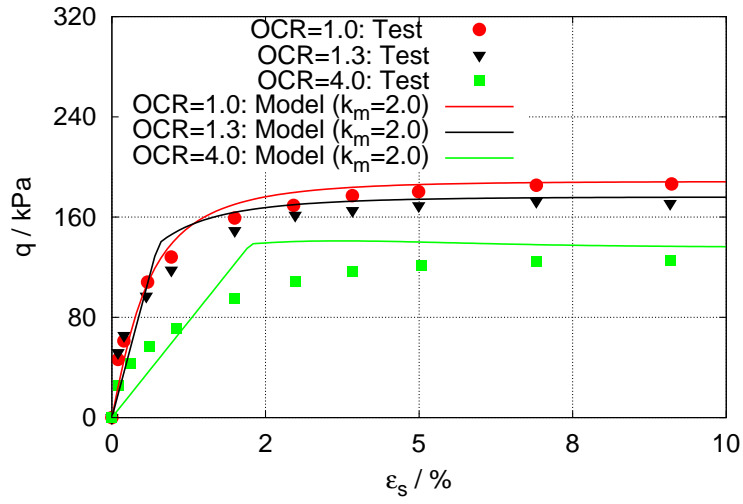


(b)  $\psi_a=0.30$

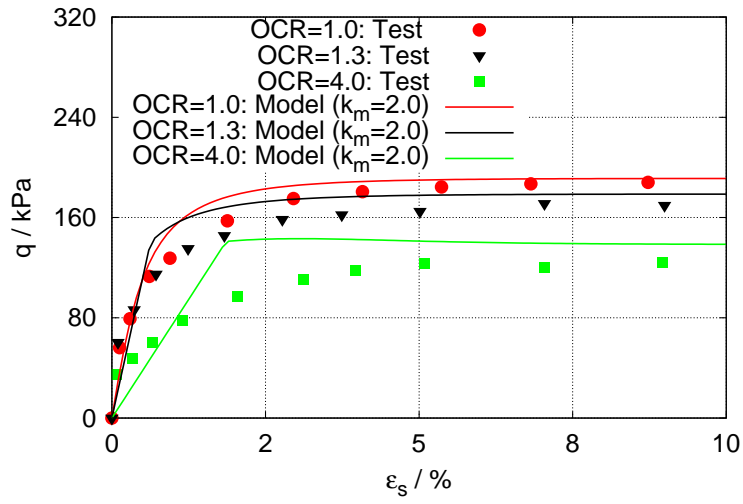


(c)  $\psi_a=0.50$

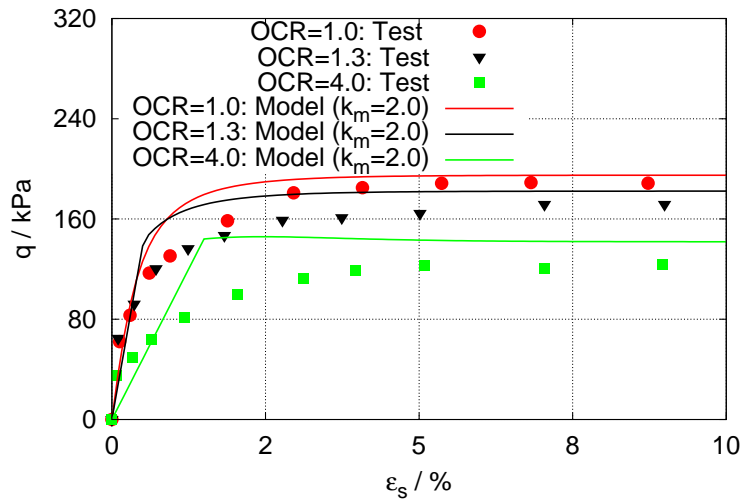
Figure 14: Experimental volumetric strain and numerical simulations of drained triaxial test (Kaolin clay-sand mixtures)



(a)  $\psi_a=0.00$

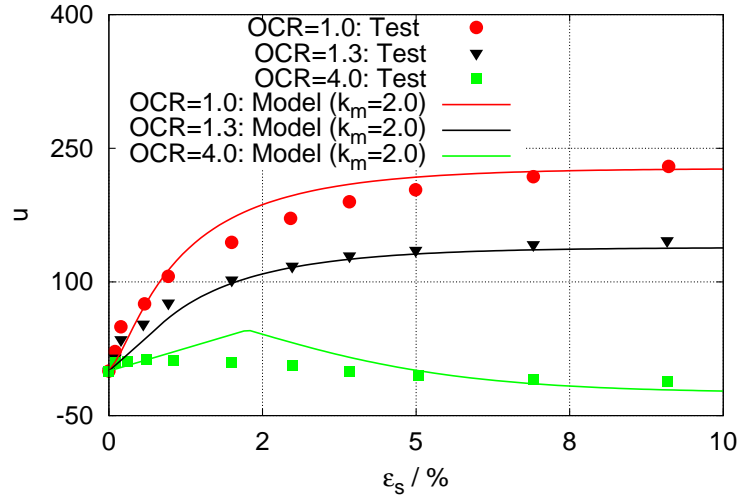


(b)  $\psi_a=0.30$

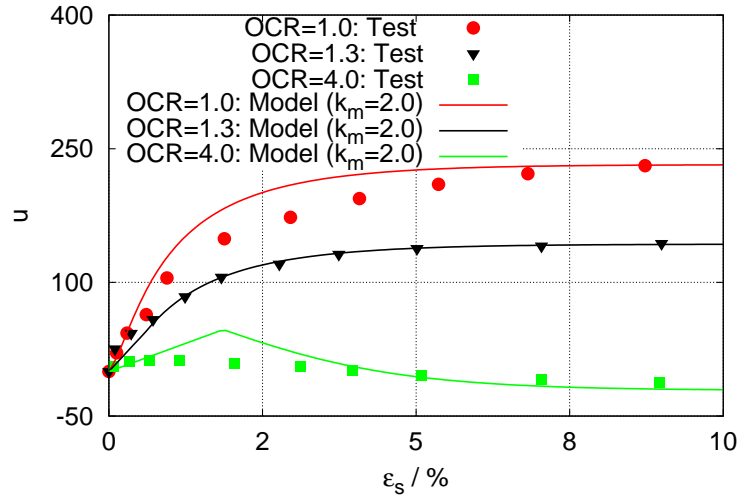


(c)  $\psi_a=0.50$

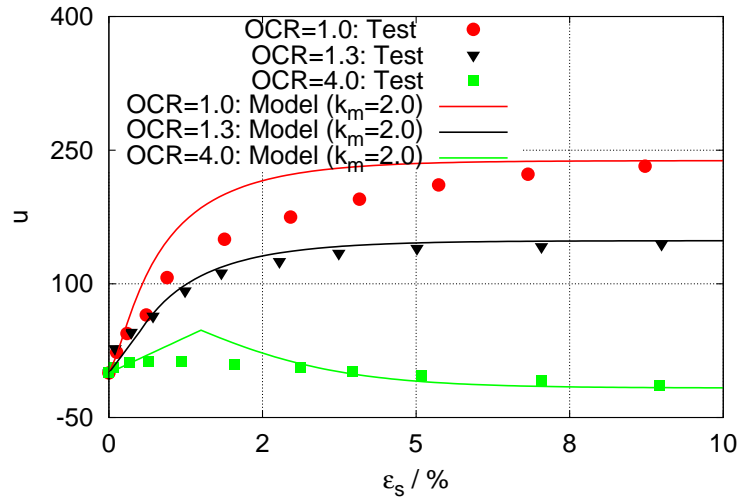
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(a)  $\psi_a=0.00$



(b)  $\psi_a=0.30$



(c)  $\psi_a=0.50$

Figure 16: Experimental data of excess pore water pressure dissipation and numerical simulations of undrained triaxial test (Kaolin clay-sand mixtures)